









संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती मखाठी विमाग

भाग अखिल भाषतीय महाती साहित्य महामंडळ बंबुमार्ग गर्वति

# चर्चासत्र मराठी : अभिजात, बोली आणि उपयोजित

दि. २४ सप्टेंबर २०१६

उद्यातन स्वार्थम अध्याः मा. घं, कुल्तीयन चांदेवस, कुश्तनुम, संत ना। कः सा क्षाव्य भोड, अध्याः, म. रा. नाहित् स्वी वी नाम्याय चीतायतु, पूर्वप्रधाः, ज. मा.स.

संत मातमे व









Government of India Department of Atomic Energy (DAE) Board of Research in Nuclear Sciences (BRNS)

Dr. S.Dutta Programme Officer (RTAC) BRNS Secretariat, 1st Floor, CC, BARC, Trombay, Mumbai-400085. Phone: +91-2225595331 Email: dsubrata@barc.gov.in

Date: 07/11/2017

No: 35/14/30/2016-BRNS/35259

#### OFFICE MEMORANDUM

Sub: **Renewal-cum-additional** R/P entitled "Gamma irradiated enhancement of Artemisinin production from A. annua L. and process optimization" under Dr. Anita Surendra Patil, Professor, Department of Biotechnology, Sant Gadge BabaAmravati University, Amravati-444 602, MS bearing sanction 35/14/30/2016-BRNS with RTAC, BRNS.

In continuation of this Department's OM of even number dated 09/08/2016 on the recommendations of the Board of Research in Nuclear Sciences (BRNS), I am pleased to convey the administrative approval of the President of India for the continuance of captioned project during **2017-2018** and sanction to incur an expenditure of **Rs. 5,59,875/- (Rupees five lakh fifty nine thousand eight hundred seventy five only)** as detailed below:

Item of expenditure	Year 2
	(2017-2018)
Staff Salary - JRF (1)	300000
Consumables	150000
Travel - PI	15000
Contingencies *	* 60000
Overheads	34875
Total(INR)	559875

\* This includes additional grant of ₹.30,000/-.

2. The grant remaining unspent under the Heads "Equipments", "Consumables" and "Overheads" for the year 2016-2017 are only allowed to be carried forward to the current financial year 2017-2018. Unspent balance under other heads will be adjusted while releasing the subsequent grant.

3. The expenditure involved is debitable to the following Head of Account: 3401 00 004 27 02 31

4. This issues with the concurrence of Scientific Secretary, BRNS and IFA.

Subrala Sulla

Dr. S.Dutta

Pay & Accounts Officer, DAE, Mumbai-400 001

Copy forwarded to:

- 1. Director of Audit, Scientific Department, AEAP,OYC,CSM Marg, Mumbai-400 001.
- 2. Joint Secretary (R&D), DAE, Anushakti Bhavan, CSM Marg, Mumbai-400 001.
- 3. Registrar, Sant Gadge Baba Amravati University, Amravati-444 602, MS.

4. Principal Investigator(PI): Dr. Anita Surendra Patil, Professor, Department of Biotechnology, Sant Gadge Baba Amravati University, Amravati-444 602, MS.

A. The actual amount of grant-in-aid that would be released will be limited to the actual expenditure up to 31st March 2017 subject to the maximum recommended by the Board.

B. In accordance with Rule 30 of the General Financial Rules 2005, this sanction will lapse if no payment in whole or part is made during a period of twelve months from the date of issue of this sanction letter.

C. The PI is authorized to incur the expenditures towards the general administrative requirements of the project and towards fees payable to Auditor/Chartered Accountant for audit of accounts of the project, from the funds sanctioned under "overhead" for the project.

D. The other Terms and Conditions governing the payment of the grant-in-aid have already been sent to you while sanctioning the grant for the first year of the project.

E. Grant for the second year will be released in FULL (unspent balance of previous year and Interest earned will be adjusted) on receipt of the bill IN DUPLICATE in prescribed FORM-II along with the following documents. The data for these documents are required to be filled online and a printout of the same should be forwarded to the undersigned after obtaining original signatures of the concerned officials.

(a) Utilization Certificate (FORM-III) for the preceding year.

(b) Statement of Accounts signed by Internal Auditor/ Accountant of the Institution/ University (FORM-IV) for the preceding year. Interest earned in previous year should be reflected in the Statement of Accounts.

(c) Copy of appointment order and joining report of the staff appointed for the project along with Minutes of the Selection Committee.

(d) An Inventory of euipment (FORM-V).

F. PLEASE NOTE THAT CLAIM(S) SHALL BE SUBMITTED TO THIS DEPARTMENT WELL IN ADVANCE BEFORE THE END OF THE FINANCIAL YEAR IN WHICH THE CLAIM IS DUE. OTHERWISE, THE SANCTIONED GRANT WILL BE LAPSE.

G. Renewal/ Extension Application: Principal Investigator (PI) is required to upload by January 15 a pdf copy of duly signed renewal/ extension application in the prescribed Form-PRA after logging into his/her account at www.daebrns.gov.in . Template for Form-PRAcan be downloaded from website www.daebrns.gov.in . All applications received shall be examined by experts from the field and PIs may be invited to a Technical Programme Discussion Meeting (TPDM). Renewal of the project will be based on the recommendations of the TPDM, Advisory Committee and the Board.

- 5. AAO (Cheque), DAE, Anushakti Bhavan, CSM Marg, Mumbai-400 001.
- 6. Co-Investigator (CI): Dr Ramesh Satdive, Nuclear Agriculture and Biotechnology, Bhabha Atomic Research Centre, Trombay, Mumbai- 400085, Email : rameshsatdive@gmail.com, Mobile :9769735571 .
- 7. Member Secretary (RTAC) : Dr Ashutosh Dash, adash@barc.gov.in; Ph:+91-2225595372 .
- 8. Principal Collaborator(PC): Dr. D. P. Fulzele, Nuclear Agriculture and Biotechnology, Bhabha Atomic Research Centre, Trombay, Mumbai- 400085, Email : dfulzele@hotmail.com, Mobile :9869327940.

Subrala Dulla

Dr. S.Dutta

#### Note:

- 1. Please quote the Sanction Number 35/14/30/2016-BRNS in all your correspondence with BRNS.
- 2. Kindly upload renewal/extension forms and update the Statement of Accounts, details of Bank Account, Equipment and staff appointed, after clicking view application the menu available on the top left corner needs to be populated by the PI. Henceforth the office memorandum (OM) or all other letters would be downloadable from PI's account. Hard copies would no more be posted. All the forms mentioned in the sanction letter and the terms and conditions are also available on the website. In case of any difficulty, please call up the undersigned.
- 3. Please note that as per the government orders under **Direct Benefit Transfer (DBT)scheme, the staff salary** has to be transferred to his/her bank account. Accordingly, Aadhar Number(UID) of the appointed staff, Bank Account details and the Mobile number linked to the bank account should be obtained and it should be intimated to this office.

Final Report and Grant Utilization Certificate of MRP

*In Vitro* Production And Process Optimization Of Camptothecin (An Anticancer Drug) From *Nothapodytes nimmoniana* Graham: A Rare Endangered Medicinal Plant From Western Ghats Of Maharashtra (F.No.42-212/2013 (SR)

SUBMITTED TO

### UNIVERSITY GRANTS COMMISSION, BAHADUR SHAH ZAFER MARG

NEW DELHI - 110 002



SUBMITTED BY

DR. ANITA PATIL(PI) Professor Department of Biotechnology SGB Amravati University, Amravati (M.S) DR. SURENDRA PATIL (Co- PI) Professor College Of Horticulture Dr. P. D. K.V Krishi Nagar, Akola (M.S)

Scanned with CamScanner

### UNIVERSITY GRANTS COMMISSION **BAHADURSHAH ZAFAR MARG NEW DELHI-110002**

# STATEMENT OF EXPENDITURE IN RESPECT TO MAJOR RESEARCH PROJECT

- 1 Name of Principal Investigator : Dr. Anita Surendra Patil 2 Department of Principal Investigator : Department Of Biotechnology University/College Sant Gadge Baba Amravati University, Amravati (M.S) 444 602
- 3 UGC Approval Letter No. & Date : F. No.42-212/2013 (SR) dated 22 March, 2013
- Title of Research Project : In Vitro Production And Process Optimization Of 4 Camptothecin (An Anticancer Drug) From Nothapodytes Nimmoniana Graham: A Rare And Endangered Medicinal Plant From Western Ghats Of Maharashtra.
- Effective Date of Starting Project 5
- : 26-04-2013

6. a Period of Expenditure : 01/04/16 to 31/03/17

Details of Expenditure b

S.	Item	Total	Amount Approved (De) Tetho			
No		1 otal	Amount Approved (KS)		lotal Grant	Expenditure
NO		Amount	Ist Creant 2nd Creat		Received	Incurred
		Approved	1st Grant	2 Grant		$(\mathbf{P}_{\mathbf{S}})$
1	Books and Journals	-	Nil		-	(13)
2	Equipment	3,50,000=0	3.50.000=0	-	3 50 000-0	2 00 700 0
	(Interest 49,700)*				5,50,000-0	3,99,700=0
3	Contingency	30,000=0	15,000=0	12,000=0	27,000=0	27.000 0
4	Field Work/Travel	30,000=0	15,000-0	12,000-0	27,000=0	27,000=0
5	Llining Comilese	20,000 0	15,000-0	12,000=0	27,000=0	27,000=0
5	rining Services	30,000=0	15,000=0	12,000=0	27,000=0	27.000=0
6	Chemicals and	1,75,000=0	87,500=0	70,000=0	1.57.500=0	1.52.057-0
	Glassware				1,57,500-0	1,55,957=0
7	Overhead	76,300=0	76.300=0	-	76 200-0	
8	Any other items	56 440=0	56,440-0	-	70,300=0	76,300=0
	(Please specify)	50,440 0	50,440-0	-	56,440=0	56,440=0
	(Trease speerry)					
	Interest*					
9	Project Fellow	5,28,000=0	2.64.000=0	1 89 600=0	1 52 600-0	1.55 1.10 0
- manufold of an average of	Salary		_,.,	1,07,000-0	4,33,000=0	4,57,143=0
	Total (Rs)	12,19,300=00	8,79,240=00	2.95,600=0	11 74 840-0	11 (9 100 0
	# A manual of ! !			=,= 2,500 0	11,74,040-0	11,08,100=0

Amount of interest Rs 49,700/- is incurred on Equipment grant

#### c. Staff : Project Fellow

Date of Appointment: 26/6/13

S.no	Expenditure Incurred	From to	Amount	Expenditure
			Approved (Rs)	Incurred (Rs)
1	Project Fellow	26/6/13 to	4,53,600=00	4,57,143=0
	(1 <sup>st</sup> and 2 <sup>nd</sup> 14000/- p.m. & 3 <sup>rd</sup> year 16,000/-)	13/03/16		

- 1. It is certified that the appointment(s) have been made in accordance with the terms and conditions laid down by the Commission.
- 2. It as a result of checks or audit objective, some irregularly is noticed, laterdate, action will be taken to refund, adjust or regularize the objected amounts.
- 3. Payment @ revised rates shall be made with arrears on the availability of additional funds.

4. It is certified that the expenditure of <u>Rs. 11,68,100/-</u> (Rupee Eleven Lakh, Sixty Eight Thousand and one hundred only) out of the grant received <u>Rs. 11,18,400/-</u> (Rupee Eleven Lakh, eighteen thousand and four hundred only) and interest <u>Rs. 56,440/-</u> (Fifty Six Thousand four hundred fourty only)received from the University Grants Commission under the scheme of support for Major Research Project entitled <u>"In Vitro Production and Process Optimization of Camptothecin (An Anticancer Drug) from Nothapodytes nimmoniana Graham: A Rare and <u>Endangered Medicinal Plant from Western Ghats of Maharashtra</u>"vide UGC letter No. <u>F.No.42-212/2013 (SR) dated 22 March, 2013</u>has been utilized for the purpose for which it was sanctioned in accordance with the terms and conditions laid down and that the balance of **Rs6740=0** remaining grant unutilized will be refund back to the University Grants Commission.</u>

1

SIGNATURE OF PRINCTICALS. Patil Professor Department of Biotechnology INVESTIGAT OR Gadge Baba Amravati University Amravati (M.S.) SIGNATURE OF THE CO-INVESTIGATOR (Dr. Surendra R. Patil) Professor (Horticulture) Department of Fruit Science College of Horticulture Dr. P. D. K. V., Akola (M.S.)

**REGISTRAR/PRINCIPAL** 

3

#### UNIVERSITY GRANTS COMMISSION BAHADUR SHAH ZAFAR MARG NEW DELHI – 110 002

# Final Report of the work done on the Major Research Project

1. Project report No. : Final Report

2. UGC Reference No. : F. No.42-212/2013 (SR)

3. Period of report from : 3 Years from 1.04.2013 to extension to 31.03.2017
4. Title of research project : In Vitro Production and Process Optimization of Camptothecin (An Anticancer Drug) from Nothapodytes nimmoniana Graham: A Rare and Endangered Medicinal Plant from Western Ghats of Maharashtra

5. (a) Name of the Principal Investigator : Dr. Anita Surendra Patil,

Professor Department of Biotechnology Sant Gadge Baba Amravati University, Amravati (M.S) 444 602

(b) Department and University/College where work has progressed:

Department of Biotechnology, Sant Gadge Baba Amravati University, Amravati (M.S) 444 602

6. Effective date of starting of project: 26-04 -2013

7. Grant approved and expenditure incurred during the period of the report:

a. Total amount approved : Rs. 12,19,300/-

b. Total expenditure Rs. : Rs. 11,68,100/-

c. Report of the work done: (Please attach a separate sheet): Attached Appendix-I

Annexure - IV

### UNIVERSITY GRANTS COMMISSION BAHADUR SHAH ZAFAR MARG, NEW DELHI – 110 002

# STATEMENT OF EXPENDITURE INCURRED ON FIELD WORK

# GRANT RECEIVED: Rs. 27,000=00

Name of the Principal Investigator: Dr. Anita Surendra Patil

Name of the Place visited	Duration of the Visit		Mode of Journey	Expenditure Incurred (Rs)
	From	То	· ·	
Melghat Forest Region	28-7-14	30-7-14	Taxi	4,885=0
UGC New Delhi	28-2-14	30-2-14	Train	5,570=0
UGC New Delhi	5-1-16	12-1-16	Train	2,125=0
Mumbai, Alibag,	8-12-16	12-12-16	Train/Taxi	14,420=0
Western Ghats Total				
				27.000=0

Certified that the above expenditure is in accordance with the UGC norms for Major Research Projects.

Dr. Anita S. Patil

SIGNATURE OF PRINCIPALAINNERSTIGATIONBlogy Sant Cadge Baba Amravati University Amravati (M.S.)

SIGNATURE OF THE CO-INVESTIGATOR

(Dr. Surendra R. Patil) Professor (Horticulture) Department of Fruit Science College of Horticulture Dr. P. D. K. V., Akola (M.S.)

REGISTRAR/PRINCIPAL REGIST(Scal) Sant Gadge Baba Anwavati University, Amravati.

Annexure - V

# UNIVERSITY GRANTS COMMISSION

# BAHADUR SHAH ZAFAR MARG

#### **NEW DELHI - 110 002**

#### UTILIZATION CERTIFICATE

It is certified that the expenditure of Rs. 11,68,100=0 (Rupees Eleven lakh sixty eight thousand one hundredonly) out of the grant received Rs. 11,18,400=0 (Rupees Eleven lakh eighteen thousand four hundred only) and interest Rs. 56,440/- (fifty six thousand four hundred fourty only) received from the University Grant Commission under the scheme of support for Major Research Project entitled "In Vitro Production and Process Optimization of Camptothecin (An Anticancer Drug) from Nothapodytes nimmoniana Graham: A Rare and Endangered Medicinal Plant from Western Ghats of Maharashtra "vide UGC letter F. No.42-212/2013 (SR) 22 March, 2013 has been fully utilized for the purpose for which it was dated sanctioned and in accordance with the terms and conditions laid down and that the balance of Rs 6740=0 remaining grant unutilized will be refunded back to the University Grants Commission.

SIGNATURE OF THE RINCIPAL INVESTIGA (Seal) Professor Sant Gadge Baba Amravati University Amravati University, Amravati (M.S.)

SIGN OF\CO-INVESTIGATOR (Dr. Syrendra R. Patil) Professor (Horticulture) Department of Fruit Science College of Horticulture Dr. P D. K. V., Akola (M.S.)

REGISTRAR rabe Baba Sant Amravati.

STATUTORY AUDITOR

3. Anita Surendra Patil, Ankit Subhash Kale, Surendra Rajaram Patil, Hariprasad Madhukarrao Paikrao (2016). Validation of accumulation of camptothecin content, an anti-cancer alkaloid in *Nathopodytes nimmoniana* graham. In phenotypic variants: method for identifying high-yielding sources of alkaloid. International Journals of Pharmacy and Pharmaceutical Sciences. Vol 8,Issue 9, pp 19-23.

### Paper presented in conferences/symposia: 04

- Mahure, S. M, Patil A.S. and Patil SR. Extraction and HPLC purification of anticancer drug camptothecin from Nathopodytes nimmoniana.101 Indian Science Congress, 2-7 Feb, 2014. Abs no 176, section XIV-plant Science, University of Jammu.
- 2. Anita Patil, Surendra Patil, Ankit Kale. Process optimization for *In Vitro* Production and Confermation of Camptothecin –an anticancer frug from Nathopodytes nimmoniana.4<sup>th</sup> *Bhartiya Vigyan Sammelan & Expo* 2015. 6<sup>th</sup> -8<sup>th</sup> February 2015,Goa Organized by Vigyan Bharti Government of goa; Goa University. National
- 3. Anita Patil "Validation of Accumulation of Camptothecin Content, An Anticancer Alkaloid In Nathopodytes nimmoniana Graham. In Phenotypic Variants: Method for Identifying High-Yielding Sources of Alkaloids" OP 45 National Conference of "Plant Diversity: Past and Present(Pdpp-2016)", held during 30<sup>th</sup>-31<sup>st</sup> December,2016 Dept of Botany,SGBAU.
- Anita Patil and Ankit Kale –Validation of Antiurolithiatic Potential of Abutilon Indicum(L.) and Passiflora Incarnata. OP51.". National Conference On plant Diversity Past and Present(Pdpp-2016)", held during 30<sup>th</sup>-31<sup>st</sup> December,2016 Dept of Botany,SGBAU.
- (d) Other impact, if any : Nil

Dr. P D. K. V., Akola (M.S.)

SIGNATURE OF CO-INVESTIGATOR Professor Department of Biotechnology Sant Gadge Baba Amravati University Amravati (M.S.) SIGNATURE OF CO-INVESTIGATOR (Dr. Surrendra R. Patil) Professor (Horticulture) Department of Fruit Science College of Horticulture Appl

REGISTRAR/PRINCIPAL REGISTRAR Sant Gadge Baba Amravati University, Amravati. An inexpensive Method for the determination of oxalate in rine and assessment of hyperoxaluria by nano-based paper strip technique (RGSTC/File-2011/DPP-166/CR-42 dated 27/2/2017)

UTILIZATION UPTO 31/8/2019

SUMMITTED TO

Rajiv Gandhi Science and Technology Commission (RGSTC) Government of Maharashtra, Mumbai-400032



SUMMITTED BY

Dr. Anita Surendra Patil Dr. Gajanan G. Muley Professor & Principal Investigator Assistant Professor and Co-Investigator Department of Biotechnology Department of Physics Sant Gadge Baba Amravati University, Amravati 444602(M.S)

> Dr.Hariprasad M. Paikrao Assistant Professor and Co-Investigator Government Institute of Forensic Science, Aurangabad

#### RAJIV GANDHI SCIENCE & TECHNOLOGY COMMISSION GOVERNMENT OF MAHARASHTRA, MUMBAI UTILIZATION CERTIFICATE FOR THE FINANCIAL YEAR 2018-2019

1. Title of Project: An inexpensive Method for the determination of oxalate in urine and assessment of hyperoxaluria by nano-based paper strip technique'

2. Name of the Institution : Department of Biotechnology, Sant Gadge Baba Amravati

	University, Amravati, Maharashtra (India)
3. Principal Investigator	: Dr. Anita Surendra Patil, Professor,
	Dept. of Biotechnology, SGB Amravati
	University, Amravati,444602
Co-Investigator	: Dr. Gajanan G. Muley, Assistant Professor,
	Dept. of Physics, SGB Amravati University, Amravati 444602
Co-Investigator	: Dr. Hariprasad Madhukarrao Paikrao, Assistant Professor, Government Institute of Forensic Science, Aurangebed (M.S)
	Selence, Autualigadad (W.S)

4. RGSTC Sanction latter No & : No: RGSTC/File-2011/DPP-166/CR-42 dated 27th March 2017

Date Sanctioning the Project

5. Amount brought forward from the previous financial year quoting RGSTC letter no. and date

: Rs. 8,32561=00

in which the authority to carry forward the said amount was given : i) Amount : **Rs. 1,72561=00** ii) Letter No : RGSTC/File-2011/DPP-166/CR-42 iii) Date : 8/1/2019

6. Amount Received during the financial year 2018-2019:

i) Amount : Rs.6,60000=00
ii) Letter No : RGSTC/File-2011/DPP-166/CR-42
iii) Date : 8/1/2019
7. Total amount that was available for the expenditure

8. Actual Expenditure (excluding commitments) : Rs. 7,70639=00 Incurred during the financial Year 2018-2019.

9. Balance amount available at the end of the financial year : Rs. 61,922=00

10. Unspent balance refunded, if any (details of cheque no.etc): Nil

11. Amount to be carried forward to the next financial year **Rs.** 61,922=00

#### **UTILIZATION CERTIFIACTE**

Certified that out of Rs. 660000=00 (Six lakh and sixty thousand only) of grants-in-aid sanctioned during the year 2018-2019 in favor of <u>PI-Dr. Anita Surendra Patil</u> under this commission/Department letter/order <u>No: RGSTC/File-2011/DPP-166/CR-42</u> dated 27/3/2017, and unspent balance of the previous year Rs. 1,72561=00 (One lakh seventy two thousand five hundred and sixty one only) and interest (Nil), a sum of Rs. 7,70639=00 (Seven lakh seventy thousand six hundred and thirty nine only) has been utilized for the purpose for which it was sanctioned and that the balance Rs. 61,922=00 (Sixty one thousand nine hundred and twenty two only) of remaining unutilized at the end of the year will be adjusted towards the grants-in-aid payable during the next year i.e. 2019-2020.

Sign of

Trate-Able (St. Patil MRP-ROSTC (Mumbai) Department of Biotechnology Bant Gadge Baba Amravai University Amravati (M.S.)

Sign of Head & BCAUDOUNTSTONFILIENERSITY Dates ANT GADGE BABA AMRAVATI UNIVERSITY, AMRAVATI,

Registrar Datgistrar Sant Gadge Dabe Amravati University

#### **Detail Statement of Expenditure** For Financial Year (1/4/2018 to 31/8/2019)

Sr.	Sanctioned	Carry forward	Funds allocated	Expenditure	Balance as on	Remark if
No	Heads	Amount for II <sup>nd</sup>	(Indicate	Incurred in	31/8/2019	any
(1)	(11)	year	Sanctioned or revised in Rs.6,60,000=00 (Total project fund)	2018-19		
1	Salaries/Wages		Amount Received 2 <sup>nd</sup> Year	2 <sup>nd</sup> Year		Rs.
2	Equipment	Nil	Nil	Nil	Nil	61,922=0
3	Staff Salary	77468=0	3,36000 =00	3,99468=00	14,000=00	forwarded
4	Consumables	4742=00	2,00,000=00	1,65854=00	38,888=00	for 2019-
5	Travels-PI	1070=00	20,000=00	12,036=00	9,034=00	2020
6	Analytical Charges	Nil	24,000=00	24,000=00	Nil	
7	Contingencies	Nil	30,000=00	30,000=00	Nil	
8	Overheads	89,281=00	50,000=00	1,39,281=00	Nil	
	Total	Rs. 1,72561=00	Rs. 6,60000=00	Rs. 7,70639=00	Rs.61,922 =00	

Sign of PI Sign of Head Accepunt officervos Drifverstry Date: Sant GADGE BABA AMRAVATI UNIVERSITY, MRP-RGSTC (Mumbai) Department of Biotechnology Sant Gadge Baba Amravati University Amravati (M.S.)

Registrar Date. Registrar Sant Gadge Baba Amravati University

# No.DST/INT/AGR/P-03/2015 Government of India Ministry of Science & Technology Department of Science & Technology (International Division)

Dated /9.06.2015

### ORDER

The composition of the team is as under:-

1.200	Indian Side Co-PI		Argentinean Side Co-PI
Co-PI-I	Dr. A.K. Gade, Asstt. Prof. Amravati University, Amravati, Maharashtra	Co-PI-I	Dr. Marcos DeritaNational University of Rosario, Argentina

The break-up of approved expenditure is as indicated below and no reappropriation (from one budget head to the other) shall be allowed under this project

SI. No.	Budget Head	1 <sup>st</sup> yr	2 <sup>nd</sup> yr	3 <sup>rd</sup> Year	total
1.	Number of visits (1 every year in each direction )	1 visit	1 visit	1 visit	3 visits
2.	Air Far Int'l' air travel (shortest route) by lowest Economy class (to & fro) for visit of Indian participants to Argentina including Domestic travel in India, Silver class medical insurance and visa fee @ Rs. 2,00,000/- per visit. (1 <sup>st</sup> & 3 <sup>rd</sup> visit of 21days each & 2 <sup>nd</sup> visit of 90 days by Indian PI/Student/Young researcher)	2,00,000	2,00,000	2,00,000	6,00,000
3.	Lodging & boarding for Argentina Scientist in India Accommodation charges @ Rs. 2000/- per day	42,000	1,80,000	42,000	2,64,000
4.	Living allowance (per diem) @ Rs. 2500 per day for short term visit and Rs. 60,000 per month for long term visits incldg. expenses towards local travel, airport transfer etc. (1 <sup>st</sup> & 3 <sup>rd</sup> visit of 21 days each & 2 <sup>nd</sup> visit of 90 days by Argentina /Student/ researcher)	52,500	1,80,000	52,500	2,85,000
5.	Consumables and Contingencies	15,000-	-15,000	15,000	-45,000-
6	Total :- (2+3+4+5+)	3,09,500	5,75,000	3,09,500	11,94,000

2. Sanction of the President is also accorded for release of initial grant of Rs. 3,09,500/- (Rupees three lakh nine thousand five hundred only) and the payment of this amount may be made to DDO, DST by means of electronic transfer(E-transfer) for the account of Registrar, SGB Amravati university, Amravati. Maharashtra.

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Contd. . . P/2

-Name/Designation of account holder -Bank Account No. - Name of Bank & Address - IFSC Code -MICR Code	<ul> <li>Registrar, SGB Amravati university, Amravati.</li> <li>20115413704 (S/B)</li> <li>Bank of Maharashtra, SGB University Branch, Tapovan Road Amravati</li> <li>MAHB0001331</li> <li>444014001</li> </ul>
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(This sanction order being 1<sup>st</sup> installment for implementation of this project, no UC/ SE is due from the grantee institute against this project at this stage).

3. Under the existing terms and conditions, sending side will bear the international air travel (air tickets and host side will bear the cost of lodging and boarding including perdiem of the scientist visiting under the project.

4. All project related Visits to be undertaken by the scientists from either side in connection with the implementation of the project shall require prior approval from this Department separately on a case to case basis before any expenditure is incurred in this regard.

5. As per MOF instructions dated 13.07.2009 all cases of air travel both domestic and international, the officials may travel only by Air India. For travel to stations not connected by Air India the officials may travel by Air India to the Hub/Point closest to their eventual destination, beyond which they may utilize the services of another airline which should also preferably be an alliance partner of Air India.

6. The institute will maintain separate audited accounts for the project and the amount of grant will be kept in a bank account earning interest. The interest earned should be reported to DST while submitting the Statement of Expenditure/Utilization Certificate. The interest thus earned will be treated as a credit to the institute to be adjusted towards further installment of the grant.

7. This sanction is subject to submission of Utilization Certificate (UC) and audited statement of the expenditure(SE) along with up to date progress report at the end of each financial year for the grants already received under the project and seeking specific approval of this Department for carry forward of unspent funds to the next financial year for utilization of project and list of exchange visits of scientists under the project.

8. As per latest instructions of Ministry of Finance (Department of Expenditure), Agencies will have to enter & upload the Utilization Certificate in PFMS portal besides sending it in physical form to this Division. The next installment will be released only after confirmation of the entry of Utilization Certificate in the PFMS

9. As per rule 211 (i) of GFR, the accounts of the project shall be open to inspection by sanctioning authority/audit whenever the institute is called upon to do so.

10. The expenditure involved shall be met out of the budget head under Demand No.86-Department of Science and Technology -3425-Other Scientific Research (Major Head) 60.798 International Cooperation (Minor Head) -12-S&T Cooperation with Other Countries-12.00.31-Grants-in-aid (Plan) during the year

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2015-2016.

11. This issues under the delegated powers to the Ministries/Departments and with the concurrence of IFD vide Dy. No. C/1235 /IFD//2015-2016, dated 11.06.2015.

12. This Sanction Order has been entered in Grant in aid Register (2015-16) at Sl. No. 207 Machuk

> (Dr. Jagdish Chander) Scientist'G'

> > Contd. . . P/3

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(Dr. Jagdish Chander) Scientist'G'



# Medical Mycology Current Trends and Future Prospects



Editors Mehdi Razzaghi-Abyaneh Masoomeh Shams-Ghahfarokhi Mahendra Rai



# Therapeutic Medicinal Plants From Lab to the Market





Editors Marta Cristina Teixeira Duarte Mahendra Rai







# Clinical Microbiology Diagnosis, Treatment and Prophylaxis of Infections THE MICROBIOLOGY OF SKIN, SOFT TISSUE, BONE AND JOINT INFECTIONS

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Edited by

### KATERYNA KON

Associate Professor, Department of Microbiology, Virology and Immunology, Kharkiv National Medical University, Kharkiv, Ukraine

### MAHENDRA RAI

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# Clinical Microbiology Diagnosis, Treatment and Prophylaxis of Infections THE MICROBIOLOGY OF CENTRAL NERVOUS SYSTEM INFECTIONS

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# The Microbiology of Respiratory System Infections

*Edited by* Kateryna Kon Mahendra Rai



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# Nanotechnology for Bioenergy and Biofuel Production



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### Preface

There is a greater need to search for alternative sources of energy due to the limited availability of fossil fuels. The use of biofuels is the common alternative in front of the whole world, and its use is significantly increased in many countries. Generally, biodiesel is used as a fuel for diesel engines due to its technical, environmental, and strategic advantages. Apart from these, biodiesel is technically competitive with conventional, petroleum-derived diesel fuel and requires virtually no changes in the fuel distribution infrastructure. Other advantages of biodiesel as compared to petrodiesel include reduction of most exhaust emissions, biodegradability, higher flash point, inherence, and the fact that it is of domestic origin. However, ethanol is mostly used for Otto cycle engines as an alternative to gasoline.

Nanotechnology is one of the most important research areas in the twenty-first century. The potential applications of nanobiotechnology in the production of sustainable bioenergy and biosensors have encouraged researchers in recent years to investigate new novel nanoscaffolds to build robust nanobiocatalytic systems. Different kinds of nanomaterials have been developed as per the need for application purpose.

Various metal nanomaterials are found to have direct or indirect application in the production of biofuels. Among these, nanoparticle-mediated enzyme hydrolysis of lignocellulosic residues for the production of ethanol is in practice. Magnetic and metal oxide nanoparticles such as TiO<sub>2</sub>, ZnO, SnO<sub>2</sub>, etc. are commonly used matrices for enzyme immobilization. Magnetic nanoparticles are found to have potential uses in the field of biofuel and bioenergy, i.e., in the production of bioethanol from lignocellulosic materials by immobilization of enzymes like cellulases and hemicellulases by physical adsorption, covalent binding, cross-linking, or specific ligand spacers. Such nanomaterials are generally called as nanocatalysts, which are fabricated by immobilizing enzymes with functional nanomaterials as enzyme carriers.
The book would be immensely useful for a diverse group of readers including physicists, chemists, microbiologists, biotechnologists, food technologists, agriculture engineers, nanotechnologists, and those who are interested in clean technologies. The students should find this book useful.

Amravati, Maharashtra, India Lorena, SP, Brazil Mahendra Rai Silvio Silvério da Silva

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Nanotechnology Applied To Pharmaceutical Technology

Mahendra Rai · Carolina Alves dos Santos Editors

### Nanotechnology Applied To Pharmaceutical Technology



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### Preface

Pharmaceutical technology is based on the use of different technologies for the preparation and development of pharmaceutical products. It is a branch of the pharmaceutical sciences constantly linked to the development and use of innovative technologies for the production and development of products with maximum effectiveness, adequate dosage, and stable formulations, aiming at different applications in living beings.

Nanotechnology is a recent, highly interdisciplinary and versatile branch of science that allows to develop products with interesting and novel structural characteristics totally different from those obtained by the compounds in their conventional form, a fact that attributes differentiated properties and specific interaction mechanisms, which adds technological value to the newly developed products. Thus, the use of nanotechnology in the pharmaceutical science promotes the design and development of products known as nanopharmaceuticals, which are composed of nanoscale properties with slow drug release systems, as stabilizers of formulations, pharmaceutical activities, or systems for detection and evaluation. The possibility of obtaining different mechanisms of response provided by the nanotechnologically developed products allows varied combinations enriching the range of products to be developed from this technology.

In this context, the book seeks to integrate the pharmaceutical sciences with nanotechnology as a way to obtain innovative products and solutions applied to the pharmacy. It provides the reader with a broad vision where nanotechnology is employed with different perspectives and applications. This approach focuses on the application of nanotechnology in pharmaceutical technology and also represents an interdisciplinary nature that can be used by people interested in expanding their knowledge in applied nanotechnology. The present book is highly interdisciplinary and would be very useful for a diverse group of readers including pharmacologists, nanotechnologists, microbiologists, biotechnologists, clinicians, and those, who are interested in development of nanoproducts. The students should find this book useful and reader-friendly.

Amravati, India Sorocaba, Brazil Mahendra Rai Carolina Alves dos Santos The original version of the book was revised. The incorrect country name in preface "Sorocaba, India" has been corrected as "Sorocaba, Brazil".

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## Mahendra Rai Avinash P. Ingle Serenella Medici *Editors*

# Biomedical Applications of Metals



Biomedical Applications of Metals

Mahendra Rai · Avinash P. Ingle Serenella Medici Editors

## Biomedical Applications of Metals



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### Mahendra Rai · Jayanta Kumar Biswas *Editors*

## Nanomaterials: Ecotoxicity, Safety, and Public Perception



Nanomaterials: Ecotoxicity, Safety, and Public Perception

Mahendra Rai · Jayanta Kumar Biswas Editors

## Nanomaterials: Ecotoxicity, Safety, and Public Perception



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### Foreword

Nanotechnology is an important and enabling technology today. With the confluence of ideas from diverse disciplines, it is anticipated to positively affect various facets of our lives. Nanomaterials derive their strength by harnessing the property, which with decrease in particle size surface area of materials increases, and some novel properties emerge. This fundamental premise paves way to many possibilities as envisioned by Richard Feynman, "There is Plenty of Room at the Bottom" for "manipulating and controlling things on a small scale". Nanomaterials are expected to enable many innovations in the form of diverse devices, processes and products in the domains of medicine, energy, environment, and others.

The last two decades or so have witnessed a burst of interest and activity in synthesis, characterization, and application of nanomaterials. There is much scope for developing ecotoxicological information. The nanomaterials could interact with different biotic and abiotic factors in the environment and cause adverse biological and ecological effects at different levels. To assess environmental risks, scientific data and knowledge on transport of nanomaterials in the environment and their effects on health are to be properly examined.

Future development of nanotechnology should be accompanied with research on (eco)toxicity associated with possible risks by analyzing behavior and fate of nanomaterials in the ecosystem inclusive of its biotic organisms and human health. There must be a reasonable concern about plausible and potential hazards, based on scientific investigation.

Given the vibrancy of nanoscience and technology, the edited book entitled *Nanomaterials: Ecotoxicity, Safety, and Public Perception* is a topical and useful book providing the state of knowledge on ecotoxicity of nanomaterials and safety issues. The book consists of 17 chapters arranged in three sections dealing with ecotoxicity, risk, and safety issues, regulations and public perceptions. All the chapters have been written by experts from different countries, who have offered authoritative treatments of nanoecotoxicological aspects. It has dealt with nanomaterials in food, soil, medicine, and unintentionally produced nanomaterials, and addresses key issues of ecotoxic effects of nanomaterials on life and environment

including humans. It has iterated the need for an increased level of understanding of risks and vulnerabilities that are likely to influence public, regulatory, and non-governmental activities. The book has been designed with adequate coverage for a multisectorial audience of academics, researchers, students, professionals, and policy-makers. I congratulate the editors for their endeavor in producing this useful book.

Bangalore, India

C. N. R. Rao Linus Pauling Research Professor

### Preface

Nanoscience and nanotechnology refer to the control and manipulation of matter at nanometer dimensions for tailoring and hybridizing the physical, biological, and engineering properties of matter. It is a burgeoning, vibrant, and promising field which has offered the possibilities and stunning new achievements as envisioned by Richard Feynman (1959): "There is plenty of room at the bottom". Nanotechnology is not only an emerging field of study, it has become an industry with transforming effects. The use of nanomaterials is expected to offer great promise for diverse devices, procedures, and products in the domains of consumer goods, medicines, energy, environment, and many others. But there is the other side of the coin; the environment is prone to suffer from nanopollution and toxic insult due to release and build up of nanocontaminants during different phases of lifecycle of nanomaterials, viz., production, transportation, disposal, entry into ecosystems, accumulation, bioconcentration, transfer across trophic boundaries, etc. The NMs raise the possibility that they could interact with and cause adverse biological effects at cellular, subcellular, and molecular levels. Assessing potential environmental/ ecological risks requires quality information on transport and fate of nanoparticles in the environment, exposures, and vulnerabilities of organisms to the nanomaterials and standard methods for assessing toxicity for aquatic or terrestrial organisms and human health. Nanomaterial toxicity is extremely complex, and our understanding about ecotoxicity is very fussy and incomplete. The systematic risk characterization and evaluation of the safety of nanomaterials require a multidisciplinary approach and convergence of knowledge and efforts from researchers and experts from toxicology, biotechnology, materials science, chemistry, physics, engineering, and other branches of life sciences. Although studies are beginning to appear in the literature addressing the toxicity of various nanomaterials and their potential for exposure, at this stage definitive statements regarding the impacts of nanomaterials on human health and the environment remain sketchy and scanty. Because of our limited understanding of the environmental implications of nanomaterials, there remains a gulf between their promises and pitfalls although the flourishing development of new nanomaterials as well as emerging new applications. No doubt present time demands a clarion call for an increased level of precautions with regard to nanomaterials, as has happened with other emerging contaminants and technologies (e.g., biotechnology). The perceptions of risks and of benefits will vary and are likely to influence public, regulatory, and nongovernmental activities regarding risk and benefit evaluations. The Triple Helix, i.e., academia, government, and industry, needs to be extended to a fourth dimension, the civil society, because of the latter's role and relevance in today's world. Since people are receptive to fictional representations as they shape their imaginaries, the "nanomania" has rather alienated the entertainment and leisure industry, and given rise to misconceptions, misunderstandings, and distorted views regarding this particular technology rather than contributed to a rational and genuine discussion. Quo vadis? Where will we go? Systematic identification and assessment of the risks posed by any new technology are essential. Systematic research into the flip side of this new technology, especially the risks that it entails, is essential. A prudent, integrated, and holistic approach is required to develop best practices based on the scientific understanding about what we know and what we don't know but need to know so that the engines of creation would not become the engines of destruction.

With suitable breadth of coverage, this edited book has been *designed*, *written*, *and organized* for the broad-spectrum multisectorial audience of researchers, students, professionals, and policy-makers. It is expected to be a pointer and reference resource and will cater to the needs of the researchers, professionals as well as the graduate and postgraduate students of nearly every discipline of science and engineering—toxicology, biotechnology, chemistry, physics, biology, medical science, material science & engineering, environmental science, and environmental engineering, who are interested to know the ecotoxicological aspects of the nanomaterial and the potential hazards and risks involved therein. This book would be an inspirational comprehensive resource for researchers, environmental professionals and practitioners as well as industrial portfolio holders, who are interested to earn a working knowledge of this fascinating field of Nanotechnology and Nanotoxicology.

This book includes important chapters written by specialists and experts in the concerned topic, and hopefully it will win wide acceptance among the undergraduate and graduate students of chemistry, biochemistry, pharmacology, microbiology, biotechnology, and allied subjects. It will be very useful for researchers working in nanotechnology, nanobiotechnology, toxicology, microbiology, ecology, and other related fields.

Last but not the least, we hope, this edited book surely will evoke some surprise and shock, hope and despair, and stimulate resolution and action through concerted and collaborative efforts of appropriate bodies to shape the nanoworld and save Mother Nature.

Amravati, India Kalyani, Nadia, India Mahendra Rai Jayanta Kumar Biswas

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His area of expertise includes microbial biotechnology and nanobiotechnology. His present research interests are nanobiotechnology in medicine and agriculture. He has published more than 380 research papers in national and international journals. In addition, he has edited/authored more than 47 books and 6 patents.



Jayanta Kumar Biswas is an Associate Professor at Department of Ecological Studies and International Centre for Ecological Engineering, University of Kalyani. He obtained M.Sc. in Zoology, M.Phil. in Ecology, and Ph.D. on ecotechnological management of aquatic systems. His interest and expertise span the following focal areas: ecotoxicology and bioremediation of toxic metal(loid)s; ecological engineering and ecotechnological remediation of water and soil contaminants; environmental microbiology; nanobiotechnology. He received many awards and fellowships including Fellow. National Environmentalists Association, Zoological Society of India, Zoological Society, Kolkata, Senior Scientist Award & Gold Medal (ZSI), Best Research Paper Award (AESA), Outstanding Reviewer Award (Chemosphere (Elsevier) & Environmental Geochemistry and Health (Springer); National Merit Scholarship (MHRD, GoI), etc. He is credited with publishing 125 research papers in reputed journals. He is serving as editorial board member of several international journals of repute, namely Science of the Total Environment (Elsevier); Environmental Geochemistry and Health (Springer); Current Pollution Reports (Springer); Energy & Environment (SAGE).

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# Part I Ecotoxicity

# Chapter 1 Nanomaterials: What Are They, Why They Cause Ecotoxicity, and How This Can Be Dealt With?



# Mahendra Rai, Indarchand Gupta, Avinash P. Ingle, Jayanta Kumar Biswas and Olga V. Sinitsyna

**Abstract** Nanomaterials have been benefiting human by their wide applications in different fields. Till date, many types of natural and engineered nanomaterials have been reported. Each of them has specific characteristics, which are helpful in deciding their use for particular application. Although they are beneficial to the human beings, there is probability of their harmful effects on the ecosystem. After the desired use of the nanomaterials, they are routinely disposed of into the environment either intentionally or unintentionally. This scenario can create the harmful environment for the whole ecosystem. The ecotoxicity of nanomaterials is an imperative point to be considered for the safety of flora and fauna. Hence, with argument on their characteristics and applications, their safety for human and environment should also be considered. Therefore, the present chapter introduces the nanomaterials, encompasses the discussion on major types of nanomaterials, which are being available naturally and others that are synthesized artificially. The parameters which make

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nanoparticles harmful to ecosystem have also been discussed. Moreover, the special emphasis is given on how the scientific community can deal the situation to avoid the harmful effects of nanoparticles so that it can be beneficial to mankind without causing any damage to ecosystem.

**Keywords** Nanoparticles · Engineered nanomaterials · Toxic effects · Ecotoxicity

#### **1.1 Introduction**

Materials, which consist of particles with at least one dimension smaller than 100 nm, are called nanomaterials. The era of the increasing popularity of the nanomaterials began with the discovery of fullerenes in 1985 by Harold Kroto, James Heath, Sean O'Brien, Robert Curl, and Richard Smalley (Kroto et al. 1985). According to The Nanodatabase (2018), commercial products containing fullerenes are already on the market. These are rackets for tennis and badminton, and skin care products. Another database, The Nanotechnology Consumer Product Inventory, was created in 2005 to document the emergence of nanotechnology products in the consumer market (Vance et al. 2015). The number of products in the Inventory is constantly growing. The most advanced category for the introduction of nanomaterials is the Health and Fitness. Oxides of titanium, silicon and zinc are the most produced nanomaterials worldwide, and silver nanoparticles (AgNPs) are the most popular additive in products (Vance et al. 2015). The nanoparticles act as antimicrobial agents, pigments, UV filters, protective coatings, reinforcing additives, etc. The global market of the nanomaterials is growing rapidly. According to a pessimistic forecast, its value will reach \$11.3 billion in 2020, and an optimistic forecast predicts the market value of \$55.0 billion in 2022 (Inshakova and Inshakov 2017).

The active use of the nanomaterials raises the question of their safety for the environment. The first publication considering the potential ecological impacts of the engineered nanomaterials appeared after more than 15 years of the widespread development of nanotechnology (Colvin 2003; Skjolding et al. 2016). A new discipline, nanoecotoxicology, began to study the environmental effects of nanoscale materials. The evaluation of the behavior and fate of the nanomaterials in complex environmental matrices is an extremely challenging problem (Boyes et al. 2017). The possible release of the nanoparticles into the environment should be considered across the entire life-cycle of a product, including manufacture, use, and disposal. The nanoparticles can possess not only advantageous but also unique toxic properties different from the bulk materials. Potential adverse effects should be taken into account from the cellular level to the whole ecosystem. The environmental hazard evaluation, based on dose-by-mass, may not be valid for the nanomaterials, and the question on the relevant dose-metrics in nanoecotoxicology is still open.

How to use the gifts of nanotechnology and protect the environment from the possible harmful effects is one of the most pressing issues of today. In the present chapter, we try to answer these major questions.

# **ETHNOBOTANY** Application of Medicinal Plants



Editors José L. Martinez Amner Muñoz-Acevedo Mahendra Rai



# ETHNOBOTANY Local Knowledge and

**Traditions** 

Editors José L. Martinez Amner Muñoz-Agevedo Mahendra Rai

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# MYCOTIC KERATITIS

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# Nanoformulations in Human Health

**Challenges and Approaches** 



Nanoformulations in Human Health

Sushama Talegaonkar • Mahendra Rai Editors

# Nanoformulations in Human Health

**Challenges and Approaches** 



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#### Foreword

In the era of the Fourth Industrial Revolution, several technologies are poised to bring swift changes in the approach and the living style of mankind. Nanotechnology is one of such fields that has multifaceted applications in healthcare ranging from medicines, surgical materials to devices and even means to prevent diseases. Initially, nanotechnology was looked upon for novel and targeted drug delivery systems, but now it is used for precision medicine for prevention of disease through appropriate medical gazette.

The book entitled *Nanoformulations in Human Health: Challenges and Approaches* edited by Dr. Sushama Talegaonkar and Dr. Mahendra Rai focusses on the application of nanoformulations in human healthcare. The authors have dedicated years in the field of nanomedicine and have carried out original research work achieving numerous milestones in their professional journey. They have helped many students and scholars in this field. This book reflects dedication, hard work, and collaborative efforts of the authors. I congratulate the authors and contributors for ensemble of in-depth information in simple language. I am sure that the book will prove to be useful to healthcare specialists including researchers, scholars, and academicians.

This book provides comprehensive information of nanomedicine with special emphasis on drug targeting, theranostics, chronic disorders, and topical afflictions. There is a specific emphasis on preformulation of active pharmaceutical ingredients, advantages of nanoformulations over conventional therapy, and formulation, characterization, and application of various nanocarriers. Parts have been dedicated entirely on various routes of drug delivery like ocular, topical, and nose-to-brain targeting along with elaborate chapters covering diseases like glaucoma, epilepsy, colonic ailments, candidiasis, neuropathic pain, respiratory disorders, tuberculosis, and cancer.

R. K. Goyal

Vice-Chancellor, DPSRU, New Delhi, India

#### Preface

Nanotechnology, i.e., the science of nanosized agents, is one of the fastestgrowing fields which is on its way of giving breakthroughs to the healthcare industry. Nano-agents usually are in the size range of 1-100 nm. Owing to their extremely small size and majorly large surface area, they possess distinct special characteristics which enable them to do wonders in biomedical applications. Characteristics like ability to cross the cell membrane and increased reactivity make them more useful than conventional therapy. Particularly in biomedical sciences, nanotechnology can be used for therapy and diagnostics. In addition to their therapeutic properties, they are more commonly used to carry the therapeutic agents to the diseased site in the body, thus achieving targeted therapy. Delivery of drugs through a nano-agent directly to the target site also results in a decrease of the required dose of the active agent. More recently, the application of nanotechnology has been explored for monitoring the progress and success of therapy inside the body. Many countries are now investing in developing nanotechnology for healthcare applications. For instance, the first generation of cancer drugs delivered via nanoparticles has already been approved by the US Food and Drug Administration (FDA).

However, nanoformulation for healthcare is also a topic of hot debate as this minute size also brings the concern of toxicity. Though regardless of the varying opinions, the research and development have actively been going on to take nanotechnology to the commercial level to be used for the benefits of the health of humans.

The main goal of the book titled *Nanoformulations in Human Health: Challenges and Approaches* is to serve as a complete reference guide to understand the role of nanotechnology in the healthcare system and as a ready reference for all aspects related to the theme. The book is intended to provide a multidisciplinary approach signifying the role of nanosizing in the treatment of various challenging diseases. It is planned to highlight the pathophysiology of disease with special emphasis on various targets, receptors, biomarkers, and transporters associated with the disease. Young researchers/ scientists who plan to initiate research in this important field would find this book extremely relevant and handy. Each chapter of this book will be immensely useful to identify the new targets for drug delivery systems and drug discovery. This book is a true amalgamation of the experience and expertise of all the contributors in the field of nanotechnology especially in designing novel nanoformulations in the treatment of various challenging diseases. It is an exhaustive compilation of the multifaceted arena of nanoformulations in the healthcare system.

This book includes six parts concerning nanoformulations: Part I includes emerging trends and challenges in the area; Part II describes their role in drug targeting; Part III incorporates applications in ocular diseases; Part IV deliberates role in topical diseases; Part V discusses their role in natural therapeutics delivery; and finally, Part VI focuses on other applications such as in respiratory diseases, tuberculosis, and cancer.

The book would cater to the needs of postgraduate students of pharmacy, nanotechnology, and biotechnology, medical students, and researchers. In addition, the book will also be very useful for the pharmaceutical industries, regulatory bodies, pharmacy, medical institutes, etc. involved in research and development activities related to drug discovery, newer treatment modalities, and technology.

New Delhi, India Amravati, Maharashtra, India Sushama Talegaonkar Mahendra Rai

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Part I

Emerging Trends and Challenges in Nanoformulations

#### **Princeton University**

#### **Department of Geosciences**

Guyot Hall, Princeton, New Jersey 08544-1003 Telephone: 609-258-4100 Fax: 609-258-1274 December 13, 2017

Prof. Dr. Murlidhar Chandekar Vice-Chancellor Sant Gadge Baba Amravati University Amravati -444602 (MS) India Email: <u>murlidhar.chandekar@gmail.com</u>

Re: Permission for Prof. Dr. S. Khadri to join our fieldwork Feb. 1-Feb. 15, 2018.

Respected Dr. Murlidhar Chandekar

For the past five years Dr. Syed Khadri, Head of Geology at Amravati University, has collaborated with our team of geologists from Princeton University (USA) and the University of Lausanne (Switzerland) on Deccan Trap studies, including high-resolution U-Pb zircon dating of Deccan volcanic eruptions. Our team has been very successful in obtaining the first U-Pb ages of Deccan basalt flows (Schoene et al., 2014, 2015, Science). In subsequent fieldwork we collected many more red boles for dating and this work has also been very successful in adding many new high-precision U-Pb age dates for the Deccan Traps. This work is now being written up for submission to a high-profile journal. This high precision Deccan geochronology will have farreaching implications for understanding the nature and tempo of eruptions, gas emissions and their devastating effects globally leading to the mass extinction at 66 Ma.

Despite the considerable success so far we have not quite hit the jackpot – which is dating the killer lava flows that caused the mass extinction. But we are very close. In fact, we have dated lava flows that are just 60,000 years below the mass extinction. This is why we need to renew our fieldwork with an intensive search for the killer lava flows precisely at the mass extinction horizon now U-Pb dated elsewhere at 66.021±0.024 Ma.

For this reason we are planning our next phase of fieldwork for February 1-15, 2018. Dr. Syed Khadri is a critical and necessary member of our international team as he has much experience in Deccan Trap studies and leads our team in fieldwork. I am contacting you in the hope that you will grant Dr. Khadri permission to join us in fieldwork and continue this critically important research project. Thank you for your consideration of this request.

Sincerely,

- de fille

Prof. Gerta Keller Department of Geosciences Princeton University Princeton NJ 08544, USA email: gkeller@princeton.edu

#### **Princeton University**

#### **Department of Geosciences**

Guyot Hall, Princeton, New Jersey 08544-1003 Telephone: 609-258-4100

February 13, 2020

Dr. Prof. Syed Khadri Professor and Head, Department of Geology SGB Amravati University, Amravati-444602 (MS) India

Dear Syed

It is a pleasure to write a letter of intent to collaborate with you and Dr. Jahnavi Punekar (IITB) on your endeavor to establish the Lonar Science Centre at Amravati with a Field Station and Planetarium at the Lonar Crater in India's Deccan Volcanic Province. I congratulate you both for taking the initiative for this important project, which will be a great asset to India's for student education, science and research, furthering the integration of multidisciplinary projects in India, as well as attracting collaboration with foreign scientists.

Thank you for inviting me to participate in this project. I will be most happy to contribute to this project with my experience working in India over more than a dozen years specializing in Deccan volcanism and associated environmental and climate changes and the likely cause of the end-Cretaceous mass extinction.

My team closely collaborated with you over the past 10 years especially on Deccan volcanism with much success. I introduced Princeton's geochronologists to date Deccan eruptions, a project that has been highly successful and you have been a most important integral part of this effort. The Western Ghats eruptions and Malwa Plateau have now been dated based on high-resolution U-Pb zircon dating that resulted in influential *Science* publications (Schoene et al., 2015, 2019; Eddy et al., in press). This effort has already brought many foreign teams to India to investigate the age, nature and environmental effects of Deccan volcanism. In addition, it has re-invigorated India's scientists working on Deccan Traps and brought them important international attention and collaboration with foreign teams.

I believe the panned Field Station and Planetarium at Amravati with the Field Station and Planetarium at the Lonar Crater in India's Deccan Traps will further invigorate the both India's science education and foreign collaboration. I wish you success in this endeavor and look forward to participate.

Sincerely,

cente fille

Prof. Gerta Keller Department of Geosciences Princeton University Princeton NJ 08544, USA email: gkeller@princeton.edu

UNIL | Université de Lausanne Institut des sciences de la Terre bâtiment Géopolis bureau 3879 CH-1015 Lausanne

Prof. Dr. Thierry Adatte Tel. 0041 21 692 3526 Prof. Dr. Murlidhar Chandekar Vice-Chancellor Sant Gadge Baba Amravati University Amravati -444602 (MS) India Email: murlidhar.chandekar@gmail.com

Lausanne, 12.12. 2018

Concerns: Permission for Prof. Dr. S. Khadri to join our fieldwork Feb. 1-Feb. 12, 2019

Respected Dr. Murlidhar Chandekar

For the past five years Dr. Syed Khadri, Head of Geology at Amravati University, has collaborated with our team of geologists from the University of Lausanne (Switzerland) and Princeton University (USA) on Deccan Trap studies, including high-resolution U-Pb zircon dating of Deccan volcanic eruptions. Our team has been very successful in obtaining the first U-Pb ages of Deccan basalt flows (Schoene et al., Science 2015). In subsequent fieldwork we collected many more red boles for dating and this work has also been very successful in adding many new high-precision U-Pb age dates for the Deccan Traps. This work has been now accepted in the prestigious journal Science. This high precision Deccan geochronology will have far reaching implications for understanding the nature and tempo of eruptions, gas emissions and their devastating effects globally leading to the mass extinction at 66 Ma. In 2018, we have have extended our dataset to the eastern part of Deccan Trap, around Hyderabad, but we still need to complete our dataset in western India (Gujarat). For this reason, we are planning our next phase of fieldwork for February 1-12, 2019. Dr. Syed Khadri is a critical and necessary member of our international team since he has much experience in Deccan Trap studies and leads our team in fieldwork. I am contacting you in the hope that you will grant Dr. Khadri permission to join us in fieldwork and continue this critically important research project.

Thank you for your consideration of this request.

Sincerely,

Prof. Dr. T. Adatte

Faculté des géosciences et de l'environnement Institut des sciences de la Terre

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Prof. Dr. Thierry Adatte Tel. 0041 21 692 3526 Prof. Dr. Murlidhar Chandekar Vice-Chancellor Sant Gadge Baba Amravati University Amravati -444602 (MS) India Email: murlidhar.chandekar@gmail.com

Re: Permission for Prof. Dr. S. Khadri to join our fieldwork January 5 - 22, 2020

Respected Dr. Murlidhar Chandekar

For the past five years Dr. Syed Khadri, Head of Geology at Amravati University, has collaborated with our team of geologists from the University of Lausanne and Princeton University (USA) on Deccan Trap studies, including high-resolution U-Pb zircon dating of Deccan volcanic eruptions. Our team has been very successful in obtaining the first U-Pb ages of Deccan basalt flows (Schoene et al., Science 2015, 2019). This high precision Deccan geochronology will have far reaching implications for understanding the nature and tempo of eruptions, gas emissions and their devastating effects globally leading to the mass extinction at 66 Ma.

We recently got some news data from the Malwa area (Narmada River). these data demonstrate that the lower Malwa Plateau basalts are temporally correlative with the first pulse of Deccan volcanism and provide new constraints on its initiation and duration. The first eruptive pulse is coeval with a  $\sim$ 200 kyr period of global warming, and we suggest that eruption through the sedimentary rocks of the Narmada-Tapti rift basin provided the necessary CO<sub>2</sub> to drive this warming through contact metamorphism of some Permian coal levels. We need to sample these coal levels (Indore-Jabalpur-Nagpur) in order to evaluate the Mercury, Sulfur and carbon contents. Another part of the fieldtrip will be will be focused on additional sampling of the Western Ghats basalt flows to refine our age model.

For this reason, we are planning our next phase of fieldwork for January 5-22 2020. Dr. Syed Khadri is a critical and necessary member of our international team since he has much experience in Deccan Trap studies and leads our team in fieldwork. I am contacting you in the hope that you will grant Dr. Khadri permission to join us in fieldwork and continue this critically important research project. Thank you for your consideration of this request.

Sincerely,

Prof. Dr. T. Adatte

HADAYU

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January 10, 2020

Dr. Syed F.R. Khadri Sant Gadge baba Amravati University Amravati, India

Dear Professor Khadri,

This letter expresses my enthusiastic interest to collaborate with you in joint scientific projects involving the geological evolution of India. In particular, I am interested in continuing our collaboration on the geochronology of the Deccan Traps, as well as your new effort to study the Lonar Crater. I believe that this partnership will benefit both of our universities through continued interaction between Indian and US students and faculty, production of high quality analytical and field data, and publication and dissemination of our research in renowned academic journals. I believe that this partnership will continue to strengthen ties between the Indian and US scientific communities, and I look forward to our continued collaboration.

Please let me know if there is any additional information I can provide.

With best regards,

Un V. Edilp

Dr. Michael P. Eddy

#### **Princeton University**

#### **Department of Geosciences**

Guyot Hall, Princeton, New Jersey 08544-1003 Telephone: 609-258-4100 Fax: 609-258-1274 November 23, 2015

Dr. Mohan Khedkar Vice-Chancellor Sant Gadge Baba Amravati University Amravati -444602 (MS) India Email: vc@sgbau.ac.in Fax: 0091-721-2662135/266949

Respected Dr. Mohan Khedkar

For the past two years Dr. Syed Khadri, Head of Geology at Amravati University, collaborated with our team of geologists and geochronologists from Princeton University (USA) and the University of Lausanne, Switzerland, on Deccan Trap studies and in particular to conduct high-resolution U-Pb zircon dating of Deccan volcanic eruptions. Our team has been extremely successful to date in obtaining the first U-Pb ages of Deccan basalt flows and the first results have been published in Science (Dec. 2014 and Jan. 2015). This was the first step in what are revolutionary results for high precision Deccan geochronology with far-reaching implications for understanding the nature and tempo of eruptions and gas emissions and their devastating effects globally leading to the mass extinction 65 m.y. ago. In our January 2015 fieldwork we collected a series of samples that have also yielded promising results and also revealed where additional fieldwork is necessary for the next step in this geochronology study, which will result in excellent publications in high-visibility international journals.

For this reason we are planning our next phase of fieldwork for January 15-31, 2016. Dr. Syed Khadri is a critical and necessary member of our international team as he has much experience in Deccan Trap studies and leads our team in fieldwork. I am contacting you once again in the hope that you will grant Dr. Khadri on duty permission to join us in fieldwork as a part of our International Research Collaboration and continue this critically important research project. Thank you for your consideration of this request.

Sincerely,

certe file

Prof. Gerta Keller Department of Geosciences Princeton University Princeton NJ 08544, USA email: gkeller@princeton.edu

#### **Princeton University**

#### **Department of Geosciences**

Guyot Hall, Princeton, New Jersey 08544-1003 Telephone: 609-258-4100 Fax: 609-258-1274 October 26, 2016

Dr. Muralidhar Chandekar Vice-Chancellor Sant Gadge Baba Amravati University Amravati -444602 (MS) India Email: <u>vc@sgbau.ac.in</u> Fax: 0091-721-2662135/266949 muralidhar.chandekar@gmail.com/vc@sgbau.ac.in

*RE: This letter requests permission for Prof. Syed Khadri to lead fieldwork in the Deccan Traps for an international team of scientists during Jan. 15-31, 2017.* 

Respected Dr. Muralidhar Chandekar:

I'm Professor Keller at Princeton University contacting you in regards to permission for Professor Syed Khadri, Head of Geology at Amravati University, to participate and lead our ongoing international research collaboration with fieldwork scheduled between Jan. 15-31, 2017. This international research collaboration is between Princeton University and Sant Gadge Baba University and includes: Prof. Syed Khadri, Profs. Gerta Keller and Blair Schone from Princeton University, Dr. Thierry Adatte from the University of Lausanne, Switzerland, Prof. Jahnavi Punekar, Indian Institute of Science Education and Research Pune, Dr. Eric Font, University of Lisbon, Portugal, Dr. Michael Eddy, MIT, Boston, and several graduate students. Professor Syed Khadri is an integral part of our team as well as our team leader in conducting fieldwork because of his extensive expertise in Deccan volcanism.

*Background:* Our research project is on the *Chronology of Deccan Volcanism* and concentrates on the most important outstanding problem in Deccan research, namely determining the absolute age, the rate and tempo of basalt eruptions and their cause-and-effect relationship to the mass extinction at the end of the Cretaceous 66 million years ago. We began this project in 2013 and have been fantastically successful in age dating even individual lave eruptions revealing that over 3200 m of lava erupted in about 750,000 years with about 2000 m over just 250,000 years leading up to the mass extinction. This was the first step with revolutionary results for high precision Deccan geochronology with far-reaching implications for understanding the nature and tempo of eruptions and gas emissions and their devastating effects globally. This paper was published in Science January 2015 (Schoene et al) and accompanied by much international publicity. Since then our research has spawned numerous other teams working on Deccan volcanism as well as all other mass extinctions associated with volcanism.

Our research is continuing with even greater success thanks to yearly fieldwork lead by Dr. Syed Khadri. We have now succeeded in even higher resolution U-Pb age dating of Deccan lava flows, which narrows the voluminous lava flows down to the time of eruption that directly ended with the mass extinction (a paper is being prepared for publication in Science or Nature).

But our research is far from over, besides our work in the Western Ghats, we must now concentrate outside this area to evaluate the geographic distribution of lava flows, which is necessary to obtain estimates of the volume and rate of lava eruptions and gas emissions that are critical for modeling the climatic and environmental effects that appear to have caused the mass extinction. Also time is of essence because so many other teams have now embarked on Deccan studies. For these reasons we plan to conduct fieldwork during January 15-31, 2017.

Prof. Syed Khadri is an important member of our scientific team and his participation and extensive knowledge of the Deccan Traps are critical for the success of this project. For these reasons we kindly ask that you may grant permission to Dr. Syed Khadri to participate and lead our fieldwork. Our past successes virtually assure even greater success based on the upcoming fieldwork. It also assures a high visibility role for all participating institutions in this project.

Thank you for your consideration of this request. Sincerely,

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Prof. Gerta Keller Department of Geosciences Princeton University Princeton NJ 08544, USA email: gkeller@princeton.edu

#### **Princeton University**

#### **Department of Geosciences**

Guyot Hall, Princeton, New Jersey 08544-1003 Telephone: 609-258-4100 Fax: 609-258-1274 September 28, 2018

Prof. Dr. Murlidhar Chandekar Vice-Chancellor Sant Gadge Baba Amravati University Amravati -444602 (MS) India Email: murlidhar.chandekar@gmail.com

Re: Permission for Prof. Dr. S. Khadri to participate in documentary film on Deccan volcanism.

Respected Dr. Murlidhar Chandekar

I write to you to ask your permission for Dr. S. Khadri to participate in a documentary film on Deccan volcanism from October 27 through October 31, 2018.

For the past five years Dr. Syed Khadri, Head of Geology at Amravati University, has been a very important collaborator in my team of geologists from Princeton University (USA) studying the Deccan Traps. These studies have been highly successful and have received major media attention. Thanks to our research, Deccan volcanism as a driver of the dinosaur mass extinction is now of intense interest by several documentary filmmakers in the USA. I have chosen one highly regarded team of NYC filmmakers to create a documentary film on Deccan volcanism highlighting our research advances. This team will accompany my fieldtrip to the Deccan Traps at the end of October 2018.

Participation of Dr. Syed Khadri in this documentary is important because of his critical contributions to the successful research. His participation is also very beneficial for Amravati University as it highlights the important science contributions of one of your faculty members in what is expected to become an outstanding and internationally distributed science documentary on the Deccan Traps and the mass extinction.

Thank you for considering this request.

Sincerely,

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Prof. Gerta Keller Department of Geosciences Princeton University Princeton NJ 08544, USA email: gkeller@princeton.edu

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#### U-Pb zircon age constraints on the earliest eruptions of the Deccan Large Igneous Province, Malwa Plateau, India



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#### ABSTRACT

Climate instability driven by emission of volatiles during emplacement of large igneous provinces (LIPs) is frequently invoked as a potential cause of mass extinctions. However, documenting this process in the geologic record requires a holistic understanding of eruption rates, the location of eruptive centers, and potential sources of climate-changing volatiles. We present new chemical abrasion-isotope dilutionthermal ionization mass spectrometry (CA-ID-TIMS) U-Pb zircon geochronology from Malwa Plateau basalts on the northern margin of the Deccan LIP, India. These basalts have been previously interpreted as either an extension of the province's main volcanic stratigraphy or as an independent eruptive center active up to millions of years prior to the main eruptive phase. Our data instead demonstrate that the lower Malwa Plateau basalts are temporally correlative with the first pulse of Deccan volcanism and provide new constraints on its initiation and duration. Paleomagnetic data further indicate that upper Malwa Plateau basalts may be age-equivalent to the second, third, and fourth pulses of Deccan volcanism. The relative thicknesses of age-equivalent packages of basalt are consistent with eruption of the Deccan LIP from a southward-migrating eruptive center. The first eruptive pulse is coeval with a  $\sim$ 200 kyr Late Maastrichtian warming event preserved globally in contemporaneous stratigraphic sections. We propose that the first pulse of Deccan magmatism was more voluminous in the north, where it erupted through organic-rich sedimentary rocks of the Narmada-Tapti rift basin. Thermal metamorphism of these sediments could have been a source of sufficient CO<sub>2</sub> to drive the Late Maastrichtian warming event, which if true would reconcile the apparent dampened warming signals associated with later Deccan eruptive pulses.

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#### 1. Introduction

The temporal relationship between emplacement of large igneous provinces (LIPs) and mass extinctions is well known (e.g., Courtillot and Renne, 2003). However, the potential causal relationship between volcanic eruptions and paleo-environmental change remains uncertain. The most commonly proposed driver of paleo-environmental degradation is the syn-eruptive release of volatiles such as CO<sub>2</sub>, CH<sub>4</sub>, and SO<sub>2</sub> (e.g., Self et al., 2006). The ocean-atmosphere residence times of these gases differ dramatically,  $10^3-10^5$  yr (CO<sub>2</sub>) and  $10^0-10^1$  yr (CH<sub>4</sub>, SO<sub>2</sub>), and climate change

\* Corresponding author. E-mail address: mpeddy@purdue.edu (M.P. Eddy). driven by release of these volatiles acts on similar timescales (e.g., Self et al., 2006). Potential sources of these volatiles during LIP eruption are basaltic magmas and their differentiation products or crustal rocks that are metamorphosed during magma transport, emplacement, and eruption (Svensen et al., 2004; Self et al., 2006). Thus, precise chronologies of LIP eruption, robust estimates for eruptive tempo, and knowledge of the underlying crustal composition are all prerequisites to assessing the efficacy of LIP emplacement in driving climate change and environmental catastrophe.

Eruptions of the Deccan LIP in India (Fig. 1A) span the latest Cretaceous through the earliest Paleogene and has, along with the Chicxulub bolide impact (e.g., Alvarez et al., 1980), been implicated in driving the end-Cretaceous mass extinction and late Cretaceous environmental instability (Courtillot et al., 1988). Recent geochronologic results from the Western Ghats region of


Fig. 1. A) Simplified geologic map of Deccan large igneous province showing the study location on the southern margin of the Malwa Plateau relative to the Western Ghats and the location of the Paleozoic-Mesozoic 'Gondwana' sedimentary sequences (modified from Lehman et al., 2010). B) Generalized stratigraphic column for the southern Malwa Plateau basalts in the vicinity of Mandu-Mhow-Singachori (Khadri, 2003), showing the elevation of formation boundaries and magnetic polarity chrons, including flows with transitional polarity in gray. The locations of sampled 'red boles' are denoted as black circles and the U-Pb eruption ages for the seven 'red boles' that are interpreted to contain volcanic ash are shown. Each eruption age is shown as a white rectangle (i.e., eruption age interpreted from each sample's selected zircon) and a blue rectangle. The latter interpretation refines eruption ages by using the law of superposition and a Bayesian Markov Chain Monte Carlo technique (e.g., Schoene et al., 2019). C) Interpreted upper crustal seismic cross sections across the northern Deccan LIP showing the buried Narmada-Tapti rift (Kaila, 1988; Verma and Banerjee, 1992). (For interpretation of the colors in the figure(s), the reader is referred to the web version of this article.)

India have refined the eruptive chronology of the Deccan LIP (Schoene et al., 2019; Sprain et al., 2019), but these datasets disagree on their preferred models of either pulsed (Schoene et al., 2019) or near-constant (Sprain et al., 2019) eruption rates. Furthermore, the general applicability of either of these eruptive models to the entire province remains tenuous due to uncertainties in the correlation between the volcanic stratigraphy of the Western Ghats (Cox and Hawkesworth, 1985; Beane et al., 1986) and outlying areas. These correlations were previously established using geochemical and morphological similarities between packages of basalt (e.g., Lightfoot et al., 1990; Mitchell and Widdowson, 1991;

Khadri et al., 1999; Jay and Widdowson, 2008) and largely ignore the possibility of lateral changes in basalt flow morphology; broad geochemical similarities between different basalt formations; and the potential for magnetostratigraphy in different areas to record different magnetic reversals (e.g., Kale et al., 2019). Based on regional geologic mapping by the Geological Survey of India, Kale et al. (2019) proposed that the Deccan LIP represents a composite volcanic pile composed of several major eruptive centers that were active over several Myr. Such an interpretation challenges the regional significance of any eruptive model inferred solely from geochronologic data within the Western Ghats (Schoene et al., 2019; Sprain et al., 2019), as province-scale models of basalt volume are dependent on long-distance correlation of geochemicallydefined basalt formations.

The most robust way to temporally correlate basalt packages within a LIP is by using independent geochronologic constraints from many different areas to develop a high-resolution age model. To this end, we present new U-Pb zircon geochronology from the Malwa Plateau on the northern margin of the Deccan LIP (Fig. 1A). The Malwa Plateau includes  $\sim$ 21,000 km<sup>3</sup> of basalt and has been interpreted to represent a spatially distinct flow field based on flow morphology (Kale et al., 2019). Our new U-Pb dates provide a new age model for the area and an independent means of correlating eruption of Malwa Plateau basalts with both eruption of Western Ghats basalts and global paleoenvironmental records.

#### 2. Malwa Plateau basalts

The Malwa Plateau comprises an area of  $\sim$ 80,000 km<sup>2</sup> along the northernmost margin of the Deccan volcanic province (Fig. 1A). The stratigraphic thickness of basalts within the plateau ranges from >500 m in the south to  $\sim200$  m in the northwest and Schöbel et al. (2014) used these constraints to estimate that  $\sim$ 21,000 km<sup>3</sup> of basalt is currently preserved. However, prior to erosion the Plateau's volume was likely much greater. The basalts in this region have been classified into five geochemically defined formations and are named from stratigraphically lowest to highest: the Narmada Fm., Manpur Fm., Mhow Fm., Satpura Fm., and Singachori Fm. (Fig. 1B: Rao et al., 1985; Khadri, 2003). The Geological Survey of India has proposed an alternative stratigraphic nomenclature that divides the basalts of the Malwa Plateau into eight formations on the basis of flow morphology (summarized in Kale et al., 2019). However, in this study we utilize the geochemicallydefined stratigraphy, because flow morphology can change laterally in time-equivalent basalt flows, and the geochemically-defined stratigraphy of the Western Ghats (Cox and Hawkesworth, 1985; Beane et al., 1986) has been successful in identifying temporally distinct eruptive packages (e.g., Schoene et al., 2019).

The stratigraphy of the Malwa Plateau spans two magnetic reversals, from normal-to-reverse polarity and reverse-to-normal polarity (Fig. 1B: Pal and Bhimasankaram, 1976; Rao et al., 1985; Khadri and Nagar, 1994; Khadri et al., 1999; Khadri, 2003; Schöbel et al., 2014). While this sequence of reversals has been linked to the C30n-C29r-C29n sequence preserved in the basaltic stratigraphy of the Western Ghats region (Rao et al., 1985; Khadri et al., 1999; Khadri, 2003; Schöbel et al., 2014), this correlation has been largely untested with geochronologic data. A single <sup>40</sup>Ar/<sup>39</sup>Ar analysis of a plagioclase separate obtained at the base of the stratigraphy along the Plateau's western margin produced a date of 67.12  $\pm$  0.44 Ma (2 $\sigma$ ), suggesting that the oldest basalts in this area are from C30n (Schöbel et al., 2014). However, the general applicability of this result remains uncertain. The upper reversal, from reverse-to-normal polarity, is only preserved at the highest elevations on the southern Malwa Plateau, suggesting that most of the flows recording this transition have been eroded away (Khadri, 2003).

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Sample	Latitude (°N)	Longitude (°E)	Elevation (m)	Age (Ma)	— (Ma, 2.5% C.I.)	+ (Ma, 97.5% C.I.)
RBCQ	22.36098	75.72529	606	66.150	0.067	0.065
RBCN	22.41410	75.58862	554	66.199	0.074	0.055
RBCL	22.38527	75.53110	447	66.250	0.052	0.039
RBCJ2	22.37889	75.53228	417	66.256	0.042	0.034
RBCI	22.37673	75.53067	383	66.275	0.053	0.066
RBBZ	22.29716	75.37197	263	66.320	0.057	0.057
RBCF	22.30058	75.50123	237	66.358	0.070	0.065
RBBZ RBCF	22.29716 22.30058	75.37197 75.50123	263 237	66.320 66.358	0.057 0.070	0.057 0.065

 Table 1

 U-Pb eruption/deposition ages.

#### 3. U-Pb zircon geochronology methods and results

#### 4. Discussion

#### 4.1. Age of the C29n/C29r magnetic reversal

To better constrain the age of the Malwa Plateau basalts, we collected interflow paleosols, locally called 'red boles', for U-Pb zircon geochronology. Some red boles contain volcanic ash (e.g., Widdowson et al., 1997) and have been successfully sampled for U-Pb zircon geochronology in the Western Ghats (Schoene et al., 2019). While, the source of silicic volcanic ash remains uncertain, descriptions of poorly dated, but broadly coeval, rhyolitic complexes in the north of the Deccan Province suggest that small volume rhyolitic volcanism was ongoing during LIP emplacement (e.g., Rao, 1971).

Samples for U-Pb zircon geochronology were obtained from three transects over an area of  $\sim$ 400 km<sup>2</sup> along the southern margin of the Malwa Plateau (Fig. 1A). They were dated using a modified version of chemical abrasion-isotope dilutionthermal ionization mass spectrometry (CA-ID-TIMS) following the methods presented in Schoene et al. (2019), and were projected onto a regional stratigraphic column (Fig. 1B) following Khadri (2003). The trace element composition of each zircon was measured using the methods presented in Schoene et al. (2010) and a quadrupole-inductively coupled plasma-mass spectrometer (Q-ICP-MS) at Princeton University. Zircon photographs, U-Pb isotopic and compositional data, and more detailed methods are presented in the supplementary material.

Of eleven red bole samples that yielded zircon, four contained limited zircon; clearly detrital, rounded, Precambrian grains; or Cretaceous zircon with no clear geochemical relationship. Samples with a significant detrital component are not unexpected in this area, as the presence of inter-trappean sedimentary beds within the Malwa Plateau basalts suggests that the region was at low elevation during basalt eruptions (e.g., Tandon et al., 1995). Seven samples contained coherent age populations of euhedral zircon that we interpret to come from silicic volcanic ash. These populations preserve geochemical trends consistent with pre-eruptive, closed-system magmatic differentiation (Fig. S2) and were used to estimate eruption dates using a Bayesian Markov Chain Monte Carlo (MCMC) technique that assumes a pre-eruptive zircon age distribution bootstrapped from the selected zircons (Table S2: Keller et al., 2018). These eruption ages are further refined utilizing the law of superposition and a second Bayesian MCMC model to generate an age model for the dated stratigraphy (Table S3: e.g., Schoene et al., 2019). The resulting eruption age estimates are presented in Table 1. Alternative estimates based on different priors and more conventional data interpretation techniques are presented in Table S2. All uncertainties are analytical only and are reported at the 95% credible interval-level. The inclusion of systematic uncertainties, such as those associated with the calibration of the EARTHTIME <sup>202</sup>Pb-<sup>205</sup>Pb-<sup>233</sup>U-<sup>235</sup>U isotope dilution tracer and the U decay constants, is necessary for comparison with other geochronologic methods, and these uncertainties are reported in Table S3.

The magnetic polarity reversal sequence preserved within the southern Malwa Plateau basalts was previously assumed to be correlative to the C30n-C29r-C29n magnetic polarity chrons (Khadri et al., 1999; Khadri, 2003; Schöbel et al., 2014). Our dates bracket the lower reversal to between 66.275 +0.066/-0.053 Ma and 66.256 +0.034/-0.042 Ma (Fig. 1B), confirming that it is C30n/C29r. reversal Using our stratigraphic age model and an elevation of 400 m, we can further refine the date of this reversal 66.266 +0.060/-0.049 Ma. This date is consistent with several previous estimates for this reversal, including a U-Pb zircon date from a transitional polarity basalt flow in the Western Ghats (Schoene et al., 2019); sanidine <sup>40</sup>Ar/<sup>39</sup>Ar dates from interbedded tuffs in the terrestrial Hell Creek Fm. in Montana (Sprain et al., 2015); and astrochronologic estimates from the late Cretaceous/early Paleogene marine sedimentary sections in northern Spain (Batenburg et al., 2012), ODP Hole 762C (Thibault et al., 2012), ODP Holes 1267B, 762C, and DSDP 525A (Husson et al., 2011), and ODP Hole 1262 (Barnet et al., 2019). Notably, however, our date is 170 kyr younger than a recent estimate for the C30n/C29r reversal based on U-Pb zircon dates of ash beds within the Denver Basin, Colorado (Clyde et al., 2016). This discrepancy may be the result of variations in sediment accumulation rates within the terrestrial Denver Basin, because the Clyde et al. (2016) study did not bracket the C30n/C29r reversal, but rather projected their stratigraphic age model down-section over  $\sim$ 30 m. Regardless, the convergence of datasets pointing to a date of 66.266 +0.060/-0.049 Ma provides another important tie point with which to compare the eruptive history of the Deccan LIP to global paleoenvironmental records.

#### 4.2. Correlation to the Western Ghats

Most studies of the Deccan LIP have attempted to correlate volcanic sections over broad areas of the province using basalt geochemistry (e.g., Lightfoot et al., 1990; Mitchell and Widdowson, 1991; Khadri et al., 1999; Jay and Widdowson, 2008). However, there are limits to this approach (e.g., Kale et al., 2019). For example, there is the possibility that the province is composed of multiple eruptive centers that each records a distinct geochemical evolution through time (e.g., Kale et al., 2019). Correlation based on time provides a more robust means to reconstruct eruptive histories, obviating the need for assumptions regarding geochemical homogeneity of coeval lavas across the province.

Fig. 2 shows a correlation scheme for the Malwa Plateau and Western Ghats using U-Pb zircon dates from this study and from Schoene et al. (2019). Importantly, our dates demonstrate that emplacement of the Narmada Fm. is temporally correlative with emplacement of the Jawhar Fm. (Fig. 2). However, the oldest date for the Narmada Fm. is ~60 kyr older than the oldest date in the Jawhar Fm., extending initiation of the first eruptive pulse within the Deccan LIP to at least 66.358 +0.065/-0.070 Ma. The overlying Manpur Fm. is temporally correlative with the remainder of



**Fig. 2.** Stratigraphic correlation diagram of the southern Malwa Plateau basalts and the basaltic stratigraphy of the Western Ghats. U-Pb zircon age constraints are shown as bars color-coded to correspond to their respective formations and are derived from Bayesian Markov Chain Monte Carlo eruption (e.g., Keller et al., 2018) and stratigraphic (e.g., Schoene et al., 2019) age models derived from the distribution of zircon ages in each sample and their relative stratigraphic positions. All of the dates are from either Schoene et al. (2019) or this study. A combination of paleomagnetic and U-Pb zircon geochronologic constraints suggest that the stratigraphy of the southern Malwa Plateau basalts record the same four eruptive pulses as outlined in Schoene et al. (2019).

the Kalsubai and Lonavala subgroups (Fig. 2). We currently have no direct age constraints for the Mhow, Satpura, or Singachori Fms. However, if the upper magnetic reversal in the Malwa Plateau represents the C29r/C29n reversal, then the Mhow and Satpura Fms. are constrained to be younger than 66.137 +0.112/-0.170 Ma and older than the U-Pb date for the C29r/C30n reversal of 65.631 +0.053/-0.030 Ma (Figs. 1 and 2; Schoene et al., 2019). The Singachori Fm. is exposed as only erosional remnant at high elevations on the southern margin of the Malwa plateau, but must be temporally correlative to the Mahabaleshwar Fm. in the Western Ghats since both formations record the C29r/C30n magnetic reversal (Fig. 2: Khadri, 2003). Given these constraints, we suggest that the most parsimonious correlation is that the Mhow, Satpura, and Singachori Fms. record the second, third, and fourth eruptive pulses of Schoene et al. (2019) and are temporally correlative to the Poladpur, Ambenali, and Mahabaleshwar Fms., respectively. However, these temporal correlations should remain speculative until they are directly tested with additional geochronologic data. Regardless, the Malwa Plateau basalts must contain one or more hiatuses in eruption, or at least extreme changes in eruption rate, as a constant rate would predict a much thicker section of basalt between the Manpur Fm. and the C29r/C30n magnetic reversal. The alignment of U-Pb zircon eruption dates between the Narmada and Manpur Fms. with the first pulse of magmatism in Schoene et al. (2019); the requirement for hiatuses or extreme changes in eruption rate within the section; and the preliminary temporal correlations with the rest of the Western Ghats stratigraphy, all suggest that the same pulses identified in Schoene et al. (2019) are likely present in the Malwa plateau but with different relative thicknesses. Moreover, these observations are not compatible with an interpretation of near-constant Deccan eruptions (Sprain et al., 2019).

A single  ${}^{40}\text{Ar}/{}^{39}\text{Ar}$  date of a plagioclase separate from the Malwa plateau has been used to suggest that eruptions in this area started by 67.12  $\pm$  0.44 Ma ( $2\sigma$ : Schöbel et al., 2014). This sample was collected from a different part of the plateau than our study area, and is thus difficult to directly compare with our results. The contact between basalt and basement is not exposed in our study area along the Narmada River, and >100 m section of

basalt is present beneath our stratigraphically lowest sample. Thus, an older history of eruptions is possible. However, we suggest that any earlier eruptions within the Deccan LIP were volumetrically minor, because our age constraints for the onset of the first pulse of magmatism align with the initial downturn of <sup>187</sup>Os/<sup>188</sup>Os in ocean sediments globally (Fig. 3A).

#### 4.3. Implications for the architecture of the Deccan eruptive edifice

An understanding of the extent of Deccan eruptions through time is necessary to confidently constrain eruption volumes and tempos. Existing models have focused on either a single source (Model A in Mitchell and Widdowson, 1991; Jay and Widdowson, 2008); a southward-migrating central eruptive center, consistent with the passage of India over the Reunion hotspot (Devey and Lightfoot, 1986; Model B in Mitchell and Widdowson, 1991); or multiple eruptive centers that are distinct in time and space throughout the province (e.g., Kale et al., 2019). Our new data from the Malwa Plateau provide an important test for each of these models.

The temporal correlation between the Malwa Plateau basalts and the stratigraphy of the Western Ghats suggests that they record the same or a similar history of pulsed eruptions (Fig. 2). The continuity of these pulses over large regions is consistent with, but does not require, sustained volcanic activity from a central eruptive center, rather than several centers that were widely dispersed and behaving independently (e.g., Kale et al., 2019). Future studies of basalt geochemistry, morphology, and flow direction placed within a robust temporal framework will be able to assess whether coeval magmas across the province share the same petrogenetic and eruptive history or reflect coincidental eruptions from multiple distinct centers.

Assuming that the shared history between the Malwa Plateau and Western Ghats reflects the behavior of a central eruptive center, we can assess whether the basalts on the northern margin of the province are consistent with a center that remained stationary or migrated to the south through time. Existing maps from the southern Deccan LIP show southward attenuation of the formations of the Western Ghats and a consistent onlapping of flow



**Fig. 3.** A) Variation in benthic foraminifera  $\delta^{18}$ O and  $\delta^{13}$ C from ODP site 1262 (Barnet et al., 2019) compared to changes in  $^{187}$ Os/ $^{188}$ Os in ocean sediments from both outcrop and ODP core (Robinson et al., 2009; Ravizza and VonderHaar, 2012) and the timing and duration of the four eruptive pulses identified for the Deccan LIP by Schoene et al. (2019) and this study defined by the 95% credible intervals of the youngest and oldest U-Pb zircon eruption age in each pulse. The uncertainty of the U-Pb zircon dates for the C29r/C29n reversal (Schoene et al., 2019), Chicxulub impact (Clyde et al., 2016; Schoene et al., 2019), and C30n/C29r reversal (*this study*) are also shown. B) Schematic cartoon showing the proposed development of the Deccan LIP as a series of shield volcanoes with eruptive centers progressively migrating to the south through time (Devey and Lightfoot, 1986; Model B in Mitchell and Widdowson, 1991) and corresponding to the four (or five) eruptive pulses identified within the province. We speculate that the first Deccan eruptive pulse erupted through volatile-rich Narmada-Tapti rift basin sedimentary rocks, while subsequent pulses erupted through only granitic and gneissic basement.

fields to the south (Devey and Lightfoot, 1986; Model B in Mitchell and Widdowson, 1991). These observations are consistent with either of the proposed eruptive histories. However, both possibilities make specific predictions about the geometry of coeval flows along the north of the province. A central eruptive center should show symmetry with the map patterns in the south, with onlapping formations through time northward, while a migrating eruptive center should show progressive offlapping of formations as it moved southward.

The recognition that the  $\sim$ 460 m thick Narmada and Manpur Fms. are temporally correlative to the Lonavala and Kalsubai subgroups of the Western Ghats provides a test of coeval magma thickness in both the northern and southern parts of the province. Using the exposure limit of these two subgroups in the southern Western Ghats and assuming a stationary eruptive center near Nashik (e.g., Model A in Mitchell and Widdowson, 1991), coeval flows are predicted to be thin to nonexistent in the Malwa Plateau. The presence of thick, temporally correlative flows to these subgroups is, instead, consistent with a southward-migrating eruptive center (e.g., Devey and Lightfoot, 1986; Model B in Mitchell and Widdowson, 1991).

Fig. 3B shows a cartoon of how we envision the geometry of the province, with each eruptive pulse identified in Schoene et al. (2019) forming a shield volcano as the eruptive center migrated progressively to the south. It remains unclear if the undated Panhala and Desur Fms. (Fig. 2) represent a fifth eruptive pulse or are part of the eruptive pulse that includes the Mahabaleshwar Fm. Regardless, this architecture needs to be considered in future Deccan volume estimates. A first order consequence of this geometry

is that the relative volume of the first pulse of magmatism is significantly increased over the estimates presented in Richards et al. (2015).

Richards et al. (2015) calculated the total volume of the province by approximating each formation as a cone with a height equal to the maximum exposed thickness and diameters of 250 km for the Kalsubai and Lonavala subgroups and the Panhala and Desur Fms. and 600 km for the Poladpur, Ambenali, and Mahabaleshwar Fms. This approach results in a total volume of 602,700 km<sup>3</sup>, of which 23% belongs to the formations that compose the first eruptive pulse of Schoene et al. (2019). Recalculating the LIP volume using the same method but treating each pulse (including a hypothetical fifth pulse consisting of the Panhala and Desur Fms.) as a cone with a radius of ca. 440 km, similar to Model B in Mitchell and Widdowson (1991), gives a total volume of 737,900 km<sup>3</sup>, with 59% erupted during the first pulse. These revised volumes suggest that both Sprain et al. (2019) and Schoene et al. (2019) underestimated the eruption rates associated with the Kalsubai and Lonavala subgroups by using the Richards et al. (2015) model and highlight the importance of stratigraphic and geochronologic controls in constraining the eruptive volume of LIPs. The discrepancy also suggests that any conclusions drawn from modeling the climate response to the emplacement of the Deccan LIP (e.g., Hull et al., 2020) should be tempered by the uncertainties associated with each volume model, including variations in the volume of offshore basalt, in the volume of basalt that was eroded following eruption, and in the stratigraphic architecture that arises from multiple and/or migrating eruptive centers.

#### 4.4. Relationship to Late Maastrichtian climate instability

A 200 kyr period of  $\sim$ 2.5–8 °C warming and subsequent cooling before the end Cretaceous extinction, named the Late Maastrichtian warming event, has been inferred on the basis of  $\delta^{18}$ O in benthic foraminifera (Fig. 3A: e.g., Li and Keller, 1998; Barnet et al., 2018), pedogenic carbonate (Nordt et al., 2002), and bivalve shells (Petersen et al., 2016), as well as changes in leaf morphology (Wilf et al., 2003). The onset of this excursion is temporally correlative with the first pulse of Deccan volcanism identified by Schoene et al. (2019) and the initial downturn in oceanic <sup>187</sup>Os/<sup>188</sup>Os toward more radiogenic values (Fig. 3A: Robinson et al., 2009; Ravizza and VonderHaar, 2012), suggesting a potential causal relationship between the two events. Direct CO<sub>2</sub> emissions from basalt are unlikely to cause this magnitude of warming, except at extreme eruption rates (e.g., Self et al., 2006), which is difficult to reconcile with the likely longer duration and lower eruption rates inferred from this first eruptive pulse (Schoene et al., 2019). More puzzling is the observation that the other, shorter, and possibly more intense pulses of Deccan volcanism did not lead to records of warming on timescales >100 kyr (Fig. 3A). One possibility is that the second eruptive pulse was associated with shorter-term warming preserved in only a few locations due to widespread hiatuses in sedimentation (e.g., Punekar et al., 2014). Moreover, there is the possibility that shorter-term (<100-kyr) warming events occurred throughout the earliest Paleogene (e.g., Coccioni et al., 2010; Lyson et al., 2019). Nevertheless the temporal correlation of these proposed warming events with eruption of the Deccan LIP is not fully constrained and their global significance remains debated due to the absence of correlative  $\delta^{18}$ O excursions in benthic foraminifera (see discussion of DAN-C2 event in Barnet et al., 2019).

Our observation that the dominant volume of erupted material in the northern part of the Deccan LIP is temporally correlative with the first pulse of volcanism may provide a means to explain the Late Maastrichtian warming event. First, the increased areal extent of the basalts that compose the first eruptive pulse suggests a much larger eruptive volume than the estimates used in Schoene et al. (2019). Second, an inferred northern eruptive center for the first magmatic pulse would place it within the Narmada-Tapti Rift (Fig. 3B), where dike swarms have been previously documented (e.g., Vanderkluysen et al., 2011) and a gravity anomaly suggests large volumes of mafic rock in the middle to lower crust (Verma and Banerjee, 1992). This rift is well characterized in seismic lines through the northern part of the province and consists of 1.5-2 km of sedimentary rock (Fig. 1C: Kaila, 1988; Verma and Banerjee, 1992). These rocks can be traced laterally into sedimentary sequences on the margin of the Deccan LIP that record terrestrial Carboniferous to Cretaceous shallow marine sedimentation, including abundant Carboniferous and Permian coal (Veveers and Tewari, 1995). We suggest that initial eruption of Deccan basalts through the Narmada-Tapti Rift may have provided a critical source of greenhouse gases via metamorphism of organic rich sedimentary rock. This mechanism is widely invoked to have driven catastrophic CO<sub>2</sub> and CH<sub>4</sub> emissions during in late Permian during emplacement of the Siberian Traps (e.g., Svensen et al., 2009); in the early Jurassic during emplacement of the Karoo-Ferrar basalts (e.g., Svensen et al., 2007); and at the Paleocene-Eocene Thermal Maximum (PETM) during emplacement of the North Atlantic Igneous Province (e.g., Svensen et al., 2004). A relatively small  $-0.6\% \delta^{13}$ C excursion in benthic foraminifera during the late Maastrichtian warming event (Fig. 3A) has been used to argue against volumetrically significant addition of isotopically light organic ( $\delta^{13}C = -25\%$ ) carbon (Barnet et al., 2018). However, quantifying the expected change in global  $\delta^{13}$ C due to emission of these gases over the 200 kyr event duration depends on several unconstrained factors including the rate of organic matter burial in carbon sinks, the ratio of organically-derived greenhouse gases relative to volcanic CO<sub>2</sub> ( $\delta^{13}C = -6\%$ ), and the rate of greenhouse gas emission. Thus, the relatively small  $\delta^{13}C$  excursion observed during the Late Maastrichtian warming event (Fig. 3A) may not be sufficient to preclude protracted addition of isotopically light carbon during this time.

If our hypothesis is correct, southward movement of the Deccan eruptive center off of the Narmada-Tapti rift and onto bedrock entirely composed of Precambrian granite and gneiss minimized this extra source of greenhouse gas and may explain the more subdued climate proxy record associated with later pulses of Deccan volcanism. In this case, if the second pulse of the Deccan LIP played an important role in climate instability immediately predating the Chicxulub impact and mass extinction, either a) evidence for shorter-term ( $\leq$ 100 kyr) global greenhouse conditions is not preserved in most studied stratigraphic sections due to sedimentary erosion and/or hiatuses, b) an efficient silicate weathering feedback buffered atmospheric greenhouse conditions, and/or c) the effect of short-term global cooling from sulfur aerosols may have been an important contributor to environmental decline and accelerated extinction rates.

#### **Declaration of competing interest**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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#### Appendix A. Supplementary material

Supplementary material related to this article can be found online at https://doi.org/10.1016/j.epsl.2020.116249.

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#### MASS EXTINCTION

## **U-Pb** constraints on pulsed eruption of the Deccan Traps across the end-Cretaceous mass extinction

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Temporal correlation between some continental flood basalt eruptions and mass extinctions has been proposed to indicate causality, with eruptive volatile release driving environmental degradation and extinction. We tested this model for the Deccan Traps flood basalt province, which, along with the Chicxulub bolide impact, is implicated in the Cretaceous-Paleogene (K-Pg) extinction approximately 66 million years ago. We estimated Deccan eruption rates with uranium-lead (U-Pb) zircon geochronology and resolved four high-volume eruptive periods. According to this model, maximum eruption rates occurred before and after the K-Pg extinction, with one such pulse initiating tens of thousands of years prior to both the bolide impact and extinction. These findings support extinction models that incorporate both catastrophic events as drivers of environmental deterioration associated with the K-Pg extinction and its aftermath.

ontinental flood basalt provinces are characterized by eruption of >1 million km3 of basalt over a period of <1 million years (1, 2), representing the largest volcanic events on Earth. Four of the five most severe Phanerozoic mass extinctions [~541 million years (Ma) ago to the present] coincided with emplacement of one of these provinces (3). Although the temporal link between flood basalts and extinctions is well established, the mechanisms by which eruptions drive extinction are poorly understood (4). Two models of environmental change from volcanic activity relate to eruptive volatile emissions (1, 4). The first is volcanogenic CO2 release, with associated global warming, ocean acidification, and carbon cycle disruption. The second is SO<sub>2</sub> injection into the stratosphere and its conversion to sulfate aerosols, causing global cooling, acid rain, and ecosystem poisoning (5). The predicted time scales of these perturbations contrast sharply. The emission of SO<sub>2</sub> from a single eruption would produce years of cooling, whereas accumulated greenhouse warming from CO<sub>2</sub> can be sustained for many thousands to tens of thousands of years. Testing the effects of this interplay on ecosystems thus requires precisely calibrated volcanic eruption rates that can be correlated to high-resolution climate proxy and biostratigraphic data.

We applied U-Pb zircon geochronology to construct a precise temporal record of eruption

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within the Deccan Traps volcanic province, India (Fig. 1). The province is temporally correlated to the K-Pg mass extinction, in which roughly three-fourths of life on Earth was eradicated, including non-avian dinosaurs (6). Previous attempts to constrain eruption rates were limited by poor stratigraphic coverage and/or high analytical uncertainties (7-12). We used U-Pb geochronology by isotope dilution-thermal ionization mass spectrometry (ID-TIMS) (13), which provides analytical uncertainties  $(\pm 2\sigma)$  as low as 40,000 years (40 ka) for individual dated zircons. Our sampling covers the nine major Deccan formations in the Western Ghats, where the most voluminous (>90% total volume) and complete Deccan exposures are preserved (14-17) (Fig. 1). We sampled both coarse-grained basalts and sedimentary beds between basalt flows that infrequently contain zircon-bearing volcanic ash (11) (fig. S1). These beds, locally termed "redboles," range from oxidized volcaniclastic material with visible lithic fragments and phenocrysts to paleosol-type horizons produced by in situ weathering of flow tops (18, 19). Of 141 sampled redboles and coarse-grained basalts (Fig. 1 and figs. S1 and S2), 23 redboles and one basalt sample yielded sufficient zircon (≥5 crystals) to estimate an eruption age, including four distinct bole horizons and one basalt previously presented by Schoene et al. (11). Pristine volcanic crystal morphology indicates minimal transportation or reworking of zircon in a sedimentary environment. Consequently, we inferred that this volcaniclastic, zirconbearing material was incorporated into redboles as air fall tuff, consistent with some redboles containing a high-SiO<sub>2</sub> (nonbasaltic) component (19), and that these zircons provide a robust means for dating Deccan eruptive stratigraphy.

To estimate the eruption date and associated uncertainty for each sample, we developed an approach using Bayesian statistics to account for the probability distribution of zircon dates and their analytical uncertainties (20) (fig. S6). Although we considered alternative data interpretations (13), they do not affect the conclusions of this study. Twenty-one of 24 dated horizons are from five stratigraphic sections along prominent roads in the Western Ghats, providing complete coverage of the upper four Deccan formations (Fig. 1 and figs. S1 and S2). The remaining three samples span the lower five Deccan formations, where redboles are rare and less likely to contain zircon.

When compiled into a composite stratigraphic section (Fig. 1), almost all samples follow anticipated "younging-up" temporal order based on the independently defined regional stratigraphy (14-17) (figs. S2 and S7). The exception is the Katraj Ghat south of Pune city, where two samples from what was mapped as upper Poladpur Formation are ~100 ka younger than samples near the Poladpur-Ambenali contact in other sections. To resolve this discrepancy, we placed the Poladpur-Ambenali contact in the Katraj Ghat section as ~100 m lower than previously mapped. This simple adjustment does not violate geochemical or geological observations in the stratigraphy, as the Poladpur-Ambenali contact is geochemically transitional in published datasets (14). Furthermore, our placement of the contact is consistent with geochemical studies of the nearby Sinhagad Fort section suggesting that the Poladpur Formation is relatively thin just south of Pune (14).

To further refine the composite stratigraphic age model, we used a Bayesian Markov chain Monte Carlo (MCMC) model in which stratigraphic superposition is imposed on U-Pb zircon dates (13, 21) (Fig. 1). The result is a deposition age estimate for each dated horizon, incorporating dates from all beds above and below each sample to produce an internally consistent age model (Fig. 1). The accuracy of refined age estimates depends solely on sample placement in proper stratigraphic order and is independent of samples' exact stratigraphic heights.

To calculate volumetric eruption rates through the Deccan Traps, we adopted the volume model of Richards et al. (22), in which units of the Wai subgroup (i.e., the Poladpur, Ambenali, and Mahabaleshwar Formations) were interpreted as more voluminous than is apparent from their proportionate thickness in the Western Ghats. This assertion carries nontrivial uncertainties, but we believe it is justified given the correlation of these formations to basalt flows on the province's periphery, including massive flows that traveled ~1000 km to India's eastern shore (23, 24). Although different volume models produce changes in the magnitude of calculated eruption rates, the timing of peak eruption rates is unaffected by either the volume model or the interpretation approach of the zircon data (13) (figs. S8 and S9). Additional uncertainty relates to the unconstrained mass and

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age of Deccan basalt that is currently submerged and inaccessible off India's western shore. We consider this uncertainty to be intractable because current volume models cannot account for this mass component of the province. Consequently, all eruption rates are likely minimum estimates, although we also cannot assess whether the offshore component erupted during the same time intervals as that of the Western Ghats.

We converted our age model into a probabilistic estimate of volumetric flux of basaltic lava using outputs from the MCMC algorithm (Fig. 2). Our results showed that the Deccan Traps erupted in four high-volume events, each lasting ≤100 ka, separated by periods of relative volcanic quiescence. The first event corresponded to the eruption of the lowermost seven formations from ~66.3 to 66.15 Ma ago; the second to the Poladpur Formation from ~66.1 to 66.0 Ma ago; the third to the Ambenali Formation from ~65.9 to 65.8 Ma ago; and the fourth and final to the uppermost Mahabaleshwar Formation, from ~65.6 to 65.5 Ma ago.

Our Deccan eruption model (Fig. 2) constrains the volcanic tempo with high resolution, providing a means to correlate eruption records with biostratigraphic and climate proxy data across the K-Pg extinction. Our model places the second pulse of Deccan volcanism (Poladpur Formation, 66.1 to 66.0 Ma) as slightly predating a published U-Pb zircon date for the K-Pg boundary (KPB), defined as the Ir anomaly and associated fallout from the Chicxulub impact, within the Denver Basin, Colorado (25). For consistency, we applied the Bayesian approach to that dataset (25) to estimate a date of 66.016 ± 0.050 Ma ago for the KPB [95% credible interval, internal uncertainties only (13)]. Comparison of our data with recently published <sup>40</sup>Ar/<sup>39</sup>Ar geochronology from the Deccan Traps and the Chicxulub impact (12, 26) is currently not possible at the necessary level of precision given systematic bias between the two dating methods, primarily related to uncertainty in ages of <sup>40</sup>Ar/ <sup>39</sup>Ar fluence monitors and the values of the <sup>40</sup>K decay constant and physical constants (13). Assuming that the Chicxulub impact coincides exactly with the main phase of extinction, the MCMC model outputs from our Deccan data demonstrate a ~90% probability that the Poladpur Formation eruption pulse began tens of thousands of years before the K-Pg mass extinction event.



Fig. 1. Stratigraphy, sampling transects, and U-Pb age model for the Deccan Traps. (A) Elevation map of study location in the Western Ghats, India. A black segmented line denotes cross section X-X' as shown in (B). Sampling transects are located by colored dots. (B) Geologic cross section through the field area, with sample locations indicated. Different basalt formations in the Deccan Traps are color-coded to the stratigraphic column in (C). Cross section is based on previous work (14–17), modified according to our geochronology. (C) Volumetric stratigraphic column and magnetic

chrons of the major formations of the Deccan Traps (22, 45). The plotted sample heights ("RB" sample prefix omitted) are based on the composite stratigraphic section compiled in fig. S2. The age model for the Deccan Traps, based on our U-Pb geochronology, is shown with 95% credible intervals. Horizontal gray bars indicate eruption ages derived from populations of zircon dates from each horizon; black horizontal bars show dates refined from the stratigraphic Bayesian model. The vertical gray-shaded bar shows an age for the Chicxulub impact (25).

## Fig. 2. Eruption rate model for the Deccan Traps, based on U-Pb geochronology.

(A) Results from the MCMC algorithm used to generate the age model in Fig. 1, converted to a probabilistic volumetric eruption rate for the Deccan Traps shown with contours up to 68%credible intervals. The U-Pb date for the Chicxulub impact is the same as in Fig. 1. Total global volcanic productivity (~3 to 4 km<sup>3</sup>/year) includes mid-ocean ridges and volcanic arcs (28). (B) Compilation of proxy records from ODP cores and outcrops. Upper data points are  $\delta^{18}$ O of species-specific benthic foraminifera from ODP 525 (46), ODP 1262 (36), and ODP 1209 (37); VPDB, Vienna PeeDee Belemnite standard. Temperature is calculated for benthic foraminifera Nuttallides truempyi in ODP 1262 (36). Osmium isotopic records come from bulk carbonate from both ODP cores and outcrop (40, 41). Age models are described in (13).



The K-Pg extinction preserves the only known mass extinction that coincides with both a large igneous province and a bolide impact. As such, several hypotheses have been forwarded in which the impact triggered or modulated volcanic eruptions. Although the most recent iteration of this hypothesis concedes initiation of Deccan eruptions several hundred thousand years before the impact, it proposes that impact-induced seismicity increased eruption rates in the Deccan Traps and at mid-ocean ridges through evacuation of preexisting magma chambers in the upper mantle and lower crust (12, 22, 27). It is unlikely that our Deccan eruptive history is consistent with this model, given the high probability that the Poladpur pulse began before the impact by tens of thousands of years, followed by an eruption hiatus of ≤100 ka after the impact.

Estimates for the entire volcanic flux on Earth today are 3 to  $4 \text{ km}^3$ /year (28), indicating on average a doubling in global volcanic activity for ≤100 ka during each of the four highvolume Deccan eruptive events, but requiring periods of >5 to 10 times the global average. In fact, groups of flows within the Poladpur and Mahabaleshwar Formations, each potentially comprising >50,000 km<sup>3</sup>, lack secular evolution in paleomagnetic poles, suggesting eruption over decades to centuries (29). Such high eruption rates of >1000 km<sup>3</sup>/year are permitted by our U-Pb geochronology, requiring hiatuses of hundreds to thousands of years within our resolved pulses so as not to exceed total volume estimates. In addition to being consistent with brief but extreme eruption rates, our data demonstrate that the Deccan Traps erupted in pulses with durations of ~100 ka, providing insight into tempos of melt production and/or transport in the upper mantle and lower crust (30, 31).

Our eruption rate model is a first step toward robustly evaluating the environmental impacts associated with Deccan volcanism. The most commonly cited contributors to environmental change associated with flood basalts are CO<sub>2</sub> (warming), SO<sub>2</sub> (cooling upon conversion to sulfate aerosols), and chemical weathering of fresh basaltic material (cooling via CO<sub>2</sub> drawdown). For single continental flood basalt flows that erupt over a few decades, volcanic SO<sub>2</sub> has been modeled to drive cooling of 5° to 10°C for the duration of the eruption (5), after which acid rain rapidly removes sulfur compounds from the atmosphere. For persistent cooling over many thousands of years, therefore, hiatuses of only several decades between eruptions are required (5).

In contrast to SO<sub>2</sub>, the time scale of CO<sub>2</sub> removal from the ocean-atmosphere system is slow: ~1 ka, ~10 ka, and ~100 ka for mixing into the deep ocean, reaction with sediments, and removal by silicate weathering, respectively (32, 33). As a result, although climate effects during an eruptive event may be dominated by cooling associated with elevated sulfate aerosols, accumulation of volcanic CO<sub>2</sub> emissions can lead to net warming on intermediate time scales between eruptive events. On time scales of hundreds of ka to >1 Ma, weathering of fresh basalt has been modeled to result in net  $CO_2$  drawdown and cooling, especially if the basalt is at low latitudes, as were the Deccan Traps (34).

As an initial attempt to correlate our eruptive history with paleo environmental data, we used two proxy records across the K-Pg transition (Fig. 2B). Benthic for aminifera  $\delta^{18}$ O compositions indicate ~2° to 4°C of deep oce an warming over ~150 ka, beginning at the C30n-C29r magnetic reversal (~66.3 Ma ago), followed by cooling over ~150 ka prior to the KPB (35–37). It has also been argued on the basis of  $\delta^{18}$ O data from Elles, Tunisia, that renewed warming began tens of thous and of years before the KPB (38) (fig. S11).

Initial warming at ~66.3 Ma ago and a coeval increase in carbonate dissolution have been interpreted as resulting from volcanogenic CO<sub>2</sub> buildup and consequent ocean acidification (35), which our geochronology shows to have occurred during the initial pulse of Deccan eruptions. Warming curtailed toward the end of the first pulse, and cooling began before and continued through the initiation of the Poladpur Formation eruptions (Fig. 2). The extrusion of the voluminous Poladpur Formation may have resulted in short periods of SO<sub>2</sub>-driven cooling that could have continued to promote the overall cooling trend, but cooling for tens of thousands of years due to SO<sub>2</sub> emissions is difficult to sustain given the predicted short residence time of sulfate aerosol (1, 5). Alternatively, an increase in surface area of exposed basalt associated with the Poladpur eruptions is possible given current Deccan stratigraphic area/volume models (22), resulting in enhanced basalt weathering,  $CO_2$  drawdown, and continued global cooling during the tens of thousands of years before the extinction. If periods of cooling did result from sulfate aerosols during the Poladpur eruptions, the short intervals of temperature decrease could have slowed silicate weathering and associated  $CO_2$  drawdown, thus permitting  $CO_2$  buildup in the atmosphere that would be manifest between punctuated eruptions within the Poladpur Formation (*39*).

Testing whether basalt weathering was important leading up to the KPB is aided through study of the Os isotope system in marine carbonates because the ocean residence time of Os is short and basaltic  $^{187}$ Os/ $^{188}$ Os is low [0.1 (10)] relative to late Mesozoic seawater [0.6 (40)]. Published Os isotopic data from marine carbonates (40) show a marked decrease toward mantle values beginning at the onset of Deccan volcanism (Fig. 2B). A second downturn in <sup>187</sup>Os/<sup>188</sup>Os, beginning tens of thousands of years before the KPB, has been interpreted as a downward redistribution of extraterrestrial Os derived from the Chicxulub impactor (40, 41). However, this decrease is synchronous with the Poladpur eruption pulse and is thus also consistent with increased weathering of a more extensive Deccan basalt pile.

Post-extinction and post-Chicxulub benthic foraminifera  $\delta^{18}$ O and carbonate Os isotopic records do not covary with the Deccan eruption record. However, the Os record does not recover to the pre-Deccan <sup>187</sup>Os/<sup>188</sup>Os value either, perhaps indicating that a steady state was reached between basalt production and weathering despite continued eruptions. Regardless, the starkly different responses of O and Os isotope records during the post-extinction recovery require models that explicitly incorporate the effects of continued Deccan eruptions, the Chicxulub impact, and biotic effects on the carbon cycle in a world with devastated ecosystems.

Although the initiation of a massive eruptive pulse shortly before the Chicxulub impact and mass extinction supports a Deccan contribution to ecosystem collapse, much remains to be discovered as to how flood basalt magmatism contributes to mass extinctions. U-Pb geochronology has shown that, similar to the K-Pg extinction, the end-Permian (~252 Ma ago) and end-Triassic (~201 Ma ago) mass extinctions occurred on short time scales (< tens of ka), hundreds of thousands of years after the onsets of the Siberian Traps and Central Atlantic Magmatic Province flood basalt provinces, respectively (42–44). The eruptions and associated intrusive magmatism are presumed to have driven rapid extinction despite this time lag and the absence of bolide impacts. This lag between the onset of magmatism and extinction may be a result of highly nonlinear rates of magmatism, as documented here for the Deccan Traps. Continuing to study other flood basalt provinces will clarify the importance of eruptive and intrusive tempo in driving ecosystem collapse and extinction. Such an understanding of biosphere sensitivity and the importance of threshold processes during climate change is as relevant today as for these catastrophic events in Earth history.

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#### SUPPLEMENTARY MATERIALS

www.sciencemag.org/content/363/6429/862/suppl/DC1 Materials and Methods Supplementary Text Figs. S1 to S11 Tables S1 to S3

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# Science

## U-Pb constraints on pulsed eruption of the Deccan Traps across the end-Cretaceous mass extinction

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#### Two timelines for extinction

The Cretaceous-Paleogene extinction that wiped out the nonavian dinosaurs 66 million years ago was correlated with two extreme events: The Chicxulub impact occurred at roughly the same time that massive amounts of lava were erupting from the Deccan Traps (see the Perspective by Burgess). Sprain *et al.* used argon-argon dating of the volcanic ash from the Deccan Traps to argue that a steady eruption of the flood basalts mostly occurred after the Chicxulub impact. Schoene *et al.* used uranium-lead dating of zircons from ash beds and concluded that four large magmatic pulses occurred during the flood basalt eruption, the first of which preceded the Chicxulub impact. Whatever the correct ordering of events, better constraints on the timing and rates of the eruption will help elucidate how volcanic gas influenced climate.

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#### **EARTH HISTORY**

## U-Pb geochronology of the Deccan Traps and relation to the end-Cretaceous mass extinction

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The Chicxulub asteroid impact (Mexico) and the eruption of the massive Deccan volcanic province (India) are two proposed causes of the end-Cretaceous mass extinction, which includes the demise of nonavian dinosaurs. Despite widespread acceptance of the impact hypothesis, the lack of a high-resolution eruption timeline for the Deccan basalts has prevented full assessment of their relationship to the mass extinction. Here we apply uranium-lead (U-Pb) zircon geochronology to Deccan rocks and show that the main phase of eruptions initiated ~250,000 years before the Cretaceous-Paleogene boundary and that >1.1 million cubic kilometers of basalt erupted in ~750,000 years. Our results are consistent with the hypothesis that the Deccan Traps contributed to the latest Cretaceous environmental change and biologic turnover that culminated in the marine and terrestrial mass extinctions.

he Deccan Traps are a continental flood basalt province that comprise >1.3 million km<sup>3</sup> of erupted lavas and associated rocks (1) that reach a total thickness of  $\sim 3000$  m near the eruptive center in Western India (2, 3). Paleomagnetic data (4-7) combined with K-Ar and <sup>40</sup>Ar/<sup>39</sup>Ar geochronology of Deccan basalts (8, 9) have been interpreted to indicate that >90% of the eruptive volume was emplaced rapidly (<1 million years), coincident with the Cretaceous-Paleogene boundary (KPB). This temporal relationship has long led to speculation that Deccan volcanism had a major role in the end-Cretaceous mass extinction (10, 11), which saw the disappearance of nonavian dinosaurs and ammonoids, as well as major biotic turnovers in foraminifera, corals, land plants, reptiles, and mammals (12-15). However, age uncertainties from existing geochronology of the Deccan Traps (6, 8, 9) are larger than their estimated total duration, and thus the onset and duration of volcanism cannot be precisely compared to geologic, extinction, or environmental records from sedimentary sections spanning the KPB worldwide.

To better establish a high-resolution eruptive history of the main phase of Deccan volcanism, we sampled volcanic rocks from throughout the 10 formations that make up the Western Ghats (2, 3, 16) and dated them by U-Pb zircon geochronology using chemical abrasion–isotope dilution–thermal ionization mass spectrometry (CA-ID-TIMS) (17). Because zircon is rare in ba-

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saltic rocks, our sampling strategy targeted volcanic airfall deposits between basalt flows and high-silica and/or coarse-grained segregations within individual flows (fig. S1). The latter have been described previously in the lower half of the Deccan sequence (18), and we successfully extracted zircon from one such sample in the Jawhar Formation (Fm) (DEC13-30; Fig. 1). Zircon was also separated from three paleosol, or "redbole," horizons within the Ambenali and Mahabaleshwar Fms (samples RBP, RBE, and RBF; Fig. 1). These distinctive red horizons are interpreted to result from weathering of basalt during periods of volcanic quiescence (5). However, many also contain an evolved, high-SiO<sub>2</sub> volcaniclastic component (19), and we sampled these horizons to search for zircon-bearing volcanic ash that may have accumulated between basalt flows. Three additional zircon-bearing samples were collected from different intervals within a ~40-cm-thick green volcaniclastic bed in the Mahabaleshwar Fm. (DEC13-08, -09, and -10).

Each sample yielded a small number (typically <50) of euhedral zircon crystals with morphologies and internal zonation indicative of an igneous origin (fig. S3). Single grains were selected for analysis, photographed, pretreated, dissolved, and analyzed using CA-ID-TIMS (17). A subset of samples with an adequate number of grains was analyzed at both Princeton University (PU) and the Massachusetts Institute of Technology (MIT) to assess interlaboratory bias. Resulting  $^{206}\mathrm{Pb}/^{238}\mathrm{U}$ dates from individual zircons from each sample scatter outside of analytical uncertainty (all uncertainties reported at the  $2\sigma$  level; data shown in Fig. 2 and figs. S2 and S4 and reported in table S1) but show a similar spread in dates for samples analyzed at both MIT and PU. Given the excellent analytical reproducibility between laboratories (fig. S2), we discuss our results as a single data set below.

The spread in <sup>206</sup>Pb/<sup>238</sup>U dates from our individual samples cannot be attributed to analytical uncertainties alone (fig. S2), and we interpret this dispersion to result from either prolonged growth of zircon before eruption and/or incorporation of zircon from slightly older eruptions at the same vent. This phenomenon is due to the ability of zircon to retain radiogenic Pb at magmatic temperatures (>700° to 900°C) and can result in zircon dates within volcanic deposits that predate eruption by  $10^3$  to  $10^6$  years (20, 21). Given that our goal is to date the deposition of the volcanic ash, taking a weighted mean of all data from single samples is inappropriate and could bias our dates too old (21). Alternatively, the youngest zircon from each deposit may serve as a maximum age for deposition (16). However, this approach assumes that chemical abrasion has completely mitigated Pb loss (17) and could bias our dates too young if this assumption is not true.

To address these potential biases, we analyzed the trace element geochemistry of the dissolved zircon after routine ion exchange separation of U and Pb (17). By asserting that cogenetic zircons from an ashfall should have the same age and the same trace element signature, we identified the population of zircon from our data set that is amenable to statistical grouping (Fig. 2) while alleviating the concern that older, inherited zircon may bias weighted mean dates too old. We find that two or more zircons from each sample meet these criteria, and we calculate weighted means from those grains. Additionally, zircons from different samples have very different trace element signatures, supporting our interpretation that each dated horizon contains a distinct population of zircon with independent age information.

As a further means of refining our age model for the middle Ambenali-lower Mahabaleshwar Fms, we employed a Markov chain Monte Carlo analysis that imposes the law of superposition as a boundary condition (17). Given that stratigraphic horizons young upward, this Bayesian approach uses the <sup>206</sup>Pb/<sup>238</sup>U dates for each horizon derived above as priors and calculates new uncertainty distributions for each sample that maximize and evaluate the probability that stratigraphically higher beds are younger. Two of three samples from the composite ashbed fail this test (DEC13-08 and -09), and thus the date arising from the third sample (DEC13-10) is used as our best estimate for the deposition of that composite ashbed. Dates derived from the redbole horizons pass the superposition test, consistent with our interpretation based on grain morphology and geochemistry that zircon from these horizons derive from primary volcanic ashfall rather than, for example, eolian transport. The depositional ages presented in Fig. 2 are those from the Monte Carlo analysis, whereas several different interpretations of the geochronological data are presented in table S3.

Regardless of the method used for U-Pb age interpretation, our main conclusions remain unaffected. The U-Pb dates reported here have corresponding uncertainties that are one to two orders of magnitude smaller than previously

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published geochronology from the Deccan Traps and can thus resolve age differences between the base and top of the main eruptive phase. Using the dates from the lower- and upper-most samples (Fig. 1), we calculate a duration of  $753 \pm 38$ thousand years (ky) for an estimated 80 to 90% of the total eruptive volume of the Deccan Traps.

These data also calibrate the timing of magnetic polarity Chron 29r, which serves as a basis for global correlation of KPB sections. The C29r/ C29n reversal was previously identified within the lower Mahabaleshwar Fm (5), from which we collected samples RBP, RBE, and RBF. We use the  $^{206}\text{Pb}/^{238}\text{U}$  date of sample RBE of 65.552  $\pm$ 0.026/0.049/0.086 million years ago (Ma) ( $2\sigma$  uncertainties given here as internal only/with tracer calibration/with <sup>238</sup>U decay constant) as our best estimate for the age of C29r/C29n reversal because it was sampled from between two basalts with transitional polarity (Fig. 1) (5). The basal age of C29r is constrained by sample DEC13-30, which was collected from a basaltic segregation vein within the Jawhar Fm near the base of the main Deccan phase and yielded a <sup>206</sup>Pb/<sup>238</sup>U date of 66.288  $\pm$  0.027/0.047/0.085 Ma. From the same outcrop, Chenet et al. (9) reported transitional magnetic polarity just tens of meters beneath lavas containing reverse polarity (C29r); from this transitional horizon they reported a K-Ar date of 67.4  $\pm$  2.0 Ma. Though their date was interpreted to represent the C30r/C30n or C31n/C30r transition, with a long hiatus in eruptions represented in that section (9), our U-Pb date from DEC13-30 is not consistent with a hiatus of that magnitude. A simpler interpretation is that the transitional polarity basalts represent the C30n/C29r transition. Thus, the age difference between DEC13-30 and RBE of 736  $\pm$ 37 ky is also the length of C29r. An independent estimate for the length of C29r from cyclostratigraphic analysis of marine Ocean Drilling Program (ODP) cores and the Zumaia section, Spain, vields durations of 713 to 725 ky (22), in good agreement with our calibration based on U-Pb geochronology.

Our results also have implications for the age of the KPB and associated mass extinction event, as several estimates for the duration of the Cretaceous portion of C29r of 300 to 340 ky have recently been published based on cyclostratigraphy of magnetically and biostratigraphically calibrated ODP sections (23). Using our date for the C30n/C29r reversal and the average cyclostratigraphic estimate yields an age for the KPB of 65.968  $\pm$  0.085 Ma (including systematic uncertainties), which agrees well with a recently reported KPB age of 66.043  $\pm$  0.086 Ma (also including full systematic uncertainties) from <sup>40</sup>Ar/<sup>39</sup>Ar geochronology on tephras that bracket the terrestrial KPB near Hells Creek, Montana (24). Determining an age for the KPB by backcalculation from our estimate for the C29r/C29n reversal is more problematic in that estimates for the Paleogene portion of C29r based on cyclostratigraphy of the Zumaia KPB section range from 206 to 398 ky (22, 25). Regardless, the combination of our geochronologic data with cyclostratigraphic estimates effectively rules out any age for the KPB younger than 65.740  $\pm$  0.086 Ma.

Although the temporal relationship between large igneous provinces and mass extinctions is

well established (26), the potential kill mechanisms remain a subject of debate. Models of proposed drivers focus on volcanically sourced  $CO_2$ ,  $SO_2$ , and halogens, which can cause global warming and/or cooling on different time scales (27); acid



Fig. 1. Geography and stratigraphy of the Deccan Traps. (A) Aerial extent of the Deccan Traps colored in red. (B) Shaded relief map of study area in the Western Ghats. Major cities are indicated by white squares. Approximate locations of sampling transects and samples are indicated by colored triangles; transects are named at the bottom of the figure. The color bar (top right) shows elevation in meters, highlighting the escarpment where the best exposures of Deccan lavas occur. (C) Schematic cross section of Deccan lavas, from Chenet et al. (4), showing general southerly dip and younging to the south. Sampling transects are indicated by colored triangles. Colors correspond to the formations named in (D). (D) Composite stratigraphic section of the Deccan Traps in the Western Ghats, with approximate formation thicknesses shown. The geologic time scale is on the left, with the gray area corresponding to the unknown location of the KPB within the Deccan Traps. The geomagnetic polarity time scale is shown, with relevant chrons labeled and polarity indicated by black, white, or gray. The duration of C29r is the difference between the ages of DEC13-30 and RBE. U-Pb ages are shown on the right and are color-coded for sample type: black, segregation vein in basalt; red, volcanic material from paleosol; green, volcanic ashbed. Colored triangles correspond to sampling transects shown in (B) and (C). The date for KPB from Renne et al. (24) includes full systematic uncertainties. U-Pb age uncertainties are 2o and include internal uncertainties only; ages with full systematic uncertainties are ~±0.085 Ma. See text and table S3 for full uncertainty budget and Fig. 2 and table S1 for data.



**Fig. 2. U-Pb zircon CA-ID-TIMS geochronological data.** (**A**) Rank order plot of U-Pb data presented in this study, color-coded by sample type in Fig. 1 and with sample name next to data. Sample locations are shown in Fig. 1. The vertical axis indicates  $^{206}$ Pb/ $^{238}$ U date, and rectangle height corresponds to  $2\sigma$  uncertainties for single-crystal zircon analyses, with internal uncertainties only. MIT and PU indicate the laboratory used. Stratigraphic younging is shown from right to left. The horizontal gray band shows the date for KPB from Renne *et al.* (*24*). Small gray rectangles behind the data indicate the youngest zircons that were indicated to be cogenetic by comparing dates and geochemistry in (B),

from which weighted means were calculated. Dates indicated beneath sample names result from the Monte Carlo Markov chain simulation that uses weighted mean dates and imposes the law of superposition to arrive at our best estimates for the time of deposition of the dated horizons (17). Asterisks indicate zircon without trace element geochemistry. U-Pb data are given in table S1. (**B**) Lutetium/hafnium (Lu/Hf) ratios of the same volume of dated zircon, younging from right to left as in (A). Gray boxes indicate zircons determined to be cogenetic due to same age and geochemistry. The full geochemical data set is presented in table S2 and plotted in fig. S5.

rain and ozone reduction (28); and ocean acidification (29). Late Cretaceous records beginning near the C30n/C29r transition, and therefore near the onset of the main phase of Deccan volcanism, show a decrease in  $\delta^{18} O$  values of foraminifera (30) and morphological changes in fossil leaves (31) that are consistent with instabilities in global temperature. A two-stage decline in seawater <sup>187</sup>Os/<sup>188</sup>Os values initiating at the C30n/C29r reversal was interpreted to record weathering of the Deccan Traps, predating a second decline in <sup>187</sup>Os/<sup>188</sup>Os and a synchronous Ir spike that were attributed to the Chicxulub impact (32). Furthermore, biostratigraphic records show increased rates of biotic turnover in mammals, amphibians, land plants, and foraminifera through the Cretaceous portion of C29r preceding the peak extinction interval (12, 13, 15, 31, 33). Additional testing of the influence of the Deccan Traps on these records will require further determination of eruption tempos and hiatuses coupled with realistic estimates of volatile release from individual eruptive phases (34) that can be temporally linked to paleoenvironmental proxies. Our results are a critical part of this discussion as they are consistent with the hypothesis that environmental and ecological deterioration began with eruption of the Deccan Traps before the Chicxulub impact and the end-Cretaceous mass extinction. Therefore, both the Chicxulub impact and eruption of the Deccan Traps should be considered in any model for the extinction.

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#### SUPPLEMENTARY MATERIALS

www.sciencemag.org/content/347/6218/182/suppl/DC1 Materials and Methods Figs. S1 to S5 Tables S1 to S3 References (35-108)

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## Supplementary Materials for

# U-Pb geochronology of the Deccan Traps and relation to the end-Cretaceous mass extinction

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## This PDF file includes:

Materials and methods

- Figure S1: Sample locations and outcrop photos
- Figure S2. U-Pb geochronological and trace element geochemical data for natural zircon standard AUS-Z2.
- Figure S3. Photographs of zircons from this study
- Figure S4: Concordia plots for U-Pb ID-TIMS geochronological data.
- Figure S5: Trace element data for zircons dated in this study.

Table S1. U-Pb isotopic data

- Table S2. Zircon trace element
- Table S3. Age interpretation table

## **MATERIALS AND METHODS**

## S1 Sampling transects and sample descriptions

Samples presented in this study were collected from three stratigraphic sections exposed by road cuts (Fig 1, S1). Background information is given on these stratigraphic sections below, and this is followed by detailed sample descriptions.

## S1.1 Igatpuri section

This section contains the lowermost exposures of the Deccan traps beginning in the Jawhar Fm, whose base is not well-exposed but appears to lie unconformably on Archean-Phanerozoic rocks (*1*). There are good exposures from the Jawhar Fm through the middle Igatpuri Fm along the Nashik-Mumbai expressway between Igatpuri and Kardi. <sup>40</sup>Ar/<sup>39</sup>Ar (8) and K-Ar (9) geochronology, in addition to detailed geomagnetic transects (*4*) have been published from this section.

## **DEC13-30** (N19.62640°, E73.46407°, elevation 255 m)

This sample was collected from a roadcut within the lower Jawhar Fm  $\sim 200$  m below the contact with the Igatpuri Fm, defined by the Thalghat Giant Plagioclase Basalt, which is a regionally continuous stratigraphic marker bed. DEC13-30 is about 50 m stratigraphically above the lower most exposed Deccan units in the western Ghats (4, 9), accounting for west dipping normal faults that cause stratigraphic repetition when driving east-west on the west-sloping highway (4, 9). The outcrop is  $\sim$ 3-4 m high and composed of massive fine-grained basalt cross-cut by sinuous veinlets of coarser-grained friable material (Fig. S1). Bondre et al. (18) describe such veins in detail as features exclusively found in composite flows, i.e., large flows that are composed of many smaller pahoehoe flows that override or inflate liquid portions of slightly older eruptions. Segregation veins are often associated with or grade into vesicular pipes that pierce the adjacent flow, which would not be expected if the segregation veins were very late dikes or hydrothermal in origin (18). The often sinuous nature of the veins led Bondre et al. (18) to speculate that they formed while the flow was still mobile and were thus deformed. We therefore interpret these veins to have crystallized very shortly after emplacement of the host flow and magmatic zircons from these veins therefore very closely date eruption of the lava. DEC13-30 was collected from the same outcrop as samples S11 and S12 of Chenet et al. (4), which were analyzed for paleomagnetism and assigned a transitional polarity (i.e., during a normal to reverse polarity switch). This outcrop is ~10 m higher stratigraphically than sample NA03, dated by K-Ar at 67.4 $\pm$ 2.0 Ma (2-sigma) (9) and which also contained a transitional polarity. These samples are ~40 m stratigraphically below a transition to reverse polarity in the middle Jawhar Fm, assigned to chron C29r (4, 9). K-Ar and <sup>40</sup>Ar/<sup>39</sup>Ar dates within the reverse polarity Jawhar Fm range from 64.7 $\pm$ 1.8 and 65.7 $\pm$ 1.4 Ma (4, 9). Using updated knowledge of fluence monitors used for <sup>40</sup>Ar/<sup>39</sup>Ar geochronology (e.g., the age of the Fish Canyon sanidine of 28.02 Ma (*35*) versus 28.2-28.3 Ma (*36*, *37*)) could shift <sup>40</sup>Ar/<sup>39</sup>Ar dates older by 400-500 ka, whereas different values of the <sup>40</sup>K decay constant would not significantly change the K-Ar dates of Chenet et al. (9). Regardless of these changes, nearly every published K-Ar or <sup>40</sup>Ar/<sup>39</sup>Ar date from the Deccan traps overlaps within 2-sigma uncertainty of each other and with our U-Pb dates, so the ages of fluence monitors or chosen decay constant for <sup>40</sup>K from dates within the Jawhar Fm do not affect comparison to DEC13-30.

## S1.2 Sinhagad Fort section

This section is located on Sinhagad road, which exposes fresh basalt on roadcuts up the Sinhagad ghat ~20 km SW of Pune. The section covers ~700 m vertically through the Deccan traps from the lower Poladpur Fm through the base of the Mahabaleshwar Fm, though the best exposures are restricted to the Ambenali Fm where roadcuts expose basalt of variable freshness. Detailed geochemical sampling and petrologic observations can be found in Beane et al. (*3*), which forms the basis for stratigraphic correlation with the Mahabaleshwar section located 40 km to the south.

## **RBP** (N18.36173°, E73.77290°, elevation 1083 m)

This sample was taken from a redbole horizon in the middle Ambenali Fm ~160 m below the Ambenali-Mahabaleshwar Fm contact in the Sinhagad fort section SW of Pune (Fig. 1, S1). This section exposes the Bushe to the Mahabaleshwar Fms and has been described petrologically and geochemically in detail (3). The redbole is ~0.5 m thick and our sample combines material from the lower contact with weathered basalt through the upper paleosol layer (Fig. S1A), which evidently contains volcanic-derived material.

### S1.3 Mahabaleshwar section

This section contains fabulous exposures of the Poladpur through Mahabaleshwar Fms as the road cuts >1000 m downward through the escarpment of the western Ghats between the towns of Mahabaleshwar and Poladpur (Fig. 1C). Detailed unit descriptions in addition to high-resolution geochemical and isotopic sampling have been reported (*16*) and these data help define the unit boundaries in this section and provide a basis for stratigraphic correlation with the Sinhagad Fort section, from which RBP was sampled. Chenet et al. (*5*) presented high-resolution magnetostratigraphic data and some stratigraphic observations from this section, including the positions of redboles.

## **RBE** (N17.92550°, E73.62360°, elevation 999 m)

This sample was taken from a redbole horizon in the Mahabaleshwar Fm  $\sim$ 80 m above the Ambenali-Mahabaleshwar contact. This redbole is  $\sim$ 1.3 m thick and consists of red clay and silty horizons, and the entire horizon was sampled for geochronology (Fig. S1A). RBE was collected closest to samples MB17, -28, -07 and -09 of Chenet et al. (*12*), and is therefore within meters stratigraphically from the contact between lavas with transitional polarity and normal polarity of C29n.

## **RBF** (N17.91957°, E73.62745°, elevation 1064 m)

This redbole horizon was sampled ~60 m stratigraphically above RBE within the Mahabaleshwar Fm. It contains an ~20 cm thick red weathering horizon with ~50 cm of dense lithified clay material underlying the base of the overlying flow (Fig. S1A). It was collected from between samples MB11 and MB17 of Chenet et al. (5) and therefore from within a reverse polarity section representing C29n within 20 meters above transitional polarity lavas.

**DEC13-08, -09, -10** (N17.91671°, E73.62527°, elevation 1088 m) Three samples were collected from in or just below a ~30 cm thick green clastic horizon between two massive basalt flows (Fig. S1A). This green layer is unlithified, clay-rich and contains abundant clasts interpreted as phenocrysts and lithic fragments floating in a finer grained matrix. At the layer's top, the overlying basalt exhibits loadcast structures, generally interpreted as denser material sinking into less dense unlithified material. The base of the layer grades into brownish clay rich material and

altered basalt that is brecciated in places, though deep weathering obscures these contacts. This horizon is interpreted as a single or composite volcanic ashbed; numerous similar layers described throughout the Deccan are colloquially called greenboles. This particular greenbole appears on the stratigraphic column of Chenet et al. (5) between basalt samples MB08 and MB05, which are both normal polarity. Other greenboles are described throughout the region, though their origin is not likely always volcanic and may also be interpreted as paleosols similar to redboles (*38*). DEC13-08 was collected near the top of this ashbed which was dominated by fine-grained material and few phenocrysts; DEC13-09 was collected at the base of the ashbed in a coarser grained trough of crystal rich material ~40 m laterally from DEC13-08. DEC13-10 was collected from a brown clay rich layer ~10 cm below the base of the green ash from within a zone of broken up heavily altered basalt.

## S2 Methods: U-Pb geochronology

## S2.1 Zircon separation and imaging

Zircons were extracted from their host rocks at Princeton University (PU) and the Massachusetts Institute of Technology (MIT) by standard crushing, gravimetric- and magnetic-separation techniques using a Bico Braun "Chipmunk" Jawcrusher, discmill, Wilfley table or hand pan, methylene iodide, hand magnet, and Frantz isodynamic separator. For unlithified clay rich samples, the crushed material was either placed into a blender or an ultrasonicator to separate zircon from clay prior to density and magnetic separation. Zircons were picked in reagent-grade ethanol from the least magnetic mineral separate. Redbole samples (RBE, RBF, and RBP) and the segregation vein sample (DEC13-30) had very low zircon yield, so most or all were picked for analysis. The greenbole samples (DEC13-08, -09, -10) contained abundant zircon (>100). A subset of zircon from RBE, DEC13-09 and -10 were picked and sent to the other lab for duplicate analysis. Picked zircons were transferred in bulk to quartz crucibles, loaded in a muffle furnace, and annealed at 900 °C for 48 hours (*39*). Annealed grains were then transferred in ethanol into plastic petri dishes from which single grains were chosen for U-Pb geochronology. Selected grains were photographed and then transferred using stainless steel picking tools into separate 3-ml Savillex PFA hex beakers containing distilled acetone and taken to the clean lab.

## S2.2 U-Pb zircon ID-TIMS analysis

Zircons were rinsed in hex beakers in either distilled acetone or MQ  $H_2O$ . Single grains were loaded into 200  $\mu$ l Savillex "micro"-capsules with 100  $\mu$ l 29 M HF + 15  $\mu$ l 3 N HNO<sub>3</sub> for a single leaching step in high-pressure Parr bombs at 195 °C for 12 h to remove crystal domains affected by Pb loss (39, 40). Grains were rinsed post-leaching with 6 N HCl, MQ H<sub>2</sub>O, and 3 N HNO<sub>3</sub> prior to spiking with the EARTHTIME (<sup>202</sup>Pb-)<sup>205</sup>Pb-<sup>233</sup>U-<sup>235</sup>U tracer and addition of 100  $\mu$ 1 29 M HF + 15  $\mu$ 1 3 N HNO<sub>3</sub>. Zircons were then dissolved to completion in Parr bombs at 210 °C for 48 h. Dissolved zircon solutions were subsequently dried down, redissolved in 100 µl 6 N HCl and converted to chlorides in Parr bombs at 195 °C for 12 h, after which solutions were dried again and brought up in 50 µ1 3 N HCl. The U-Pb and trace element aliquots were then separated by anion exchange column chromatography using 50  $\mu$ l columns and AG-1 X8 resin (200-400 mesh, chloride form from Eichrom) (41) and dried down with a microdrop of 0.05 M H<sub>3</sub>PO<sub>4</sub>. The dried U and Pb aliquot was loaded in a silica gel emitter (42) to an outgassed, zonerefined Re filament. Isotopic determinations were performed using an IsotopX PhoeniX-62 TIMS at PU or a VG Sector 54 TIMS at MIT, with Pb analyses performed in peak-hopping mode on a Daly-photomultiplier ion counting detector. A correction for mass-dependent Pb fractionation was applied cycle-by-cycle, calculated from the deviation of measured <sup>202</sup>Pb/<sup>205</sup>Pb from the known tracer  $^{202}$ Pb/ $^{205}$ Pb (0.99924 ± 0.00054 (2 $\sigma$ )). A Daly-photomultiplier Pb dead time of 40.5 ns was used at PU, as determined by >100 measurements of NBS981 and NBS982 standards and using a least squares fit to measured data over the range of 100 kcps to 2.5 Mcps <sup>208</sup>Pb. Corrections for interfering isotopes under masses 202, 204, and 205 were made cycle-bycycle by measuring masses 201 and 203 and assuming they represent  $^{201}BaPO_4$  and  $^{203}Tl$  and using natural isotopic abundances to correct for <sup>202</sup>BaPO<sub>4</sub>, <sup>204</sup>BaPO<sub>4</sub>, <sup>205</sup>BaPO<sub>4</sub>, and <sup>205</sup>Tl. Though mass 201 peaks can be in the hundreds to thousands of cps, this correction is nearly negligible. More importantly, though peaks under mass 203 were typically <10 cps, higher abundances can have an important effect on the 205 peak. Despite the assumed accuracy of the correction, it was found at PU that very large 203 peaks often correlate with negatively discordant data. In such cases, these data are reported but not used in age interpretations (e.g., several analyses from sample DEC13-30 in this study). Typical ion yields for  $^{205}$ Pb were 100-200 kcps for  $\sim 20$  pg <sup>205</sup>Pb, sustainable for 3-4 hours of analysis (160-220 ratios collected) and typical <sup>206</sup>Pb/<sup>205</sup>Pb ratios were 0.1-2.0.

UO<sub>2</sub> measurements at PU were performed in static mode on Faraday cups with  $10^{12}$  ohm resistors using an oxide composition of <sup>18</sup>O/<sup>16</sup>O of 0.00205 (43, 44). Mass fractionation of U was determined cycle-by-cycle, calculated from the deviation of measured <sup>233</sup>U/<sup>235</sup>U from the known tracer <sup>233</sup>U/<sup>235</sup>U (0.995062 ± 0.000108 (2 $\sigma$ )) and assuming a <sup>238</sup>U/<sup>235</sup>U of 187.818 ± 0.045 (2 $\sigma$ ) for sample U (45). Typical U ion beams on 10<sup>12</sup> ohm resistors at PU were 200-1000 mV of <sup>233</sup>U at PU and 25-150 mV of <sup>233</sup>U on 10<sup>11</sup> resistors at MIT for ~1 ng <sup>233</sup>U, resulting in measurement precision of 0.004-0.02% (2 SE) for <sup>270</sup>U/<sup>265</sup>U ranging from 2 to 0.06

Data reduction was performed using the programs Tripoli and U-Pb\_Redux (46, 47) and the U decay constants of Jaffey et al. (48). All Pb<sub>c</sub> was attributed to laboratory blank with a mean isotopic composition determined by total procedural blank measurements, carried out separately for each analyst and dissolution set. The composition of the tracer used at both MIT and PU is that for ET2535 v. 3.0:

Composition	Value	$\pm 1\sigma$ (abs)
<sup>202</sup> Pb/ <sup>205</sup> Pb	0.99924	0.00027
<sup>204</sup> Pb/ <sup>205</sup> Pb	0.000105	0.000009
<sup>206</sup> Pb/ <sup>205</sup> Pb	0.00048	0.00017
<sup>207</sup> Pb/ <sup>205</sup> Pb	0.00043	0.00014
<sup>208</sup> Pb/ <sup>205</sup> Pb	0.00104	0.00033
<sup>233</sup> U/ <sup>235</sup> U	0.995062	0.000054
<sup>238</sup> U/ <sup>235</sup> U	0.00307993	4.0E-7
Concentration of <sup>205</sup> Pb (mol/g)	1.0312E-11	2.6E-14
Concentration of <sup>235</sup> U (mol/g)	1.0336E-9	2.6E-12

A thorough discussion of the Th-disequilibrium correction is given in section S6.1. Uncertainties in reported U-Pb zircon dates are reported at the  $2\sigma$  or 95% confidence level and exclude uncertainties in tracer calibration and decay constants unless otherwise noted.

In addition to semi-weekly measurements of Pb standards NBS981 and 982 to monitor Pb fractionation and ion counter deadtime, we generated U-Pb and trace element standard data for natural zircon standard AUS-Z2 (49) to monitor laboratory reproducibility and interlaboratory bias (Fig. S2). Data from PU yield a statistically significant cluster in concordia space with a weighted-mean  $^{206}$ Pb/ $^{238}$ U CA-ID-TIMS age of 38.905 ± 0.013/0.017/0.045 Ma

(MSWD = 1.0, n = 9/10;  $2\sigma$  uncertainties stated as internal/+tracer calibration/+decay constant); data from MIT give  $38.879 \pm 0.020/0.022/0.047$  Ma (MSWD = 1.6, n = 3/3). These values are in excellent agreement with the published age of  $38.8963 \pm 0.0044/0.012/0.043$  Ma (MSWD = 1.0, n = 12/12) (49).

### **S3 Methods: Zircon trace element geochemistry**

The trace element compositions of the same zircon fragments dated by ID-TIMS were characterized following the analytical protocol of Schoene et al. (50) at PU. Trace element washes isolated during U-Pb column chemistry were dried down in pre-cleaned 2.0 ml polypropylene vials (CETAC #SP5540) and redissolved in 1.0 ml 1.5 M HF + 0.1 M HNO<sub>3</sub> + 1 ppb Ir. Measurements were performed on a Thermo Fisher ELEMENT2 sector field-inductively coupled plasma-mass spectrometer (SF-ICP-MS) with a sample introduction system consisting of a CETAC Aridus II desolvation nebulizer + ASX-100 autosampler with an uptake rate of 100  $\mu$ l/min. Measured elements included Zr, Hf, Y, Nb, Ta, REEs, Pb, U, Th and Ir, with iridium monitored as an internal standard during mass spectrometry. The instrument was tuned in low resolution mode with an optimal signal intensity of 0.5-2 Mcps on the peak height (not the integrated signal) for 1 ppb Ir. A matrix-matched, gravimetric external calibration solution was prepared with the relative abundance of targeted elements representing that observed in natural zircon (e.g., Zr/Hf = 50). A dilution series was generated using this solution to cover the range of concentrations observed in unknowns (e.g.,  $[Zr] = 10^1 - 10^4$  ppb solution), which was then used to generate a concentration-intensity calibration curve for each trace element at the beginning of the analytical session. Samples and interspersed instrumental and total procedural blanks were analyzed in sets of 24 over  $\sim$ 3 h, with a line washtime of 120 s and uptake time of 90 s. Following data acquisition, solution concentrations were converted to stoichiometric concentrations in zircon by normalizing solution concentration data assuming all trace elements partition into the  $Zr^{4+}$  site in  $ZrSiO_4$ , where  $\Sigma Zr+Hf+...+Th = 497,646$  ppm. Uncertainties are reported at the 95% confidence level and include subtraction of the mean and standard deviation of blank measurements. Uncertainties in ratios were calculated using standard uncertainty propagation techniques and assuming uncorrelated uncertainties in measured elemental abundances. Trace element data were also produced for AUS-Z2 during the course of this study,

which show excellent reproducibility and identical concentrations on Zr, Hf, and the REE compared to published laser ablation ICPMS data (49) (Fig. S2).

#### **S4 Results: U-Pb geochronology**

Results from U-Pb geochronology for each dated sample are presented below. Because of very limited zircon yield, these grains were not imaged by cathodoluminescence prior to analysis. Photographs of nearly all the zircons were taken (except on a day when the camera was malfunctioning), and these are presented in Fig. S3. Concordia plots are in Fig. S4 and rank order plots are in Fig. 2. Given the complicated nature of this dataset, only qualitative descriptions are presented below, and the age interpretations are discussed section S6. All uncertainties are reported at the  $2\sigma$  level with internal errors only except where indicated. Age interpretations are presented in the main text as  $\pm X/Y/Z$  where X indicates internal uncertainties only for comparison with  $^{206}Pb/^{238}U$  dates from labs using the ET(2)535 tracer solution, Y includes tracer calibration uncertainties for comparison with other  $^{206}Pb/^{238}U$  dates determined with a different tracer solution, and finally Z includes full systematic uncertainties including decay constants for comparison with other radioisotopic dates or astrochronologically-determined timescales that may be pinned to radioisotopic dates other than the  $^{206}Pb/^{238}U$  system (*35, 51*).

### **DEC13-30**

This sample yielded <30 zircons with several morphologies. One was a population of subhedral, clear grains <200  $\mu$ m in length; 5 of these were sampled for analysis and were inherited, ranging in age from 160-1500 Ma (Table S1, Fig. S3). A second population of grains were euhedral and stubby but coated in brown material (perhaps hematite; Fig. S3) such as to conceal these grains upon first viewing. Blocky fragments of zircon containing euhedral crystal terminations were also in this population and were often adhered to clayey material likely deriving from weathered glass or feldspars. Less than twenty of these zircons were recovered and 11 were successfully discordant, and these analyses also contained abnormally high count rates under masses 201 and 203, which typically indicates that interferences other than BaPO<sub>4</sub> and Tl are present in the sample (so-called "organics"); those three analyses are not included in age interpretations or on Fig. 1, but are in Table S1.

## RBP

This sample yielded dozens of very small (<50  $\mu$ m) euhedral zircons. Eleven grains were successfully dated, two were inherited ca. 2500 Ma, and nine were ca. 65 Ma. In addition to being small, these grains were low-U and despite very low analytical blanks ( $\leq 0.33$  pg) had low radiogenic to blank Pb ratios, such that the data are less precise than other samples. Analyses with Pb\*/Pb<sub>c</sub> (radiogenic to common Pb)  $\leq 1$  were not considered in age interpretations. Despite the larger uncertainties, a weighted mean of the 4 most precise analyses gives a date of  $65.651\pm0.091$  Ma, whose mean is older than the overlying samples RBE, RBF and DEC13-08, -09, and -10.

## RBE

This sample yielded abundant euhedral prismatic zircons that varied in length from about 50-100  $\mu$ m, and these are interpreted to be volcanic in origin. Eleven grains were dated at PU and five were dated at MIT; all analyses were concordant. The zircons show ~250 ka of dispersion in <sup>206</sup>Pb/<sup>238</sup>U dates, which is well beyond the analytical precision. Both labs show similar dispersion and both the youngest and oldest grains from each lab overlap with each other.

## RBF

Less than 20 euhedral prismatic zircons 50-200  $\mu$ m in length were recovered from ~1 kg of material collected throughout the redbole horizon, and these are interpreted to be volcanically derived. Ten of the most pristine zircons were selected for geochronology and all were successfully analyzed. All are concordant and spread over ~200 ka, which is well beyond analytical uncertainty.

#### DEC13-08, -09, -10

Each sample from this volcaniclastic horizon yielded abundant euhedral prismatic zircon 50-400  $\mu$ m in length that are interpreted to be of volcanic origin. The morphology and size of grains from DEC13-08 and -09 are identical though zircon from DEC13-10 is on average smaller. One xenocrystic grain was analyzed in DEC13-08, at 680 Ma; the rest of the 44 grains analyzed are

ca. 65.5 Ma. Mean dates from each sample have a 150-300 kyr range, which is well beyond analytical uncertainties, and is seen by both MIT and PU laboratories.

### **S5 Results: Zircon trace element geochemistry**

Trace element data were obtained on the exact same volume of material dated by U-Pb in this study for all but three zircons presented (*50*, *52*). These are plotted as element ratio cross plots in Fig. S5A and against <sup>206</sup>Pb/<sup>238</sup>U date in Fig. S5B. Most important for this study is that when plotted as trace element cross plots, zircon from some samples plot in very different fields, consistent with their origination from different igneous sources. We use Zr/Hf in zircon as a proxy for igneous differentiation, similar to SiO<sub>2</sub> in whole rock geochemical data (*53*). Zr/Hf in our zircon suite varies between ~65 and 35, typical of igneous zircon. DEC13-30 zircons have Zr/Hf >50, consistent with their derivation from a less differentiated magma (*30*). Zircons from RBE, and a subset of zircons from RBF and RBP, also have high Zr/Hf, suggesting they originated from a relatively mafic volcanic ash, though the chemistry of their host magma is unknown. Zircons from the green ashbed samples DEC13-08, -09, and -10 have lower Zr/Hf and decreasing Zr/Hf correlates with increasing Nb/Gd and Yb/Gd. Increasing Nb/Gd with decreasing Zr/Hf, due the very incompatible nature of Nb, is also consistent with igneous differentiation controlling the geochemistry of zircons in the green ashbed and more generally for all zircons in this study.

When plotted against crystallization age (Fig. S5B), it can be observed that despite having prolonged crystallization in zircon from each sample that largely overlap, the geochemistry of the zircons distinguishes the stratigraphically older samples RBP and RBE from the green ashbed samples DEC13-08, -09, and -10. The older samples generally have higher Zr/Hf, lower Lu/Hf and Nb/Hf ratios than the younger ones. Several samples show weak correlations between age and geochemistry (also evident in Fig. 2), such as RBF, DEC13-08 and -09, that may result from fractional crystallization, magma mixing or assimilation in the magmatic system during zircon crystallization prior to eruption. This can affect zircon geochemistry either by mixing cores and rims of grains with different geochemical compositions or through equilibrium partitioning in an evolving liquid (*50, 52-59*). Outliers from geochemically equivalent clusters or from trends in time may be due to incorporation of zircon from a slightly older or geochemically distinct batch of magma – which are often called

antecrysts (60) – and the significance of such grains for our age interpretations are discussed in section S6.3.

### S6 U-Pb Age interpretations

The zircons from all of our samples except for DEC13-30 are interpreted as deriving from ashfall during periods of quiescence of basaltic volcanism. Though the source of the (presumably moderately silicic) volcanism is unknown, both the external morphology, internal textures, and magmatic Th/U of the zircons (Fig. S3) are consistent with this interpretation. Zircons show no evidence for pitting or rounding associated with alluvial or eolian transport, though inherited grains (i.e. >100 Ma) are markedly less euhedral than ca. 66 Ma grains. Previous studies on the geochemistry and petrography of redbole and greenbole horizons have concluded that a high-Si volcanic component can be found in a subset of boles (19, 38, 61, 62). Furthermore, our greenbole samples (DEC13-08, -09, -10) come from a bed that contains phenocrysts and small lithic fragments consistent with derivation from ashfall. That (euhedral) zircons from the redbole horizons yield consistent dates yet unique geochemistry compared to the greenbole also argues for a non-detrital source for those zircons. Ongoing work on the mineralogy and geochemistry of our samples will be used to assess the amount of volcanic material within the boles. Following a volcanic interpretation for our zircons, there are two steps in achieving accurate depositional ages for our samples, and thus for bounding flows of the Deccan Traps. The first is to obtain accurate crystallization ages for each zircon grain and the second is to interpret those data in terms of depositional ages. We discuss those in sequence below and present multiple different interpretations of our data in Table S3. We show that the major conclusions about the age and duration of the Deccan Traps are unaffected by our data interpretations.

## S6.1 Th/U disequilibrium correction in zircon

It is well documented that during crystallization of minerals useful for U-Th-Pb geochronology, fractionation of Th from U can cause the <sup>238</sup>U decay chain to depart from secular equilibrium (63-65). In the case of zircon, where U is preferentially incorporated in its lattice over Th, this can result in a depletion of <sup>206</sup>Pb and therefore ages that are too young by up to ~110 kyr (66-68). A nominal correction can be made for this by measuring or assuming a Th/U of the liquid from which the zircon crystallized and thus calculating a partition coefficient based on the Th/U of the

zircon (it is standard to calculate a model Th/U for each zircon based on the <sup>208</sup>Pb/<sup>206</sup>Pb and assuming concordance between the U-Pb and Th-Pb systems). Alternatively, one can assume a constant partition coefficient for Th/U between the liquid and zircon and calculate the resulting Th/U of the liquid. For recent discussions and approaches, see also refs. (67, 69-73). For most zircon and magma compositions, the difference between these approaches makes negligible difference in the resulting ages, but for young or high-precision ages this correction is an important contributor to the overall uncertainty budget (73-75). Because we have no information about the liquid from which our zircons crystallized, we explore both "constant magma Th/U" and "constant Th/U partition coefficient" approaches. We reduced our data in its entirety using both approaches and report the differing age interpretations in Table S3. The constant magma Th/U approach assumes and Th/U of the liquid to be 2.8±0.5, as this encompasses the majority of igneous liquids. For the constant partition coefficient approach, we use the average partition coefficient determined experimentally by Rubatto and Hermann (76) for andesitic to dacitic magmas of 0.33. In the end we prefer the constant partition coefficient approach because it seems more parsimonious to suggest that highly variable Th/U in zircons from single handsamples results from changing magma compositions and/or incorporation of antecrystic zircon rather than rapidly fluctuating partition coefficients. However, there is not sufficient experimental or empirical data for zircon/liquid partitioning over a wide range in temperatures and liquid compositions to permit a totally satisfactory interpretation (77). For our dataset, the choice between these models results in a negligible difference in dates for individual zircons or weighted mean dates for all samples except DEC13-30. Zircons from this sample have relatively high Th/U (1.9-2.6) compared to the other samples (which are  $\sim$ 1) such that the liquid Th/U resulting from a constant partition coefficient approach is ~5-7 as opposed to an assumed 2.8. This results in a difference of 10s of kyr for DEC13-30 zircons; the importance of this difference is best illustrated in the calculated duration of magnetic polarity Chron 29r, which differs by ~30 kyr. While the resulting durations overlap within uncertainty for both approaches (see Table S3), we regard this as the largest uncertainty in this study. All the ages presented in the main text, Figs. 1 and 2 and Table S1 use the constant partition coefficient approach.

### S6.2 Interpretation of zircon crystallization age spectra

Each of the samples in this study except RBP show a spread in ages that is larger than analytical uncertainty. This phenomenon is well-documented in igneous systems and can result from prolonged crystallization of zircon within a magmatic system or inheritance of much or slightly older zircons into a younger batch of magma or eruptive column (20, 21, 52, 57-59, 78-88). We interpret our data to be the result of a similar phenomena and provide very reproducible dates from zircon standard AUS-Z2 as a contrasting example (Fig. S2) to illustrate that the observed spread in dates in Deccan samples is not the result of analytical error.

Further constraints on the origin of the zircon age spectra come from the geochemistry of the dated zircons (Table S3, Fig. 1, S5). If the spread in zircon dates were a result of analytical error in the U-Pb geochronology of an otherwise homogeneous population of zircons, such a population also should have identical trace element characteristics, e.g., if they crystallized at the same time from a homogeneous liquid. While three of our seven samples exhibit this behavior, most do not. The observation of heterogeneous trace element characteristics in a suite of zircons that have a spread in dates is consistent with their derivation from liquids with different compositions or at different temperatures (50, 52, 54, 55, 58, 59). In several samples (RBF, DEC13-08, and -09), there are crude trends in zircon geochemistry with time (Fig. 1, S5) suggesting zircon was derived from an evolving system. Alternatively, such data could be interpreted as mixing of cores of older zircon with rims of younger zircon that each have distinct geochemistry. In either case, it implies that the observed spread in dates is real and not the result of analytical error in U-Pb systematics. It also implies that the youngest grain or grains from the volcanic derived zircons in each sample are the best estimate of the deposition of the material. We expand on this interpretation in the following section.

While pre-eruptive zircon growth is commonplace in moderate to high-silica volcanic systems, prolonged growth of zircon in a segregation vein represented by DEC13-30 is more peculiar. These veins are interpreted to represent in situ segregation of fluid-rich and/or silica-rich liquid from a rapidly crystallizing basalt flow. Such veins have been observed in a number of thick basaltic flows or shallow-level intrusions (69, 89-91) and also in lava lakes in Hawaii (92). Thermal arguments easily rule out that these lava flows crystallized in 200 kyr, which is the range in zircon dates observed in our sample of the Jawhar Fm. However, this sample also has abundant xenocrystic zircon ranging in age up to 1600 Ma. This observation argues against

a purely fractionation-derived origin for the segregation veins and thus the origin of these veins may involve localized shallow crustal melting. If the sampled igneous zircons that crystallized ca. 66.2 Ma contained even micron-scale cores of 1600 Ma material, the ages could be shifted older as observed. Such inclusions would not be large enough to change the trace elements of the zircon (i.e. a difference in 0.5% in U-Pb date is easily measureable, but is not in our geochemical data). Because there were too few zircons, we did not collect CL images to confirm this speculation but these grains were full of inclusions of dark brown material that would conceal any cores optically (Fig. S3). We thus interpret the spread in dates in DEC13-30 as mixing of older cores within euhedral igneous zircon, suggesting the youngest grains are the best estimates of deposition and rapid crystallization of the host basalt.

Finally, another possibility is that the spread in zircon dates results from Pb-loss. Diffusion of Pb in radiation-damaged zircon is likely the cause of Pb-loss in natural zircon (93-95). The ability to remediate Pb-loss has recently been revolutionized by the chemical abrasion technique (39), which has become common practice in ID-TIMS laboratories and modified into a single-stage leaching technique (40, 96, 97). Though some age scatter in chemically abraded zircons has still been attributed to Pb-loss (21, 97), especially in highly metamict samples (98, 99), it is generally assumed that in young samples subjected to chemical abrasion, Pb-loss is negligible. For the data presented here, we can make several arguments to support this claim. The first is that in general stratigraphic order is maintained in our geochronologic dataset (with the exception of the composite ashbed sample, where two of three give slightly older dates; see discussion below), which would be quite fortuitous given the anticipated non-systematic nature of Pb-loss in grains with widely varying U-content and size. Second is in samples that exhibit trends in geochemistry with time, Pb-loss cannot be the cause of spread in dates and must result from some zircon growth/inheritance process. We thus discount Pb-loss as an important mechanism of generating the spread in dates in our dataset.

## S6.3 Best estimates for deposition of dated horizons

The previous section outlined the reasoning for interpreting the observed spread in zircon dates as a real geologic phenomena involving prolonged growth, mixing of cores and rims of zircons, and/or inheritance of slightly older grains into a volcaniclastic deposit. Recent approaches to arrive at depositional ages for volcanic deposits exhibiting this behavior, assuming negligible Pbloss, are to use the youngest zircon date as the best estimate or a small subset of the youngest grains for a weighted mean (21, 100, 101). The former method is the more conservative of the two in that it will always overlap with the latter interpretation but have larger uncertainties. This method also minimizes any effects of pre-eruptive zircon on the accuracy of the resulting depositional age. The benefit of the latter approach is that it provides a higher-precision result and if anticipated analytical scatter has resulted from a homogeneous true age spectra, this method will maximize the accuracy of the result. Here, we build upon the previous discussion that integrated U-Pb data and geochemical data from the same zircons to arrive at age interpretations that capitalizes on both approaches.

Accurate weighted means require that the data in fact represent a single population and that all scatter is analytically derived. Simply because a set of geochronological datapoints can be used to calculate a weighted mean with a reasonable MSWD (as defined by Wendt and Carl (102)) does not mean they meet that criteria given the possibility of subtle geologic dispersion that is masked by large analytical uncertainties (75, 103). Bringing geochemical data to the problem can add confidence that a population of zircon is in fact uniform in that if such a population crystallizes from a melt at exactly the same time, it should have the same geochemical signature (50, 52, 58, 59). From our dataset, a subset of the youngest zircons from each sample are the same in both age and geochemistry. In several cases (DEC13-09, RBF, Fig. 1, S5), increasing N would not jeopardize the MSWD of the weighted mean <sup>206</sup>Pb/<sup>238</sup>U date, but would violate geochemical equivalence, and so N is limited. We thus have an independent means of determining which subset of U-Pb analyses are most likely to represent a population amenable to weighted means without biasing ages too old due to antecrystic zircons. Weightedmeans for these samples are calculated and the analyses chosen are outlined by gray boxes in Fig. 2 and in bold in Table S1. A full assessment of using the alternative youngest single grain approach is given in Table S3 and the choice of method does not effect the main conclusions of this paper.

As a final step in building an age model for the middle Ambenali Fm through the lower Mahabaleshwar Fm, we use a Markov Chain Monte Carlo simulation that imposes stratigraphic order onto the existing geochronological dataset (*104*). This Bayesian approach is able to refine uncertainty estimates for depositional ages given that it is impossible for lower beds to be younger than higher ones and by using the geochronological data as estimated priors. Similar models have been used to calibrate the radiocarbon timescale (105), and have been used recently to integrate radioisotopic dates with timescales derived from orbital tuning of cyclic stratigraphic sections (100). Our model produces some number (typically a million) of data points for each sample with a Gaussian distribution and standard deviation equal to that of the uncertainties from the weighted means described above. It then randomly selects one datum from each sample. If the sequence of data points follows stratigraphic order, the path is kept, if it does not (i.e. one or more stratigraphically higher datapoints is older than those beneath), it throws out the result. This is repeated a million times and the resulting paths are used to define new means and standard deviations of each dated bed. This method is especially useful if beds have largely overlapping uncertainties as with our data from the upper Deccan, but has no effect if the beds do not overlap within uncertainties such as with DEC13-30. It is worth noting that this model does not make any assumptions about sedimentation rate, only superposition.

This model can also be used to explore the probability that stratigraphic order is violated if the model produces no, or an exceedingly small number of, acceptable paths. In our dataset, this is the case for DEC13-08 and -09, for which less than five out of a million acceptable paths are produced if they are included in the model. We hence conclude that of the three samples collected from this thick ashbed, the stratigraphically lowest, DEC13-10, sampled the youngest part of the magmatic system or at least the population of zircons that crystallized the latest on average in the system. Whole zircon grains record an average time of growth that may span over tens or hundreds of thousands of years depending of the zircon saturation temperature (106, 107), cooling rate of the magmatic system or incorporation of antecrystic material (58, 72). Given the potentially complex amalgamation history of the sample tuff, we interpret the older dates found in DEC13-08 and -09 to represent some combination of all these processes manifest within the age spectra. Geochemistry of the zircons from each sample are also different: DEC13-10 has zircon that are relatively homogeneous except for two analyses with low Lu/Hf (Fig. 2), DEC13-09 exhibits a lot of scatter but a generally increasing trend in Lu/Hf with time, while DEC13-08 shows the opposite trend. This provides evidence that each sample is in fact recording a different part of an eruptive system and the differences in age is not just a sampling artifact. We therefore use the age of DE13-10 as our best estimate for the deposition of that tuff and link it to the Monte Carlo age model derived using the other samples as well. The resulting dates from the Monte Carlo model are plotted in Fig. 1 and also in Table S3, which also provides a

comprehensive view of alternative age interpretations and their effect on the duration of C29r and the minimum duration of the Deccan Traps.

On a final note, it is worth discussing alternative interpretations of the zircon date and geochemical data. For example, we could have increased the number of zircons from RBE and DEC13-10 included in their weighted means while still retaining geochemical and date equivalence (the former being a mostly qualitative assessment, while the latter is evaluated by the acceptable range of MSWDs (*102*)). If we maximize the number of grains included in the acceptable weighted means for RBE (N=12) and DEC13-10 (N=9), the resulting ages become older and more precise. These ages become exceedingly unlikely to pass the superposition test from the Monte Carlo analysis with <10 of one million paths being acceptable, and as such we err on the side of including fewer zircons in our weighted means given the potential for preeruptive zircon growth that is unresolvable even with our very precise data. However, if that interpretation was chosen, it would shorten our calculated duration of C29r by ~35 kyr but still overlap within uncertainty with our preferred interpretation. In no reasonable interpretation do DEC13-08 and -09 pass the superposition test, so we regard their omission as unavoidable.

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See main text for references 1-34

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**Fig. S1: Sample locations and outcrop photos.** A. photos of sample locations as described in text. DEC13-08, -09, 10; shown is ~40 cm thick green horizon located at flow boundary. RBF, RBE, RBP: redbole horizons also located at flow boundaries, note that RBF has red paleosol falling into cracks in underlying basalt flow. In each, unaltered basalt flow bottom is observed at top of photo. DEC13-30: BS sampling freshly hammered friable segregation vein within massive basalt flow. Note vein is ~10 cm thick subhorizontal band in photo but other veins in the same outcrop are subvertical. B. shaded relief map of Western Ghats with sample locations. Color bar is altitude in meters, white squares are major cities. C. View from Mahabaleshwar of typical outcrop pattern of the Deccan Traps on weathered hillsides. Easily accessible exposures are located along roadcuts running up the cliff faces. D. Composite stratigraphic column showing sample locations, magnetic polarity timescale. See main text for references.



b Corrected for initial Pa/U disequilibrium using initial fraction activity ratio [231Pa]/[235U] = 1.1 c Isotopic dates calculated using the decay constants  $\lambda 238 = 1.55125E-10$  and  $\lambda 235 = 9.8485E-10$  (Jaffey et al. 1971).

d % discordance = 100 - (100 \* (206Pb/238U date) / (207Pb/206Pb date))

e Th contents calculated from radiogenic 208Pb and the 207Pb/206Pb date of the sample, assuming concordance between U-Th and Pb systems

f Total mass of radiogenic Pb.

g Total mass of common Pb

h Ratio of radiogenic Pb (including 208Pb) to common Pb

i Measured ratio corrected for fractionation and spike contribution only. i Measured ratios corrected for fractionation, tracer and blank

Fig. S2. U-Pb geochronological and trace element geochemical data for natural zircon standard

AUS-Z2. A. Concordia diagram for U-Pb data from MIT and PU. For comparison, <sup>206</sup>Pb/<sup>238</sup>U age reported by (49) was  $38.896 \pm 0.004/0.012/0.043$  Ma (MSWD = 1.0, n = 12/12). All uncertainties stated at the 2s level. B. Rank-order plot of AUS-Z2<sup>206</sup>Pb/<sup>238</sup>U dates (same analyses as plotted in A) with axis in millions of years. C. Spider diagram of chondrite-normalized REE (108) concentrations (Nd to Lu) in AUS-Z2, from solution ICPMS (this study; using methods of (50)) and LA-ICPMS (49). Note that the solution ICPMS data are not the exact same zircons shown in A and B. The single black curve represents the mean of 10 LA-ICPMS analyses from (49).



**Fig. S3.** Photographs of zircons from this study. A. Optical images of zircons under transmitted or reflected light from PU and MIT, respectively. Note sharp crystal terminations are interpreted to reflect igneous growth with negligible "reworking" or sedimentary transport. Hence, these are interpreted as deriving from igneous segregations (for DEC13-30) or ashfalls (all others). U-Pb data given in Table S1 and geochemical data for the same zircons is given in Table S2. B. Cathodoluminescence images of zircons from DEC13-08, illustrating oscillatory zoning indicative of igneous growth. A lack of obvious resorbed cores indicates that inheritance of much older grains is probably not the cause of spread in measured dates.



Fig. S4: Concordia plots for U-Pb ID-TIMS geochronological data. Data provided in Table S1. Note the scales for top four plots are the same, as are scales for bottom two plots. Grayed out ellipses are those not included in age interpretations due to large uncertainties or reverse discordance. Gray band around concordia curve includes uncertainties in the U decay constants. Age interpretations shown are based on the data illustrated alone, not the ages resulting from the Monte Carlo simulations (see text). See Table S3 for alternative age interpretations. All uncertainties are at the  $2\sigma$  level.



**Fig. 5: Trace element data for zircons dated in this study.** A. Element ratio cross-plots using Zr/Hf as an igneous differentiation index. Important here is that different samples plot in different spaces in these plots, illustrating that zircons from different hand samples derived from different sources. Also important is that the dataset as a whole follows trends consistent with igneous differentiation (e.g. Nb/Gd vs. Zr/Hf). B. geochemical data plotted versus ID-TIMS <sup>206</sup>Pb/<sup>238</sup>U date. Again, note that despite largely overlapping age distributions between samples, zircons have distinct geochemical differences and thus reflect different volcanic sources. Scale only goes up to 65.8 Ma, so DEC13-30 is not plotted.

- **Table S1**. U-Pb isotopic data obtained by CA-ID-TIMS at MIT and PU. See text for details.Included at Science Online as .xlsx file.
- **Table S2**. Zircon trace element obtained by solution ICP-SF-MS. Trace element data obtainedfor same volume of dissolved zircon analyzed for U-Pb isotopes. See text for details.Included at Science Online as .xlsx file.
- Table S3. Age interpretation table. Multiple interpretations of U-Pb data given as discussed in text.

   Included at Science Online on vlay file

Included at *Science Online* as .xlsx file.

### Table S1: U-Pb isotopic data

	Date	s (Ma)						Composition			Isotopi	c Ratios							
Sample name	206Pb/	±2σ	207Pb/	±2σ	207Pb/	±2σ		Th/	Pbc	Pb*/	206Pb/	208Pb/	206Pb/		207Pb/		207Pb/		Corr.
	238U	absolute	235U	absolute	206Pb	absolute	% disc	U	(pg)	Pbc	204Pb	206Pb	238U	±2σ %	235U	±2σ %	206Pb	±2σ %	coef.
m	а		b		a,b		d	e	g	h	i	k	a,k		b,k		a,b,k		
DEC13-08																			
z4 BS	65.549	0.061	65.6	1.0	66.0	36.6	4.9	1.09	0.57	8.51	464	0.35	0.0102201	0.093	0.0666984	1.608	0.0473538	1.537	0.772
z3 BS	65.574	0.028	65.69	0.30	69.9	10.8	9.9	1.10	0.63	29.10	1538	0.35	0.0102241	0.044	0.0668331	0.477	0.0474306	0.454	0.557
710 KMS	65 602	0 031	65 49	0 30	61 5	10.8	-1.9	1.05	0 34	31 05	1657	0 34	0 0102285	0 048	0 0666274	0 472	0 0472646	0 450	0 476
210_KHS	65 602	0.031	65.45	0.30	69.0	25.0	9 7	0.95	0.95	11 59	650	0.34	0.0102205	0.076	0.0669259	1 120	0.0474124	1 097	0.704
208_KH3	03.002	0.049	03.03	0.72	09.0	23.5	0.7	1.00	0.83	10.10	535	0.27	0.0102283	0.070	0.0008338	1.139	0.0474124	1.067	0.704
200_KM3	05.054	0.067	05.71	0.91	00.5	52.5	8.0	1.22	0.00	10.10	221	0.39	0.0102554	0.105	0.0008342	1.420	0.0474026	1.559	0.089
Z1_BS	65.643	0.025	65.688	0.089	67.3	3.1	6.5	0.94	0.42	98.80	5367	0.30	0.0102348	0.038	0.0668308	0.140	0.04/3/95	0.127	0.356
z11_KMS	65.667	0.029	65.83	0.12	71.6	4.2	11.8	0.96	0.41	91.44	4939	0.31	0.0102386	0.044	0.0669765	0.193	0.0474651	0.175	0.447
z05_KMS	65.671	0.054	65.62	0.45	63.7	16.4	1.3	1.09	0.60	20.93	1111	0.35	0.0102393	0.083	0.0667581	0.709	0.0473071	0.688	0.298
z02_KMS	65.688	0.037	65.46	0.33	57.3	12.1	-9.2	1.01	0.53	26.96	1454	0.32	0.0102419	0.056	0.0665967	0.521	0.0471808	0.505	0.311
z07_KMS	65.694	0.029	65.761	0.080	68.2	2.7	7.5	1.12	0.33	141.30	7345	0.36	0.0102429	0.045	0.0669075	0.125	0.0473963	0.108	0.446
z04_KMS	65.697	0.036	65.84	0.19	70.9	6.5	10.9	1.07	0.50	50.59	2674	0.34	0.0102433	0.055	0.0669864	0.300	0.0474505	0.271	0.560
z12 KMS	682.81	0.20	1180.56	0.35	2260.28	0.60	69.8	0.08	0.42	904.66	56608	0.026	0.1117344	0.031	2.198492	0.050	0.1427684	0.012	0.729
DEC13-09																			
DEC13-03	~ ~ ~ ~				<b>63</b> 0														
Z12_KMS	65.568	0.031	65.50	0.37	63.0	13.3	0.5	1.02	0.48	23.62	1273	0.33	0.0102231	0.048	0.0666339	0.586	0.04/2939	0.559	0.587
z2_MIT	65.589	0.041	65.73	0.18	70.8	6.4	11.0	0.92	0.42	77.24	4159	0.29	0.0102264	0.062	0.0668746	0.282	0.0474495	0.267	0.306
z5_MIT	65.594	0.049	65.65	0.45	67.6	16.6	6.9	1.20	0.49	30.31	1540	0.38	0.0102272	0.075	0.0667882	0.714	0.0473844	0.695	0.283
z2_BS	65.608	0.034	65.67	0.20	67.8	7.4	7.2	1.11	0.43	80.05	4187	0.35	0.0102293	0.052	0.0668095	0.311	0.0473898	0.311	0.060
z02_KMS	65.623	0.027	65.64	0.17	66.2	6.1	5.0	0.97	0.27	56.70	3064	0.31	0.0102318	0.042	0.0667800	0.273	0.0473577	0.253	0.501
z10 KMS	65.623	0.028	65.52	0.11	61.9	3.9	-1.3	1.05	0.39	90.75	4802	0.34	0.0102318	0.043	0.0666592	0.174	0.0472717	0.162	0.328
z1 MIT	65.634	0.038	65.77	0.36	70.8	13.2	10.8	1.15	0.51	36.93	1894	0.37	0.0102335	0.059	0.0669184	0.565	0.0474478	0.552	0.252
73 MTT	65 635	0.032	65 59	0.22	64 0	8.0	1 9	1 01	0.35	60 70	3203	0.32	0 0102336	0 049	0.0667303	0 349	0 0473137	0 335	0 329
23_NIT	65 64	0.052	66.7	1 2	95 9	41 7	25.0	0.02	0.35	11 02	600	0.32	0.0102330	0.160	0.0672474	1 706	0.0477484	1 757	0.323
24_PLT	65.645	0.10	65 500	1.4	62.0	41.7	23.5	0.52	0.30	216 08	11674	0.25	0.0102343	0.100	0.06673474	1.750	0.0472002	1./3/	0.200
ZU7_KMS	65.645	0.024	05.590	0.057	63.8	2.0	1.5	0.97	0.34	216.98	11674	0.31	0.0102352	0.037	0.0667342	0.090	0.0473093	0.078	0.356
ZU4_KMS	65.646	0.033	65.55	0.22	62.0	7.8	-1.3	1.13	0.26	43.28	2262	0.36	0.0102354	0.051	0.0666840	0.348	0.04/2/29	0.326	0.465
z08_KMS	65.650	0.034	65.56	0.30	62.4	10.8	-0.6	1.02	0.41	29.37	1579	0.33	0.0102360	0.051	0.0667004	0.474	0.0472818	0.450	0.485
z5_BS	65.657	0.032	65.82	0.40	71.8	14.2	12.1	1.10	0.78	22.21	1178	0.35	0.0102371	0.049	0.0669722	0.625	0.0474690	0.596	0.608
z6_MIT	65.66	0.16	66.2	1.7	86.5	60.1	26.6	0.99	0.60	7.74	426	0.31	0.0102373	0.240	0.0673902	2.599	0.0477644	2.535	0.308
z06_KMS	65.659	0.045	65.71	0.41	67.5	14.8	6.7	1.04	0.34	22.16	1189	0.33	0.0102375	0.069	0.0668528	0.652	0.0473829	0.620	0.497
z1 BS	65,669	0.024	65.766	0.077	69.3	2.4	9.0	0.94	0.50	126.23	6845	0.30	0.0102390	0.037	0.0669135	0.121	0.0474189	0.097	0.652
709 KMS	65 673	0.025	65 760	0.067	68.9	23	8 5	1.06	0 33	164 82	8689	0 34	0 0102396	0.038	0 0669074	0 106	0 0474116	0 090	0 431
z11 KMS	65 721	0.034	65 81	0.29	69 1	10.5	8 7	1 11	1 31	31 04	1633	0.36	0 0102472	0.052	0.0669622	0 458	0 0474155	0 441	0 375
211_0.05	05.721	0.054	05.01	0.25	05.1	10.5	0.7	1.11	1.51	51.04	1055	0.50	0.0102472	0.052	0.0005022	0.450	0.04/4155	0.441	0.575
DEC12 10																			
DEC 13-10																			
z11_KMS	65.508	0.081	65.72	0.99	73.4	34.8	14.1	0.93	0.93	8.96	505	0.30	0.0102137	0.125	0.0668623	1.549	0.0474999	1.466	0.696
z9_MIT	65.581	0.049	65.49	0.41	62.3	14.9	-0.7	0.98	0.32	32.87	1756	0.31	0.0102252	0.075	0.0666260	0.645	0.0472787	0.623	0.335
z7_MIT	65.586	0.084	65.00	0.66	43.4	24.6	-41.9	1.16	0.27	21.05	1086	0.37	0.0102259	0.129	0.0661054	1.052	0.0469060	1.028	0.242
z2_MIT	65.594	0.035	65.65	0.30	67.6	11.0	6.9	1.01	0.27	50.68	2677	0.32	0.0102272	0.054	0.0667882	0.466	0.0473847	0.460	0.145
z4_BS	65.597	0.059	65.27	0.59	53.5	21.1	-16.5	1.11	0.45	16.04	854	0.35	0.0102277	0.090	0.0663972	0.937	0.0471049	0.886	0.596
73 BS	65.630	0.038	65.80	0.57	72.1	20.3	12.5	1.01	0.51	14.97	815	0.32	0.0102329	0.058	0.0669534	0.892	0.0474752	0.852	0.708
75 BS	65 636	0 099	66 13	0.95	84 0	33 4	24 4	1 00	0 34	11 12	612	0 32	0 0102337	0 152	0.0672930	1 480	0 0477123	1 406	0 524
72 BS	65 641	0 049	65 67	0 60	66.8	21 1	5.8	1 13	0 91	14 90	789	0.36	0.0102346	0.075	0.0668145	0 946	0 0473690	0 885	0.826
22_05	65.644	0.049	65.07	0.00	71 1	12.0	11 2	0.00	0.31	42.20	2245	0.50	0.0102350	0.073	0.0660388	0.540	0.0474550	0.005	0.020
20_011	05.044	0.048	05.79	0.35	/1.1	12.9	11.5	0.99	0.27	42.29	2245	0.32	0.0102330	0.075	0.0009588	0.330	0.0474550	0.340	0.190
21_65	05.070	0.020	05.75	0.11	00.4	4.2	7.0	0.96	0.51	/5.//	4098	0.51	0.0102400	0.051	0.0008954	0.178	0.0475999	0.172	0.182
ZIO_MII	65.71	0.10	66.4	1.1	90.5	38.4	29.7	0.91	0.30	12.04	665	0.29	0.0102456	0.158	0.06/558/	1.6/1	0.04/8450	1.619	0.370
z6_MIT	65.712	0.205	64.7	2.2	26.2	83.5	-126.4	0.91	0.24	5.90	335	0.29	0.0102457	0.314	0.0657593	3.577	0.0465705	3.481	0.348
z10_KMS	65.717	0.069	65.58	0.80	60.4	28.7	-3.8	1.00	0.47	11.33	623	0.32	0.0102465	0.106	0.0667136	1.255	0.0472427	1.203	0.530
z07_KMS	65.76	0.11	65.6	1.7	60.2	62.1	-4.3	1.50	0.58	5.85	297	0.48	0.0102528	0.174	0.0667485	2.722	0.0472382	2.604	0.696
z09_KMS	65.824	0.083	66.6	1.3	94.0	45.4	32.1	0.94	0.69	6.80	386	0.30	0.0102633	0.127	0.0677748	2.008	0.0479155	1.918	0.726
RBF																			
72 BS	65 54	0 18	63 0	3 1	-33 1	116 5	316 6	0.92	0.36	2 82	172	0.29	0 0102183	0 277	0 0639925	5 024	0 0454405	4 804	0 803
202 KMS	65 557	0.10	66 1	1 1	85.6	29 5	25.0	0.92	0.30	7 99	452	0.25	0.0102105	0.124	0.0672589	1 705	0.0477454	1 621	0.005
203_KH3	05.557	0.081	00.1	1.1	40.0	50.5	23.5	0.87	0.20	7.88	432	0.28	0.0102214	0.124	0.0072389	1.703	0.0477434	2.1021	0.701
Z3_B5	65.594	0.093	65.1	1.4	48.9	52.1	-26.9	0.99	0.33	6.11	345	0.32	0.0102272	0.142	0.0662656	2.292	0.04/0138	2.183	0.778
ZU6_KMS	65.595	0.041	65.62	0.52	66.6	18.6	5.5	0.95	0.39	16.80	926	0.30	0.0102274	0.063	0.066/612	0.812	0.04/3646	0.779	0.541
z4_BS	65.603	0.047	65.36	0.62	56.5	22.4	-10.5	0.97	0.37	14.26	784	0.31	0.0102287	0.072	0.0664886	0.979	0.0471652	0.939	0.585
z5_BS	65.639	0.034	65.88	0.26	74.7	9.4	15.4	1.04	0.44	35.48	1893	0.33	0.0102342	0.052	0.0670338	0.415	0.0475262	0.396	0.410
z01_KMS	65.667	0.037	65.75	0.24	68.8	8.4	8.3	0.92	0.51	37.42	2052	0.30	0.0102386	0.057	0.0668959	0.374	0.0474080	0.351	0.449
z02_KMS	65.674	0.067	65.86	0.58	72.7	20.4	13.1	0.98	0.32	17.84	976	0.31	0.0102397	0.102	0.0670142	0.908	0.0474867	0.859	0.519
z04 KMS	65.714	0.041	66.20	0.51	83.7	18.0	24.1	1.43	0.22	19.58	967	0.46	0.0102460	0.062	0.0673657	0.791	0.0477068	0.759	0.540
71 BS	65 751	0 044	65 92	0.64	72 0	22.8	12 1	1 03	2 23	13 10	718	0 33	0 0102519	0.067	0.0670726	1 007	0 0474718	0 960	0 713
21_00	05.751	0.044	03.52	0.04	72.0	22.0	16.1	1.05	2.23	13.15	/10	0.00	0.0102019	0.007	0.00/0/20	1.007	0.07/4/10	0.000	0.715
DBE																			
RBE																			
z6_MIT	65.39	0.33	64.6	2.8	33.6	105.5	-79.5	0.92	0.46	5.09	291	0.29	0.0101959	0.507	0.0656420	4.514	0.0467146	4.403	0.274
z7_MIT	65.39	0.22	62.8	2.6	-36.5	104.0	294.0	0.94	0.54	5.04	287	0.30	0.0101959	0.345	0.0637613	4.332	0.0453757	4.283	0.183
z3_MIT	65.49	0.25	64.3	2.8	20.4	103.3	-181.7	0.92	0.37	5.00	287	0.29	0.0102112	0.376	0.0653804	4.416	0.0464586	4.300	0.347
z03_KMS	65.49	0.16	67.8	2.3	150.4	79.5	57.2	1.01	1.05	3.70	215	0.32	0.0102113	0.239	0.0690638	3.569	0.0490752	3.393	0.757
z10 KMS	65.502	0.057	65.30	0.92	57.9	33.2	-7.7	1.06	0.42	9.47	516	0.34	0.0102128	0.088	0.0664242	1,453	0.0471930	1.390	0.737
z5 85	65.511	0.065	65.4	1.0	60.2	36.3	-3.8	0.94	0.69	8.31	469	0.30	0.0102143	0.100	0.0664982	1.592	0.0472386	1.523	0.702
23_05	65.511	0.003	65.4	2.0	61 1	00.5	-3.0	1 03	0.05	2 00	103	0.30	0.0102143	0.100	0.0665799	4 352	0.0472550	1.323	0.702
22_03	02.3/	0.15	03.4	2.0	01.1	33.2	-2.5	1.02	0.50	5.00	101	0.35	0.0102229	0.224	0.0003/68	4.332	0.04/2009	4.104	0.045

z1_BS	65.58	0.19	66.2	3.6	88.5	128.6	28.3	0.88	0.43	2.25	142	0.28	0.0102248	0.289	0.0673641	5.669	0.0478046	5.426	0.849
ZZ_MIT	63.39	0.16	65.0	1.6	42.0	39.0	-45.0	0.96	0.33	0.20	430	0.31	0.0102238	0.242	0.0000892	2.364	0.0400930	2.302	0.3//
ZUL_KMS	65.600	0.050	05.05	0.55	67.4	19.4	0.0	0.95	0.18	16.72	922	0.30	0.0102281	0.077	0.0667886	0.864	0.0473808	0.815	0.660
Z4_MLI	05.01	0.15	65.2	1.0	51.5	57.4	-20.8	0.92	0.36	8.18	457	0.29	0.0102303	0.223	0.0663588	2.461	0.0470655	2.405	0.294
ZU8_KMS	65.622	0.078	66.2	1.3	88.3	45.2	28.0	0.93	0.37	0.95	395	0.30	0.0102317	0.120	0.0674029	1.997	0.0477998	1.905	0.787
ZU4_KMS	65.648	0.070	65.49	0.75	59.8	27.0	-4.8	0.88	0.19	12.40	699	0.28	0.0102357	0.107	0.0666245	1.189	0.0472294	1.134	0.547
ZUS_KMS	65.66	0.24	67.7	4.4	142.2	150.5	54.8	0.91	1.08	1.90	122	0.29	0.0102370	0.369	0.0689961	6.721	0.0489040	6.413	0.844
Z3_BS	65.670	0.063	65.68	0.67	66.1	23.0	4.8	0.95	0.23	13.83	765	0.30	0.0102391	0.096	0.0668254	1.048	0.04/3561	0.965	0.875
ZU2_KMS	65.724	0.056	65.61	0.70	61.4	25.3	-2.4	0.90	0.20	13.83	//4	0.29	0.0102476	0.086	0.0667470	1.109	0.0472609	1.062	0.574
RBP																			
z1_BS	65.51	0.20	62.3	3.4	-60.5	131.5	213.5	1.08	0.32	2.74	162	0.34	0.0102144	0.305	0.0632501	5.631	0.0449305	5.391	0.797
z05_KMS	65.55	0.65	67.5	11.7	137.1	402.2	53.2	1.03	0.43	0.75	58	0.33	0.0102196	0.997	0.0687291	17.956	0.0487979	17.119	0.849
z4_BS	65.65	0.13	64.7	2.3	27.8	85.4	-114.0	1.19	0.31	3.83	214	0.38	0.0102365	0.197	0.0657456	3.718	0.0466026	3.560	0.814
z5_BS	65.67	0.38	63.5	7.3	-18.2	277.2	528.3	1.17	0.33	1.26	83	0.37	0.0102395	0.581	0.0645214	11.941	0.0457212	11.461	0.834
z2_BS	65.79	0.18	62.8	3.1	-49.0	118.3	242.4	1.02	0.27	3.38	198	0.33	0.0102575	0.279	0.0638180	5.079	0.0451436	4.860	0.795
z06_KMS	65.86	0.59	67.2	10.8	116.3	374.3	44.8	1.26	0.36	0.84	61	0.40	0.0102685	0.908	0.0684517	16.629	0.0483693	15.870	0.845
z03_KMS	65.93	0.47	71.6	8.1	263.8	257.3	75.3	1.00	0.30	1.07	75	0.32	0.0102805	0.719	0.0730180	11.783	0.0515359	11.210	0.810
z08_KMS	66.03	0.49	66.3	9.0	76.9	317.1	17.2	1.02	0.43	1.02	73	0.32	0.0102951	0.745	0.0674950	13.965	0.0475700	13.346	0.841
z11_BS	66.10	0.61	67.2	11.5	105.2	398.3	38.9	0.95	0.36	0.73	58	0.30	0.0103066	0.931	0.0683833	17.646	0.0481425	16.854	0.859
z3_BS	1830.44	0.54	2093.88	0.52	2363.87	0.82	22.6	0.14	0.32	362	22189	0.040	0.3283635	0.034	6.8629084	0.059	0.1516517	0.035	0.572
z01_KMS	2512.95	1.57	2595.67	0.89	2660.88	0.82	5.6	0.30	0.40	105	6068	0.083	0.4767175	0.075	11.8886797	0.095	0.1809533	0.037	0.856
DEC13-30																			
z7 BS	66.249	0.0489	65.34	0.38	32.2	12.3	-89.2	2.22	0.33	48.46	2028	0.71	0.0103298	0.074	0.0664657	0.594	0.0466874	0.514	1.056
z4 BS	66.286	0.0347	66.28	0.19	66.1	6.6	3.8	2.33	0.34	72.51	2968	0.74	0.0103356	0.053	0.0674534	0.292	0.0473545	0.274	0.397
z10 KMS	66.289	0.0770	65.4	1.2	32.0	42.9	-90.4	2.62	0.43	10.31	417	0.84	0.0103361	0.117	0.0665010	1.873	0.0466839	1.789	0.732
z5 BS	66.291	0.0420	66.21	0.41	63.2	13.8	-0.3	1.69	0.29	27.06	1260	0.54	0.0103364	0.064	0.0673787	0.636	0.0472983	0.580	0.871
z9 BS	66.300	0.0406	65.64	0.21	41.5	7.1	-49.5	2.22	1.10	56.92	2380	0.71	0.0103379	0.062	0.0667778	0.331	0.0468700	0.295	0.619
z6 BS	66.325	0.0238	66.21	0.26	62.2	9.1	-2.0	2.47	0.33	45.67	1830	0.79	0.0103417	0.036	0.0673843	0.399	0.0472780	0.381	0.512
z10 BS	66.343	0.0306	66.37	0.43	67.4	15.4	5.5	1.90	0.53	26.51	1186	0.61	0.0103445	0.046	0.0675484	0.676	0.0473804	0.645	0.678
z08 KMS	66.390	0.0243	66.34	0.20	64.4	7.1	1.3	2.27	0.31	60.43	2501	0.73	0.0103520	0.037	0.0675129	0.309	0.0473215	0.295	0.381
z07 KMS	66.465	0.0372	66.38	0.22	63.4	7.7	-0.4	2.35	0.47	58.69	2394	0.75	0.0103637	0.056	0.0675602	0.344	0.0473010	0.321	0.456
z09 KMS	66.48	0.12	66.42	0.38	64.3	11.4	1.0	2.13	0.66	52.64	2236	0.68	0.0103658	0.184	0.0676013	0.597	0.0473202	0.476	0.741
z01 KMS	159.92	0.10	160.11	0.77	162.9	11.6	2.5	1.17	0.29	27.45	1427	0.37	0.0251173	0.064	0.1707917	0.520	0.0493388	0.497	0.405
z11_KMS	230.90	0.70	232.8	3.5	251.9	36.4	8.6	0.94	0.43	8.40	473	0.30	0.0364672	0.308	0.2576671	1.699	0.0512684	1.580	0.463
z3 BS	242.48	0.30	237.8	3.7	191.4	38.9	-26.2	0.98	0.30	7.66	430	0.31	0.0383311	0.126	0.2638500	1.736	0.0499459	1.674	0.523
z12_KMS	795.10	0.42	815.0	2.0	869.6	6.8	8.6	0.01	0.59	22.51	1555	0.003	0.1312680	0.056	1.2313501	0.355	0.0680639	0.326	0.551
z05_KMS	1496.1	1.4	1509.2	2.6	1527.7	5.6	2.1	0.79	0.38	27.58	1534	0.23	0.2612288	0.106	3.4210552	0.332	0.0950240	0.296	0.467

m sample name followed by specific zircon name. \_KMS, and \_BS are analyses from PU; \_MIT are analyses from MIT a Corrected for initial Th/U disequilibrium using radiogenic 208Pb and using the average (Th/U) partition coefficient between silicate melt of zircon from Rubatto and Hermann (2007) of 0.33. b Corrected for initial Pa/U disequilibrium using initial fraction activity ratio (231Pa)(235U] = 1.10000. c Isotopic dates calculated using the decay constants A238 = 1.5512EF-10 and A235 = 9.8486E-10 (Jaffey et al. 1971). d % discordance = 100 - (100 \* (208Pb/238U date) / (207Pb/208Pb date)) e Th contents calculated from radiogenic 208Pb and the 207Pb/230 Ed ef of the sample, assuming concordance between U-Th and Pb systems. f Total mass of radiogenic Pb, g Total mass of common Pb, assuming all common Pb is from blank. Blank compositions are reported in the methods section. h Ratio of radiogenic Pb (including 208Pb) to common Pb. j Measured ratios corrected for fractionation, and spike contribution only. k Measured ratios corrected for fractionation, tracer and blank.

Table S2: zircon	trace	element	data
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Table S2:	zircon tra	ce element	t data																												
а	date b	±2sig	Th/U b	Y-89 c	±2sig d	Zr-91	±2sig	Nb-93	±2sig	La-139	±2sig	Ce-140	±2sig	Pr-141	±2sig	Nd-146	±2sig	Sm-152	±2sig	Eu-151	±2sig	Gd-158	±2sig	Tb-159	±2sig	Dy-163	±2sig	Ho-165	±2sig	Er-166	±2sig
DEC13-08 z4_BS z3_BS z10_KMS z08_KMS z06_KMS z1_BS z11_KMS z05_KMS z02_KMS z07_KMS z04_KMS	65.549 65.674 65.602 65.634 65.643 65.667 65.671 65.688 65.688 65.694 65.697	0.061 0.028 0.031 0.049 0.067 0.025 0.025 0.029 0.054 0.037 0.029 0.036	1.09 1.10 1.05 0.85 1.22 0.94 0.96 1.09 1.01 1.12 1.07	5481 5360 4512 3298 7068 3698 10361 7922 9843 12059 8682	137 486 50 83 3 89 128 38 382 863 587	474669 475777 477730 479596 473207 478035 459159 468676 467459 459030 467157	2085 14451 16131 24341 19455 10702 15138 3463 3827 22292 24734	42.1 35.7 39.0 32.4 40.2 47.6 124.3 61.0 41.0 77.0 62.0	0.7 0.3 0.5 0.5 0.8 3.0 0.1 1.2 2.6 1.3	443.2 132.3 50.5 23.7 8.9 1.1 0.9 28.9 72.2 11.4 21.6	12.2 12.7 6.5 5.7 12.6 3.1 5.5 8.9 10.0 5.1 6.1	733.5 265.7 143.9 74.5 76.8 20.9 66.4 96.7 160.5 99.8 81.6	16.1 23.1 0.4 0.6 3.2 0.2 4.8 1.9 8.5 18.7 3.7	61.1 24.0 12.4 4.7 4.3 0.9 1.5 9.9 23.2 7.1 9.2	1.8 2.1 0.1 0.5 0.1 0.1 0.4 1.2 1.4 1.0	195.2 86.2 50.3 19.4 26.2 5.7 11.5 50.3 111.4 40.7 48.8	3.0 11.8 5.0 0.2 1.5 0.4 2.5 1.6 5.9 8.9 3.4	41.0 26.2 22.8 11.5 30.1 8.0 22.7 33.9 52.9 40.2 34.0	0.8 1.6 2.3 1.6 2.6 0.9 1.2 1.8 2.5 8.4 2.8	0.55 0.68 0.42 0.50 0.92 0.31 0.84 0.64 1.51 0.99 0.76	0.04 0.25 0.06 0.04 0.08 0.11 0.01 0.02 0.01 0.09 0.09	243.4 134.5 98.5 57.4 133.2 49.5 132.8 141.4 192.3 195.3 148.4	21.2 1.7 5.2 1.1 1.6 1.5 6.1 7.5 8.6 28.4 13.7	33.9 33.2 26.7 18.5 43.3 23.6 51.6 45.9 51.0 67.6 48.6	0.04 2.0 1.2 0.8 1.3 0.6 2.2 0.04 2.4 6.1 3.4	359.8 373.0 298.5 211.0 475.9 316.9 630.7 513.7 538.2 742.8 550.6	11.1 23.1 0.2 2.4 14.8 14.5 23.8 21.5 0.5 64.6 35.9	129.6 135.9 108.5 77.0 172.3 116.0 239.8 185.5 194.7 276.5 201.0	0.6 5.6 2.5 1.2 0.3 0.4 2.1 0.6 5.3 30.5 5.5	588.3 628.7 496.1 355.6 753.4 546.2 1137.1 822.1 868.4 1234.9 903.2	8.4 34.7 1.9 10.7 21.4 3.5 20.7 12.0 52.8 160.0 35.8
DEC13-09 112_KMS 22_MT 25_MT 25_MT 202_KMS 202_KMS 202_KMS 21_MT 207_KMS 204_MT 207_KMS 208_KMS 25_BS 208_KMS 25_BS 26_KMS 21_BK 209_KMS 21_LKMS	65.568 65.584 65.608 65.623 65.623 65.634 65.635 65.635 65.645 65.656 65.656 65.656 65.659 65.659 65.659 65.653 65.653 65.721	0.031 0.041 0.049 0.034 0.028 0.038 0.038 0.038 0.038 0.038 0.034 0.034 0.032 0.155 0.045 0.045 0.025 0.034	1.02 0.92 1.20 1.11 0.97 1.05 1.15 1.01 0.92 0.97 1.13 1.02 1.10 0.99 1.04 0.94 1.06 1.11	7912 9226 10259 2244 6866 6705 13199 11652 4867 5205 4754 2919 6741 5803 6136 3498 2810 1298	63 1558 1029 347 221 1421 1216 3552 27 3309 385 17 689 165 983 113 873 396	469799 463870 462625 480417 468820 472700 456780 459081 477202 473948 477314 480047 471779 473505 472215 477821 480546 482499	3283 4632 1016 18435 8142 4819 30544 21639 53223 73437 33091 39244 29356 31873 15927 8216 769 4245	94.8 59.1 54.9 92.4 125.6 93.1 63.3 62.7 19.1 140.1 56.4 35.2 82.3 29.2 114.7 53.7 64.4 48.5	0.6 4.1 4.8 2.2 2.8 4.2 2.7 4.0 0.4 17.3 2.5 1.2 4.9 0.7 2.2 4.9 0.7 2.0 1.1	6.3 b.d.l. b.d.l. 27.1 15.3 b.d.l. b.d.l. b.d.l. b.d.l. 44.8 372.9 2.7 b.d.l. 44.8 44.3 0.5 337.2	10.5 b.d.1 b	68.4 54.4 89.9 7.0 115.5 83.1 67.9 73.0 48.5 10.3 46.4 515.9 68.1 58.3 158.9 87.0 8.7 562.0	4.3 1.7 7.3 0.6 3.8 20.7 5.5 5.5 5.5 5.5 7.6 9.3 3.9 9.3 3.9 2.3 40.2 5.1 3.3 121.2	2.5 0.5 0.4 0.2 12.9 8.8 0.5 0.4 0.2 0.1 1.8 64.4 2.6 16.4 5.3 0.4 61.3	0.1 0.1 0.0 1.4 2.0 0.1 0.2 0.1 0.2 2.4 0.3 2.7 0.4 0.2 13.1	16.8 6.4 8.5 1.9 67.0 45.2 7.9 7.0 4.7 1.9 12.5 208.0 17.3 6.8 79.6 17.1 2.8 201.7	2.2 0.7 0.4 0.1 10.3 0.2 1.4 0.8 0.9 2.0 2.5 3.7 19.3 1.2 39.9	22.8 17.2 26.8 4.5 35.9 31.3 27.9 23.8 13.9 7.8 17.1 40.6 21.6 14.6 40.6 9.7 6.2 37.2	1.8 1.3 1.9 0.2 0.9 6.4 0.1 5.7 0.1 4.3 0.4 5.3 0.4 5.3 0.4 0.7 14.6 0.6 2.4 4.6	0.69 0.63 0.59 0.34 0.86 0.81 0.29 0.44 0.54 0.54 0.54 0.54 0.54 0.54 0.54	0.30 0.07 0.10 0.11 0.14 0.14 0.15 0.12 0.37 0.27 0.35 0.04 0.35 0.04 0.13 0.06 0.22 0.02 0.06	118.5 111.1 162.5 30.4 134.9 125.1 167.0 146.5 74.5 60.7 80.3 173.4 110.1 93.0 139.2 61.2 40.0 152.4	1.0 3.3 9.4 1.2 19.5 20.6 3.9 34.8 0.5 31.0 4.9 7.5 5.0 2.1 24.2 4.4 14.9 25.1	43.2 44.2 57.8 15.2 41.7 44.4 64.7 57.0 25.6 32.9 28.8 23.1 39.7 32.2 40.2 22.2 20.2 15.2	0.7 2.1 2.8 0.4 3.9 7.1 5.3 11.1 1.1 1.5 2.2 0.5 3.6 0.5 7.0 0.9 7.0 3.3	491.1 539.6 650.8 185.7 461.9 480.2 772.0 674.4 299.3 420.3 307.7 206.4 426.1 380.3 416.7 292.6 245.4 138.0	7.7 12.1 112.0 7.1 59.9 55.4 35.7 135.4 3.4 185.5 15.2 25.2 39.7 23.7 89.2 21.7 89.2 21.7 89.2 21.7	186.0 208.2 236.1 174.5 171.6 284.2 246.7 110.3 160.3 114.0 74.9 155.4 136.7 151.1 105.7 88.7 45.9	0.1 7.5 22.4 2.3 24.0 26.2 11.5 53.7 3.2 71.0 4.4 5.9 13.6 3.0 24.9 11.7 32.6 6.7	849.9 971.3 1041.7 316.8 777.7 757.1 1307.9 1139.7 502.1 763.5 506.1 337.6 700.7 622.5 682.6 481.7 401.5 217.3	23.8 47.0 56.7 4.9 121.7 144.1 95.5 206.5 23.4 344.8 24.1 33.2 104.9 3.5 136.6 30.5 148.8 40.7
DEC13-10 211_KMS 29_MIT 22_MIT 24_BS 24_BS 25_BS 25_BS 28_MIT 21_BS 210_MIT 26_MIT 210_MIT 20_KMS 207_KMS 209_KMS	65.508 65.581 65.594 65.597 65.630 65.641 65.641 65.644 65.644 65.712 65.712 65.712 65.717 65.757 65.824	0.081 0.049 0.084 0.035 0.059 0.038 0.099 0.049 0.048 0.048 0.020 0.104 0.205 0.069 0.114 0.083	0.93 0.98 1.16 1.01 1.11 1.00 1.13 0.99 0.96 0.91 0.91 1.00 1.50 0.94	8546 8629 6787 10123 6711 8754 9504 7361 7087 9101 8449 2809 9310 2269 7733	1832 72 156 42 1233 2567 1344 404 115 737 274 41 1187 260 31	467144 463771 474325 46432 472289 466445 469736 472851 466692 46692 466570 483750 464094 483750	20322 5642 614 25557 47703 35451 29220 5341 18702 19773 20260 20657 31230 3357 3644	83.8 65.5 16.8 50.7 65.4 93.2 107.4 76.6 18.1 108.7 79.2 3.7 91.0 4.9 108.4	2.4 1.2 1.4 5.7 4.6 5.7 0.6 0.3 0.3 0.5 0.4 1.6 0.6 1.3	b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. 0.5 b.d.l. 1.2 b.d.l. 4.1 b.d.l.	b.d.l. b.d.l. b.d.l. 23.2 b.d.l. b.d.l. 7.2 b.d.l. b.d.l. 14.0 b.d.l. 14.0 b.d.l. 40.2 b.d.l.	58.9 49.9 57.6 78.7 89.9 62.3 65.8 58.8 58.8 58.8 58.9 31.2 48.9 24.9 74.1 64.4 64.7	7.4 2.3 5.8 13.4 14.2 9.7 1.4 2.0 6.0 1.9 0.8 17.6 6.1 2.8	1.9 0.4 1.7 1.3 5.5 2.5 0.7 1.2 0.4 0.7 1.6 2.3 4.0 26.4	0.3 0.1 0.6 0.2 1.4 0.5 0.2 0.1 0.3 0.2 0.1 0.3 0.2 0.5 0.4 18.9	12.7 4.5 10.0 12.3 29.9 10.0 13.2 8.7 11.8 4.9 3.9 10.9 10.9 12.2 19.8 129.0	1.1 1.6 3.3 6.6 0.8 3.0 1.4 0.8 1.0 1.2 1.5 3.9 1.3 161.6	19.5 17.0 22.3 28.6 25.4 20.5 17.3 20.1 23.0 18.2 16.9 12.2 20.0 14.2 21.6	4.4 1.5 0.8 0.6 6.3 8.2 3.9 0.4 0.2 2.0 1.8 2.5 2.4 0.5 3.8	0.42 0.53 1.64 0.72 0.35 0.36 0.47 0.40 0.69 0.66 0.83 1.62 0.45 1.17 0.24	0.14 0.02 0.23 0.11 0.04 0.03 0.88 0.06 0.21 0.01 0.05 0.07 0.09 0.78 0.32	124.8 106.8 116.7 150.6 119.9 121.8 120.3 117.9 105.0 105.6 103.1 49.8 140.0 51.1 121.9	18.0 3.5 8.5 7.5 12.3 28.2 14.9 17.5 6.8 17.2 7.9 3.6 30.4 0.8 7.7	45.6 40.5 52.4 40.5 45.6 45.4 43.3 35.2 45.3 39.8 15.5 50.0 13.4 42.6	6.2 0.2 7.1 13.7 8.2 1.0 0.0 5.0 0.7 0.4 1005 2.9 0.7	562.0 537.4 443.7 662.9 466.0 540.1 569.1 509.1 446.8 549.6 525.2 179.5 593.9 133.3 508.5	46.1 0.2 1.7 14.1 75.7 703.2 74.3 8.2 2.0 47.6 24.6 10.3 115.5 9.0 49.4	202.7 205.9 245.1 169.8 201.5 215.4 188.7 163.7 214.0 198.7 66.6 214.2 49.0 187.6	25.1 7.5 10.0 3.9 28.6 44.7 37.2 7.8 5.5 10.7 19.3 3.9 41.8 4.1 2.3	916.6 970.5 717.7 1115.1 912.9 975.4 836.5 758.7 991.9 931.4 304.5 988.4 227.3 834.7	122.9 1.9 37.1 23.0 133.9 179.5 170.3 35.0 3.2 78.6 33.9 9.5 64.0 21.7 47.2
RBF z2_BS z03_KMS z3_BS z06_KMS z4_BS z5_BS z01_KMS z02_KMS z04_KMS z1_BS	65.537 65.557 65.594 65.695 65.639 65.667 65.674 65.714 65.751	0.181 0.093 0.041 0.034 0.034 0.037 0.067 0.041 0.041	0.92 0.87 0.99 0.95 0.97 1.04 0.92 0.98 1.43 1.03	2520 2200 3227 6964 8936 10517 9590 8713 6262 12156	49 116 23 401 679 828 125 535 486 2716	484682 484661 482756 471277 466867 461585 463394 465423 477521 450004	7633 14618 12337 22802 21655 27732 3792 17287 33166 29474	4.6 5.8 7.7 76.6 103.6 95.4 85.2 18.5 201.8	0.3 0.4 0.3 4.4 3.6 2.6 1.1 0.5 1.0 16.9	14.7 29.8 b.d.l. 22.8 b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l.	19.9 8.3 b.d.l. 26.4 b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l.	52.2 85.6 14.8 42.6 51.1 50.8 54.6 35.5 205.5	0.9 8.2 1.3 3.6 6.5 8.8 3.6 7.7 3.3 41.4	5.8 10.1 0.6 0.7 2.1 0.5 0.7 1.1 1.0 3.7	0.2 1.1 0.1 0.2 0.2 0.1 0.2 0.3 0.1 0.7	25.0 45.0 6.7 4.1 12.9 7.1 6.9 7.1 13.3 29.0	0.6 1.8 0.3 0.8 2.2 0.8 0.5 3.7 0.3 3.8	15.7 16.8 12.6 16.1 22.3 22.9 20.0 16.2 26.4 55.6	1.1 0.1 0.2 2.8 3.8 3.3 1.3 2.3 4.2	1.54 1.53 1.40 0.62 0.87 0.75 0.79 0.54 1.97 1.82	0.12 0.25 0.16 0.08 0.03 0.07 0.25 0.09 0.69 0.44	56.0 58.2 55.6 99.0 117.8 140.9 124.3 116.9 119.4 311.3	4.2 0.7 6.7 9.9 5.8 9.7 5.2 10.9 2.6 44.9	16.5 13.7 18.6 35.7 43.9 52.7 47.5 43.0 40.0 99.6	1.3 0.9 0.2 1.3 5.6 4.4 4.8 4.9 18.0	183.1 146.3 217.0 445.7 539.8 658.5 595.5 532.7 445.3 1090.7	8.5 1.9 11.1 24.3 25.9 31.5 1.4 46.1 12.6 178.4	64.7 51.5 76.3 166.1 210.4 242.7 223.0 204.7 155.0 371.5	1.1 2.2 5.1 10.0 7.9 16.2 12.9 30.7 11.7 44.4	292.6 235.5 343.4 766.7 958.5 1130.6 1027.4 915.2 667.5 1540.9	34.3 8.4 24.6 47.5 72.4 83.0 66.6 36.9 43.5 262.8
RBE 26_MIT 27_MIT 23_MIT 210_KMS 25_BS 22_BS 24_BS 24_BS 24_MIT 201_KMS 24_MIT 208_KMS 204_KMS 23_BS 23_BS 202_KMS	65.394 65.492 65.502 65.511 65.566 65.578 65.600 65.614 65.622 65.648 65.670 65.724	0.330 0.225 0.245 0.057 0.065 0.146 0.188 0.050 0.145 0.078 0.070 0.063 0.056	0.92 0.94 0.92 1.06 0.94 1.02 0.88 0.96 0.95 0.92 0.93 0.88 0.95 0.95 0.90	2721 2899 2765 3347 1737 3374 2395 2783 2853 2783 2853 2754 2267 2347 2659 2315	118 91 996 102 830 48 217 719 173 418 152 107 152 118	483597 483672 484385 482581 485549 483345 484015 484431 484015 484191 484556 484312 484322	4825 30160 76218 12246 12035 3571 22963 17614 7253 37629 10843 31932 2033 10263	2.8 4.0 8.3 9.7 7.6 5.0 4.9 8.3 5.4 3.9 9.3 9.5 5.2 5.3	0.2 0.3 1.5 0.4 0.4 0.4 0.4 0.6 0.1 0.6 0.1 0.6 0.1 0.5 0.3 0.1	4.4 b.d.l. 2.6 0.5 b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. 3.4 0.7 b.d.l. 1.2	14.5 b.d.l. 5.8 3.2 b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. 6.4 6.2 b.d.l. 5.2	30.0 23.7 11.6 19.6 12.7 12.2 7.1 9.1 12.9 7.7 19.6 9.2 4.8 10.3	0.6 1.8 2.8 2.8 1.6 1.4 2.0 0.5 1.8 0.6 2.1 1.3 1.9	2.3 1.2 0.5 2.0 0.5 0.8 0.3 0.6 0.9 0.3 1.8 1.2 0.3 1.0	0.2 0.3 0.1 0.4 0.1 0.2 0.1 0.2 0.1 0.2 0.1 0.1 0.1 0.4 0.0 0.3	13.3 9.6 6.6 13.2 3.2 8.0 4.1 5.9 7.2 3.9 12.1 7.9 3.8 7.3	3.0 0.8 1.2 0.3 1.3 0.8 0.3 0.8 0.1 1.0 1.3 0.8 0.5 0.4	12.6 11.3 10.7 15.5 4.5 14.4 9.4 10.7 11.8 9.0 10.3 9.1 9.4 8.7	1.0 2.0 2.7 0.3 1.6 1.2 0.2 1.3 1.4 2.3 1.0 1.1 0.9 2.0	1.51 1.34 1.61 2.11 0.85 2.09 1.45 1.54 1.54 1.54 1.54 1.21	0.40 0.37 0.40 0.39 0.18 0.23 0.44 0.16 0.06 0.09 0.41 0.05 0.29 0.03	50.5 50.0 43.7 61.6 25.3 38.4 44.0 48.7 41.2 43.1 40.7 41.5 39.4	1.0 0.4 6.8 5.5 6.9 2.6 3.6 3.9 1.8 1.4 0.1 1.5 2.3 0.7	16.1 17.3 15.8 21.3 10.2 18.7 12.8 14.7 15.6 15.3 14.0 14.4 15.0 12.9	1.3 0.6 1.8 0.7 2.4 1.1 1.7 0.4 0.5 0.1 1.9 1.8 1.2 0.5	175.6 186.7 172.2 228.7 121.6 226.3 158.0 162.2 187.0 164.8 159.4 163.7 179.2 152.0	2.9 8.9 25.7 7.2 34.0 11.8 14.8 3.8 7.1 5.0 17.3 13.0 12.7 0.2	62.6 66.9 82.0 43.3 78.3 55.4 57.7 67.1 59.5 56.7 56.9 62.9 54.4	0.9 1.5 7.4 4.9 11.4 3.7 3.1 4.0 4.9 0.2 6.0 4.7 6.1 2.3	281.7 301.4 274.9 374.1 202.9 352.5 249.4 256.7 267.9 259.6 267.4 265.4 265.4 265.4 265.4 265.4	0.4 12.8 35.5 10.2 61.5 2.7 17.0 13.7 9.0 6.5 30.5 20.5 15.0 6.3
DEC13-30 z4_BS z10_KMS z5_BS z6_BS z10_BS z08_KMS z07_KMS z09_KMS	66.286 66.289 66.291 66.325 66.343 66.390 66.465 66.478	0.035 0.077 0.042 0.024 0.031 0.024 0.037 0.122	2.33 2.62 1.69 2.47 1.90 2.27 2.35 2.14	9432 15016 7478 11086 8856 10979 15036 10805	1445 976 569 813 47 622 276 451	414504 399515 433312 409009 430999 414471 399181 413427	7254 35614 22966 9809 1885 29239 3415 14488	23.4 144.4 51.7 32.9 57.6 23.6 24.7 17.8	0.3 2.6 1.0 0.8 1.1 0.4 0.4 0.3	b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l.	b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l. b.d.l.	137.9 204.4 108.8 164.8 112.6 121.9 226.8 95.1	17.6 15.1 7.7 9.5 3.7 3.6 2.6 2.7	2.7 3.9 4.3 2.6 3.8 4.7 9.5 8.0	0.5 2.4 1.4 0.3 0.7 0.3 0.2 0.4	31.3 37.3 37.6 50.8 65.1 93.8 74.1	4.2 12.6 6.7 6.6 1.3 5.1 7.1 4.2	69.2 119.8 66.8 77.6 97.4 118.4 149.0 103.6	12.5 22.4 5.3 0.6 9.6 9.3 4.9 1.7	2.17 4.67 2.26 2.21 3.84 6.41 12.86 12.15	0.83 3.41 1.00 0.21 1.27 1.31 0.35 0.95	272.3 468.2 237.5 308.7 323.5 385.5 507.6 334.5	31.6 52.0 12.9 1.2 4.1 19.7 11.4 4.7	84.6 145.1 69.3 98.0 98.0 115.3 155.1 104.1	11.1 9.3 0.6 6.6 0.6 1.1 1.4 1.7	809.5 1275.0 636.6 914.4 821.2 1017.0 1368.3 937.0	85.3 173.7 6.4 32.7 2.9 17.7 59.1 13.3	236.7 366.8 173.4 267.9 223.3 274.2 373.2 269.3	33.4 3.8 0.2 15.0 1.6 9.1 6.9 2.6	897.9 1337.2 642.4 989.0 789.2 953.9 1365.7 1023.4	89.4 68.8 9.5 7.7 47.9 20.8 15.5
RBP z1_BS z4_BS z5_BS z2_BS	65.512 65.653 65.673 65.787	0.199 0.128 0.380 0.183	1.08 1.19 1.17 1.02	3574 1427 4521 3191	19 7 144 3	481590 483983 481110 482336	19178 30868 28066 1443	2.0 2.0 b.d.l. 1.9	0.6 0.2 44.6 0.4	b.d.l. b.d.l. b.d.l. b.d.l.	b.d.l. b.d.l. b.d.l. b.d.l.	20.1 14.3 b.d.l. 23.2	2.7 0.9 b.d.l. 2.1	1.2 0.5 3.9 1.3	0.8 0.2 29.3 0.3	9.0 3.0 b.d.l. 7.8	1.1 0.9 b.d.l. 0.9	11.8 4.2 b.d.l. 12.5	2.1 0.9 b.d.l. 0.6	2.09 1.09 32.91 2.04	0.15 0.61 6.18 0.52	57.3 24.1 81.6 54.6	7.7 0.9 72.3 5.7	19.9 7.5 26.1 18.2	2.9 0.1 4.3 1.5	229.8 81.7 61.9 205.4	7.9 6.5 28.1 1.8	80.4 31.8 148.2 74.7	3.0 1.7 6.1 2.3	361.0 147.6 266.2 331.9	19.7 10.4 24.6 13.6

a. sample name followed with specific zircon listed beneath. \_BS and \_CMS indicate grains measured at PU and \_MIT are those measured at MIT. b. dates and ThU derived from TMS analysis, see Table S1 c. total concentration of element tisted with atobep measured indicated. Data are presented as ppm in zircon, by using concentrations measured in solutions and normalizing to stoichiometric 2+HF+REE... d. uncertainties include analytical uncertainties and blank subtraction, based on measured variability in total procedural blanks

Tm-169	+2cia	Vb-172	+2sig	10-175	+2sig	Hf.178	+2sia	Ta-181	+2sia	Th.232	+2sia	
1111 200	12318	10 172	12.00	20 175	1100	111 270	123.8	10 101	77.76	111 2.52	22.00	а
												DEC13-08
110.0	2.0	954.1	22.5	165.6	0.7	10200	286	5.7	0.1	2286	24	24 BS
114.8	5.7	934.3	110.2	180.1	9.6	10388	227	5.5	0.3	3008	88	73 BS
93.2	14	727.8	21.2	139.6	11	10516	177	6.5	0.1	2547	18	210 KMS
66.0	3.0	545.7	14.6	103.8	19	11057	186	83	0.1	2059	32	208 KMS
135.4	0.4	1051.4	45.2	199.6	1.0	9397	277	5.4	0.1	4750	100	206 KMS
104.2	5.9	886.6	17.2	158.1	5.8	11319	303	7.8	0.4	2298	53	z1 BS
215.3	8.6	1704.4	28.3	324.0	18.8	12380	92	19.4	0.3	10957	285	711 KMS
150.7	0.2	1218.2	37.3	222.6	2.4	10822	245	9.5	0.1	6536	221	205 KMS
152.8	3.8	1216.1	17.0	235.1	1.8	10158	163	6.6	0.1	6246	109	202 KMS
220.2	27.2	1729.9	129.0	225.2	20.0	11102	649	12.9	0.6	10227	690	107 KMS
164.7	93	1287.9	70.6	244 1	12.9	11019	377	10.2	0.0	6924	416	204 KMS
												DEC13-09
159.5	17	1201.2	72.6	241.7	2.1	10557	07	12.0	1.6	5673	108	*12 KMS
196.0	1.7	1536.2	22.0	241.7	11.0	10004	505	13.0	1.5	9409	108	212_KWI3
100.9	9.5	1520.2	107.6	207.3	11.9	10010	300	14.5	0.3	0560	200	22_WIT
154.0	0.0	1340.0	107.0	250.8	10.0	10810	/21	12.0	0.7	9309	033	23_WIT
140.5	15.5	1162.9	197.0	32.0	25.7	12155	170	17.4	0.0	6247	050	12_03 x03_KM6
140.5	19.2	1103.8	80.1	218.1	24.9	11075	612	12.2	0.4	2710	627	*10 KMS
241.2	17.5	1032.4	162.4	210.1	24.8	11075	380	13.2	0.0	11242	627	210_KWI3
241.2	42.7	1732.4	260.0	221.7	34.4	11070	1250	11.5	2.1	10794	333	21_WIT
213.8	42.7	760.0	12.6	145.2	1.2	10650	1230	13.7	2.1	10784	2284	23_WIT
33.4	2.0	109.9	12.0	143.2	1.2	10039	4430	3.0	0.1	2803	39	24_WIT
149.7	/1.5	1295.2	357.9	142.7	106.6	10265	1439	18.2	1.9	28/5	1143	207_KMS
63.6	0.0	525.9	2.0	143.7	0.9	10200	343	5.6	0.3	1546	197	204_KW3
121.0	4.4	1062.1	110.6	200.2	16.2	110570	278	11.0	0.1	1340	24	206_KWI3
151.0	0.1	1002.1	110.0	200.5	10.5	11001	339	11.0	0.0	3034	333	23_63
114.6	1.0	934.3	38.3	179.0	5.8	11308	200	8.4	0.1	4412	3/3	20_MIT
04.1	12.2	772.2	62.2	140.0	14.2	112/2	200	12.2	0.2	4780	501	200_KWI3
34.1	12.5	773.2	02.5	140.9	14.2	11413	235	5.5	0.4	2701	03	21_65
/6.1	20.1	670.6	246.9	118.4	42.6	11013	645	8.5	0.6	1503	542	209_KMS
40.0	7.0	343.3	85.4	60.4	10.9	11032	196	7.8	0.1	505	73	ZII_KMS
												DEC13-10
178.5	23.5	1357.9	143.2	263.1	47.9	11450	357	15.5	1.5	6618	923	z11_KMS
187.8	0.5	1508.2	78.7	278.9	8.7	12117	340	13.9	1.0	9140	437	29_MIT
141.2	4.5	1116.2	33.1	213.7	11.1	9209	261	4.8	0.5	4263	110	z7_MIT
206.9	9.9	1639.7	51.5	301.5	8.7	11085	169	8.7	0.1	7750	255	z2_MIT
140.8	26.3	1091.6	178.8	211.0	36.7	10521	1104	11.0	1.0	4920	832	z4_BS
173.6	35.0	1362.1	337.1	259.5	57.7	11273	697	16.6	1.1	7332	1769	z3_BS
190.0	23.8	1545.3	229.7	287.6	38.0	11648	510	19.0	1.0	9023	1410	z5_BS
152.4	0.9	1211.4	62.5	228.4	11.5	10845	220	13.5	0.1	6222	848	z2_BS
144.8	0.8	1149.2	29.4	212.8	5.2	10536	422	2.0	0.2	4042	249	z8_MIT
187.8	9.3	1574.6	59.2	295.7	11.5	12031	154	14.9	0.1	5674	863	z1_BS
185.5	3.3	1522.8	59.9	273.7	5.3	11904	732	16.0	1.9	7567	422	z10_MIT
59.4	1.8	482.9	13.9	95.1	5.8	8836	342	1.2	0.1	940	43	z6_MIT
185.8	17.9	1409.1	203.8	271.5	32.9	11458	861	17.5	0.4	8626	2016	z10_KMS
46.9	5.1	396.3	40.4	84.7	3.5	8605	475	1.9	0.8	3154	156	207_KMS
158.6	7.6	1259.1	32.4	238.5	28.1	10581	487	12.4	0.9	5291	258	209_KMS
												RBF
56.7	1.7	449.4	4.8	92.8	2.2	8430	261	1.3	0.2	682	23	z2_BS
46.1	3.4	379.1	23.5	76.7	7.1	8940	467	1.8	0.0	639	74	203_KMS
65.5	3.8	538.6	25.1	106.8	0.5	8964	32	2.2	0.2	1230	33	z3_BS
147.6	15.1	1161.1	5.6	222.8	16.2	11374	358	11.2	0.2	4791	326	206_KMS
183.3	3.3	1486.8	22.0	279.5	18.7	11154	620	12.6	1.5	6690	555	z4_BS
205.0	18.6	1658.9	87.2	310.5	34.4	11098	596	16.6	0.1	9815	874	z5_BS
200.1	15.2	1591.6	71.0	299.9	18.3	11965	103	16.8	0.9	8376	850	201_KMS
180.3	18.0	1395.5	124.8	264.0	9.6	11469	535	16.0	0.2	8162	978	z02_KMS
127.7	7.8	989.8	47.9	191.6	21.0	8019	628	3.6	0.1	2985	238	z04_KMS
268.7	54.4	2055.1	315.0	369.9	70.8	12313	897	12.3	0.4	16510	2431	z1_BS
												RBE
54.2	1.7	455.8	7.9	90.0	0.3	9189	223	1.8	0.1	883	75	z6 MIT
57.5	1.0	486.7	9.6	95.4	1.9	8645	116	2.0	0.2	1116	67	z7_MIT
51.6	8.4	442.1	61.7	87.4	10.8	8618	768	2.0	0.1	687	98	z3_MIT
69.2	3.2	572.1	28.8	112.4	1.7	9058	341	1.7	0.1	1066	28	z10_KMS
39.4	11.0	332.6	70.5	64.3	17.3	9068	539	1.6	0.1	416	92	z5 BS
66.1	1.9	539.0	8.8	107.8	1.5	8553	274	1.5	0.1	884	56	z2_BS
47.9	5.3	407.4	42.3	81.3	12.6	8492	274	1.7	0.1	639	56	z1_BS
49.0	1.6	403.8	13.0	81.7	1.8	8466	166	1.9	0.0	846	75	z2 MIT
58.5	0.4	475.7	30.5	96.7	1.9	8601	298	1.8	0.0	885	59	z01_KMS
51.1	0.4	428.5	8.0	86.3	0.3	8769	469	1.5	0.1	808	139	z4_MIT
49.5	5.2	413.5	37.5	81.4	8.5	9023	157	1.5	0.1	659	16	z08_KMS
50.4	3.3	417.8	17.4	83.7	6.6	9248	165	1.7	0.0	597	24	z04_KMS
54.9	5.1	450.6	25.0	90.4	10.0	8512	167	1.6	0.2	739	63	z3_BS
46.9	1.0	398.5	22.1	80.0	0.1	9031	186	1.8	0.1	703	38	z02_KMS
												DEC13-30
154.5	18.3	1029.9	102.2	169.9	25.6	7361	142	2.7	0.3	62393	8886	z4_BS
222.7	8.9	1444.7	73.0	229.9	15.1	7688	50	2.7	1.2	69147	3676	z10 KMS
105.8	5.2	676.7	1.5	167.1	2.8	7033	309	1.9	0.5	46857	1830	z5_BS
162.6	17.6	1089.9	71.1	171.7	10.0	7551	463	3.1	0.5	65563	2322	z6 BS
129.6	5.9	842.5	3.7	133.3	7.2	7134	176	1.9	1.0	46846	120	z10_BS
148.7	4.3	980.9	19.6	139.8	3.6	7622	15	1.7	0.2	59880	1653	208 KMS
214.8	2.6	1415.1	0.1	217.0	2.3	7351	129	3.0	0.4	69759	1131	207_KMS
161.7	8.7	1086.4	25.2	175.8	2.7	7639	207	1.8	0.1	61196	466	209_KMS
												-
												RBP
70.5	5.0	538.3	17.0	111.5	2.2	8832	298	1.5	0.4	2126	126	z1 BS
30.8	2.3	255.2	16.4	53.8	3.4	10717	320	1.7	0.1	834	29	24 BS
249.8	106.7	664.8	660.5	552.0	228.1	9828	17941	hdl	h d l	hdl	h d l	25 BS
61.7	2.2	497.2	3.9	99.3	33	8896	89	13	0.2	1814	39	22 BS

### Table S3: U-Pb age interpretation

Constant (Th/U)	partition coeffici	ient betwee	en magma and zi	ircon of 0.33	and imposed (	uncertainty	of ±0.3 o	n Th/U of res	ulting magma	(preferred in	nterpreta	tion)						
			- 	1									PREFERRED II	NTERPRETATION				
	# of analyses	deltaT (ka)	Youngest grain	+ 20	Weighted me	ean + 2/3	N	MSW/D	Geochemica	lly-guided wei	ghted mea	ns (IWMs)	Monte Carlo	ages from IWMs	Monte Carlo ag	trom single grain	Calculated durations and ages	+ 2a
DEC13-08 (PU)	11	150	65.574	0.028	65.586	0.018	4	1.2	65,583	0.020	3	1015000	failed superpo	osition	failed superpos	ition	duration of C29r (Monte Carlo	single grain)
DEC13-09 (PU)	12	150	65.568	0.031	65.606	0.015	4	2.9									754	56
DEC13-09 (MIT)	6	70	65.589	0.041	65.619	0.019	6	1	65.579	0.022	3	0.56	failed superpo	osition	failed superpos	ition	Minimum duration of Deccan (	Monte carlo single grain)
DEC13-10 (PU)	9	320	65.508	0.081	65.608	0.028	4	2.6									809	69
DEC13-10 (MIT)	6	220	65.581	0.049	65.602	0.024	5	1.6	65.566	0.037	3	1.3	65.535	0.027	65.477	0.059		
RBF (PU)	10	200	65.557	0.081	65.592	0.027	5	0.34	65.569	0.058	3	0.24	65.545	0.026	65.513	0.045	PREFERRED INTERPRETATION	(from Monte Carlo of IWMs)
RBE (PU)	11	250	65.502	0.057	65.512	0.039	5	0.5	65.511	0.037	9	0.50	65.552	0.026	65.532	0.044	duration of C29r	
RBE (MIT)	5	220	65.39	0.33	65.54	0.09	5	1.0					~ ~ ~ ~		<b>CE CO</b>		736	37
KBP (PU)	5	270	65.51	0.20	65.658	0.09	4	1.4	65.658	0.09	4	1.4	65.661	0.087	65.60	0.11	Minimum duration of Deccan	
DEC13-30 (PU)	15	180	66.286	0.035	66.309	0.017	3	2.1	66.288	0.027	2	0.03	66.288	0.027	66.286	0.035	/53	38
	Total average		Deters in held and		A make and in all	محجاه من الحاد			A set uses in al						(		Age of C30h/C29f reversal	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
	including ye	nocrysts but	Monte Carlo sim	e used for	chosen baser	daed in triese	weignteu i nce includi	ng the	chosen base	don both are	weignteu n	mical	Ages reported	Chain model that i	moses the		Age of C29r/C29n reversal	0.027/0.047/0.085
	deltaT - the	a difference	RRE RRE the vo	ungest mean	voungest gra	in		ing the	equivalence	of the selected	and geoene	norted dates	constraint of	stratigraphic contin	niposes trie		65 552	0 026/0 047/0 085
	between th	e oldest and	dates were signi	ficantly loss	youngestigra				for each sam	ole represent	combined s	weighted-	that stratigra	phically higher hed	s must be		05.552	0.020/0.047/0.005
	voungest gra	ain - does not	precise than the	2nd youngest					means of an	alvses from Mi	T and PU.	WM =	vounger. "Fai	led superposition" i	indicates that		Calculated durations for C29r a	re age of DEC13-30 minus RBE.
	include x	enocrysts	so the 2nd young	pest was					intelligent w	eighted means			< 0.1% of rand	iomly chosen paths	met these		Minimum duration of Deccan i	s calculated as DEC13-30 minus DEC13-
			chosen							0			criteria for th	ese samples, and w	ere thus		10. Uncertainties are added in	quadrature. Uncertainties for ages are
													discarded. Mo	onte Carlo IWMs ar	e the preferred		given as ±X/Y/Z, where X = ana	lytical, Y= +tracer calibration, Z = +decay
													interpretation	n that appear in the	text and in		constants	
													Figs. 1,2					
Constant magma	Th/U of 2.8 ± 0.	5 (NOT pre	fered interpreta	ation)														
	1		Voungest grain		Weighted m				Geochemica	llv-quided wei	abted mea	ne (IM/Me)	Monte Carlo	ages from IMMs	Monte Carlo a	as from single grain	Calculated durations	
	# of analyses	deltaT (ka)	date (Ma)	± 2σ	date (Ma)	± 2σ	N	MSWD	date (Ma)	± 20	N	MSWD	age (Ma)	± 2σ	age (Ma)	± 2σ	duration (kyr)	± 2σ
DEC13-08 (PU)	11	150	65.543	0.062	65.585	0.019	5	1.3	65.577	0.021	3	1.5	failed superpo	osition	65.496	0.043	duration of C29r (Monte Carlo	single grain)
DEC13-09 (PU)	12	150	65.565	0.033	65.604	0.016	4	2.5									687	63
DEC13-09 (MIT)	6	70	65.584	0.051	65.615	0.02	6	0.93	65.577	0.023	3	0.47	failed superpo	osition	failed superpos	ition	Minimum duration of Deccan	
DEC13-10 (PU)	9	320	65.508	0.082	65.561	0.048	2	2.6									736	64
DEC13-10 (MIT)	6	220	65.577	0.085	65.599	0.025	4	1.3	65.563	0.039	3	1.2	65.532	0.027	65.514	0.041		
RBF (PU)	10	200	65.560	0.081	65.592	0.029	5	0.29	65.57	0.058	3	0.2	65.542	0.026	65.531	0.04	duration of C29r (monte carlo	from IWM)
RBE (PU)	11	250	65.497	0.059	65.512	0.039	5	0.5	65.51	0.037	9	0.47	65.550	0.026	65.545	0.042	697	43
RBE (MIT)	5	220	65.39	0.33	65.54	0.09	5	1.0									Minimum duration of Deccan	
RBP (PU)	5	270	65.51	0.20	65.651	0.09	4	1.4	65.651	0.09	3	1.4	65.543	0.087	65.62	0.11	715	43
DEC13-30 (PU)	15	180	66.232	0.047	66.269	0.022	4	1.9	66.247	0.034	2	0.78	66.247	0.034	66.232	0.047		
	Total numbe	r of analyses.	Dates in hold are	e used for	Analyses incl	uded in these	weighted r	neans are	Analyses incl	uded in these	weighted n	neans are	Ages reported	d here are outputs f	from the Monte		Calculated durations for C29r a	re age of DEC13-30 minus BBE.
	including xe	nocrysts, but	Monte Carlo sim	ulations. For	chosen based	d age equivale	nce includi	ng the	chosen base	d on both age	and geoche	mical	Carlo Markov	Chain model that i	mposes the		Minimum duration of Deccan i	s calculated as DEC13-30 minus DEC13-
	deltaT - the	e difference	RBE, RBF, the yo	ungest mean	youngest gra	in.		0	equivalence	of the selected	grains. Re	ported dates	constraint of	stratigraphic contin	uity - namely		10. Uncertainties are added in	quadrature. Uncertainties for ages are
	between th	e oldest and	dates were signi	ficantly less					for each sam	ple represent	combined v	veighted-	that stratigra	phically higher beds	s must be		given as ±X/Y/Z, where X = ana	lytical, Y= +tracer calibration, Z = +decay
	youngest gra	ain - does not	precise than the	2nd youngest,					means of ana	alyses from Mi	T and PU.	WM =	younger. "Fai	led superposition" i	indicates that		constants	
			and the Dead converse	gest was					intelligent w	eighted means			<0.1% of rand	lomly chosen paths	met these			
	include x	enocrysts	so the zha young	Beschus														
	include x	enocrysts	chosen	Best mus									criteria for th	ese samples, and w	ere thus			
	include x	enocrysts	chosen	Best was									criteria for the discarded.	ese samples, and w	ere thus			
	include x	enocrysts	chosen	gest mas									criteria for the discarded.	ese samples, and w	rere thus			
	include x	enocrysts	chosen	gest was									criteria for the discarded.	ese samples, and w	rere thus			





आणि संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती

साहित्य अकादेमी

ह्यांच्या संयुक्त विद्यमाने आयोजिलेल्या

आधुनिक मराठी साहित्यावर लोकसाहित्याचा प्रभाव विषयक चर्चासत्रास आपण सादर आमंत्रित आहात.

सोमवार, 22 जानेवारी 2018

दृकश्राव्य सभागृह, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती

### कार्यक्रम

उद्घाटनाचे सत्र	:	स.10.00 ते 11.00 वा.
स्वागताचे भाषण	:	कृष्णा किंबहुने
		प्रादेशिक सचिव, साहित्य अकादेमी, मुंबई
प्रास्ताविक	:	मनोज तायडे
		मराठी विभाग प्रमुख, संत गाडगे बाबा अमरावती विद्यापीठ
उदघाटनाचे भाषण	:	मरलीधर चांदेकर
		ु मा. कुलगुरु, संत गाडगे बाबा अमरावती विद्यापीठ
अध्यक्षीय भाषण	:	विश्वनाथ शिंदे
		सुप्रसिद्ध मराठी साहित्यिक
	आ	गर
चहा	:	स. 11.00 वा.
पहिले सत्र	:	सकाळी 11.30 ते दुपारी 1.30 वा.
		कथनात्म साहित्य
अध्यक्ष	:	शशिकांत सावंत
निबंध	: .	दा.गो. काळे, रणधीर शिंदे
जेवण	: ·	दु. 1.30 ते 2.30 वा.
दुसरे सत्र	:	दपारी 2.30 ते 4.00 वा.
		नाटक आणि कविता
अध्यक्ष	:	मधुकर वाकोडे
<b>•</b> •		

निबंध ः प्रकाश खांडगे, गणेश चंदनशिवे, रफिक सुरज आभार

> या चर्चासत्राच्या निमित्ताने साहित्य अकादेमी आपल्या प्रकाशनांचे प्रदर्शन व विक्रीही आयोजित करीत आहे संपर्क : (022)-24135744, 24131948 / (0721) 2662173 ई-मेल : rs.rom@sahitya-akademi.gov.in



### साहित्य अकादेमी



तथा संत गाडगे बाबा अमरावती विश्वविद्यालय, अमरावती

के संयुक्त तत्त्वावधान में आयोजित

### आधुनिक मराठी साहित्यपर लोकसाहित्य का प्रभाव

विषयक परिसंवाद में आप सादर आमंत्रित हैं.

सोमवार, 22 जनवरी 2018

द्रकश्राव्य सभागार, संत गाडगे बाबा अमरावती विश्वविद्यालय, अमरावती

### कार्यक्रम

उद्घाटन सत्र	:	पूर्वाह्न 10.00 से 11.00 बजे
ंस्वागत भाषण	:	कृष्णा किंबहुने क्षेत्रीय सचिव, साहित्य अकादेमी, मुंबई
आरंभिक वक्तव्य	:	मनोज तायहे मराठी विभाग प्रमुख, संत गाडगे बाबा अमरावती विश्वविद्यालय
उद्घाटन वक्तव्य	:	मुरलीधर चांदेकर मा. कुलगुरु, संत गाडगे बाबा अमरावती विश्वविद्यालय
अध्यक्षीय वक्तव्य	:	विश्वनाथ शिंदे सुप्रसिद्ध मराठी साहित्यकार
	धन्य	ग्वाद ज्ञापन
चाय	:	पूर्वाहन 11.00 बजे
प्रथम सत्र	:	पूर्वाहन 11.30 से अपराहन 1.30 बजे
		कथनात्म साहित्य
अध्यक्ष	:	शशिकांत सावंत
आलेख	:	दा.गो. काळे, रणधीर शिंदे
भोजन	:	अपराहन 1.30 से 2.30 बजे
द्वितीय सत्र	:	अपराहन 2.30 से 4.30 बजे
		नाटक एवं कविता
अध्यक्ष	:	मधुकर वाकोडे
आलेख	:	प्रकाश खांडगे, गणेश चंदनशिवे, रफिक सुरज
	धन्य	ग्वाद ज्ञापन
इस अवसर	पर र	साहित्य अकादेमी अपने प्रकाशनों की
manfa	री एत	विकी भी आयोजित कर रही है

उत्तरापेक्षी : 022-24135744, 24131948 / (0721) 2662173 ई-मेल : rs.rom@sahitya-akademi.gov.in

क्र.संगाबाअवि/५१७/ 108 /२०१८ दिनांक :- २८/०७/२०१८

### विषय :- नुक्कड कथा अभिवाचन कार्यक्रम आयोजित करण्याबाबत.

डिजिटल साहित्य विश्वात सुमारे दिडशे लेखक आणि सव्वा लाखाहून अधिक वाचकांनी चालवलेली अनोखी चळवळ म्हणजे 'नुक्कड... !' नुकतेच औरंगाबाद येथे नुक्कड साहित्य लेखकांचे साहित्य संमेलन दिमाखात संपन्न झाले आणि याच सोहळ्यात जानेवारी २०१९ मध्ये तिसरे नुक्कड साहित्य संमेलन विदर्भात घेण्याचे निश्चित झाले आहे.संपूर्ण विदर्भातील नव्या लेखक-वाचकांपर्यंत नुक्कड साहित्य संमेलनाचा संदेश पोववण्यासाठी नुक्कड कथा अभिवाचन मालिका करण्याचे योजिले आहे. त्यात वाशिम, बुलढाणा, अकोला, अमरावती व यवतमाळ यांचा समावेश आहे. या मालिकेतील एक भाग म्हणजेच नुक्कड कथा अभिवाचनाचा कार्यक्रम अमरावतील घ्यावयाचा आहे. सदर कार्यक्रम पदव्युत्तर मराठी विभाग, संत गाडगे बाबा अमरावती विद्यापीठ व विदर्भ साहित्य संघ, शाखा अमरावती यांच्या संयुक्त विद्यमाने घ्यावयाचे योजिले आहे.

या कार्यक्रगाचे खास वैशिष्टचे हे की, सादर होणाऱ्या कथा वैदर्भीय लेखक-लेखिकांनी लिहिलेल्या आहेत आणि अभिवाचकही विदर्भातीलच आहे. त्यामध्ये साहित्यिक-नाटककार-मालिका लेखक विक्रम भागवत यांचा समावेश आहे. यांच्यासोबत डॉ. स्वाती धर्माधिकारी, प्रतिमा पंकज, आशा अभिजित मुळे, अभिषेक बेल्लारवार आणि अभय नवाथे यांचा सहभाग असणार आहे.

सदर कार्यक्रम १८ ऑगस्ट २०१८ रोजी सायंकाळी ५.०० वाजता विमलाबाई श्रोतृगृह येथे संपन्न होईल.

या कार्यक्रमाच्या आयोजनासाठी विद्यापीठाला करावयाचा खर्च खालीलप्रमाणे आहे.

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उपरोक्त खर्च रु. १५,०००/- हा विद्यापीठाच्या 4) साधारण निधी H) बेस्ट प्रॅक्टिसेस इन युनिव्हर्सिटी viii) मराठी भाषा संवर्धन उपक्रम या शीर्षातून करण्यात येईल. कृपया, उपरोक्तू कार्यक्रमाच्या आयोजनास व खर्च **रु. १५,०००/-** ह्यास प्र**शासकीय व वित्तीय मान्यता मिळावी**, विनंती

पुढील कार्यवाहीसाठी सादर.

Arr2 281. 7/18

02/N पान क्र. दि - १८ ऑगस्ट, 209८ रोजी जुक्ट्र केमा आभेषाखन कार्यक्रमास मबासकीय मान्मता प्राप्त बनाली आहे. वित्तीय मान्यता मिळव्याताही सादर. Horiziana M/109 माः विल्त व लेखा आद्येकारी. 01/08/2018 A. g. 313. 201.8.18 Ku- MUD (10/00) 3/8/18 84632 नियमान पालन अपश्नन प्रस्तावित प्रस्तावास २०१८ -19 मध्य रू 15000/- भी नोंद घेवुन वितीम C 22778 218/18 भान्मतरतव स्नाप् 10/8/18 11/28/18 11/28/18 311843-19 21.93. 3193701 To. or m. 32. A-2887 3-9-0 all - 38-12 - 29-14-8-18 A-- 319 78,1087 वि.प्र. मराही ख्यांमी नुक्तड कशा आमिवा वन् करीता आछी रू 15000/ ची मांगणी केली अन्तून आग्रेम मंजूरीरस्व साहर 1418 37 - 9 Ch. 37 Patrice 18 14/08/18 14/08/18 39-03-tizen 39-9 G. 3F





संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती आणि मराठी विज्ञान परिषद, मुंबई यांच्या संयुक्त विद्यमाने आयोजित

### विज्ञान कथालेखन कार्यसाळेचे उद्घाटन

दिनांक : 30 ऑगस्ट, 2018 वेळ : सकाळी 11.00 वाजता

उद्घाटक : प्रा. डॉ. मुरलीधर चांदेकर कुलगुरू, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती

डॉ. मनोज तायडे

मराठी विभागप्रमुख, पदव्युत्तर मराठी विभाग, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती

अ.पां. देशपांडे

कार्यवाह मराठी विज्ञान परिषद, मुंबई

स्थळ : दृकश्राव्य सभागृह, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती

मराठी विभाग, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती आणि मराठी विभाग, श्री शिवाजी महाविद्यालय, आकोट, जि. अकोला संयुक्त विद्यमाने आयोजित

> विद्यापीठ अनुदान आयोग पुरस्कृत कवितालेखन कार्यशाळा शहानूर दि. १ फेब्रुवारी ते ३ फेब्रुवारी २०१९

### कार्यक्रमाची रूपरेषा

दि. १.२.२०१९

सकाळी १२.३० ते २ : उद्घाटन समारंभ अध्यक्ष : मा. डॉ. मुरलीधर चांदेकर, कुलगुरू, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती उद्घाटक : मा. श्री. हर्षवर्धन देशमुख, अध्यक्ष, श्री शिवाजी शिक्षण संस्था, अमरावती प्रमुख अतिथी : कवी विठ्ठल वाघ, अकोला प्रमुख उपस्थिती : प्राचार्य डॉ. ए. एल. कुलट, श्री शिवाजी महाविद्यालय, आकोट

सूत्रसंचालनः डॉ. प्रणव कोलते

दुपारी २ ते २.३० : भोजन

दुपारी २.३० ते ३.४५ : मार्गदर्शन सत्र वक्ते : बालिका ज्ञानदेव, पुणे सूत्रसंचालनः डॉ. मोना चिमोटे

दुपारी ३.४५ ते ४ :

दुपारी ४ ते ५.३० : मार्गदर्शन सन्न वक्ते : संतोष पद्माकर पवार सूत्रसंचालन: डॉ. मोना चिमोटे

चहापान

रात्री ५.३० ते ६.३० : खुली चर्चा

रात्री ८.३० ते ९.३० : कवितावाचन

दि. २.२.२०१९

सकाळी ७ ते १०	:	जंगल सफारी
सकाळी १० ते १०.३०	:	नाश्ता-चहा
सकाळी १०.३० ते १२ वक्ते	:	मार्गदर्शन सत्र बबन सराडकर, अमरावती सूत्रसंचालनः डॉ. हेमंत खडके
दुपारी १२ ते १.३० वक्ते	:	मार्गदर्शन सत्र बालिका ज्ञानदेव, पुणे
दुपारी १.३० ते २.३०	:	भोजन
दुपारी २.३० ते ४ वक्ते	:	मार्गदर्शन सत्र डॉ. विठ्ठल वाघ, अकोला सूत्रसंचालनः डॉ. हेमंत खडके
दुपारी ४ ते ४.१५	:	चहापान
दुपारी ४.१५ ते ५.४५ वक्ते	:	मार्गदर्शन सत्र संतोष पद्माकर पवार
संध्या. ५.४५ ते ६.३०	:	खुली चर्चा
रात्री ८.३० ते ९.३०	:	कवितावाचन

### दि. १.२.२०१९

### सकाळी १२.३० ते २ : उद्घाटन समारंभ

### प्रतिमापूजन :

संत गाडगे बाबा
 डॉ. पंजाबराव देशमुख

### स्वागत

अध्यक्ष : मा. डॉ. मुरलीधर चांदेकर, कुलगुरू, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती
 उद्घाटक : मा. श्री. हर्षवर्धन देशमुख, अध्यक्ष, श्री शिवाजी शिक्षण संस्था. अमरावती

३) प्रमुख अतिथी : कवी विठ्ठल वाघ, अकोला

8) प्रमुख उपस्थिती : श्री. गजानन पुंडकर, उपाध्यक्ष, श्री शिवाजी शिक्षण संस्था, अमरावती

५) प्रमुख उपस्थिती : श्री. केशवराव मेतकर, उपाध्यक्ष, श्री शिवाजी शिक्षण संस्था, अमरावती

६) प्रमुख उपस्थिती :

७) प्रमुख उपस्थिती : प्राचार्य डॉ. ए. एल. कुलट

८) प्रमुख उपस्थिती : डॉ. मनोज तायडे, मराठी विभागप्रमुख

९) प्रमुख उपस्थिती : डॉ. विलास तायडे

30)

प्रास्ताविक : डॉ. मनोज तायडे, मराठी विभागप्रमुख, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती

### मनोगतः

- १) स्वागतपर : प्राचार्य डॉ. ए. एल. कुलट, श्री शिवाजी महाविद्यालय, आकोट
- २) शुभेच्छापर : श्री. केशवराव मेतकर, उपाध्यक्ष, श्री शिवाजी शिक्षण संस्था, अमरावती
- ३) शुभेच्छापरं : श्री. गजानन पुंडकर, उपाध्यक्ष, श्री शिवाजी शिक्षण संस्था, अमरावती
- ४) प्रमुख अतिथी : कवी विठ्ठल वाघ, अकोला
- ५) उद्घाटनपर : मा. श्री. हर्षवर्धन देशमुख, अध्यक्ष, श्री शिवाजी शिक्षण संस्था, अमरावती
- ६) अध्यक्षीय भाषण : मा. डॉ. मुरलीधर चांदेकर, कुलगुरू, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती
- ७) आभारप्रदर्शनः डॉ. विलास तायडे, मराठी विभागप्रमुख, श्री शिवाजी महाविद्यालय, आकोट

सूत्रसंचालनः डॉ. प्रणव कोलते

दि. ३.२.२०१९

सकाळी ८.३० ते ९ : चहा-नाश्ता

सकाळी ९.३० ते ११ : मार्गदर्शन सत्र

वक्ते : बबन सराडकर, अमरावती

सकाळी ३१ ते ३२.३० : समारोप समारंभ

अध्यक्ष : मा. डॉ. राजेश जयपूरकर, प्र-कुलगुरू, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती प्रमुख अतिथी : कविवर्य विठ्ठल वाघ प्रमुख उपस्थिती : डॉ. मनोज तायडे, प्राचार्य डॉ. ए. एल. कुलट, डॉ. विलास तायडे, सूत्रसंचालनः डॉ. हेमंत खडके

दुपारी ३२.३० ते ३२.४५ : प्रमाणपत्र वाटप

दुपारी ३२.४५ ते ३.३० : भोजन

पदव्युत्तर मराठी विभाग, संत गाडगेबाबा अमरावती विद्यापीठ, अमरावती पदव्युत्तर मराठी विभाग, शासकीय विदर्भ ज्ञान विज्ञान संस्था, अमरावती आणि विदर्भ साहित्य संघ, शाखा : अमरावती यांचे संयुक्त आयोजन

### ।। कवितासमय।।

(सात समकालीन कवींचे कवितावाचन) अध्यक्ष : चंद्रकान्त पाटील (पुणे)

### । सहभागी कवी।

वसंत आबाजी डहाके (अमरावती) । प्रभा गणोरकर (अमरावती) प्रफुल्ल शिलेदार (नागपूर) । श्रीधर नांदेडकर (औरंगाबाद) राजा होळकुंदे (लातूर) । रवी कोरडे (जितूर) सूत्रसंचालनः प्रो. हेमंत खडके (अमरावती)

### । विशेष उपस्थिती ।

प्रो. रणधीर शिंदे (कोल्हापूर) । डॉ. आसाराम लोमटे (परभणी) डॉ. रामचंद्र काळुंखे (औरंगाबाद) । श्री. तुषार बोडखे (औरंगाबाद) प्रो. मनोज तायडे (अमरावती) कवितावाचनाच्या या कार्यक्रमास आपली उपस्थिती प्रार्थनीय आहे.

### ।। विनीत।।

प्रो. केशव तुपे प्रमुख, पदव्युत्तर मराठी विभाग शासकीय विदर्भ ज्ञान विज्ञान संस्था, अमरावती.

प्राचार्य डॉ. रमेश अंधारे अध्यक्ष, विदर्भ साहित्य संघ शाखा : अमरावती

प्रो. मोना चिमोटे प्रमुख, पदव्युत्तर मराठी विभाग संत गाडगेबाबा अमरावती विद्यापीठ, अमरावती.

> स्थळ : टाऊन हॉल, राजकमल चौक, अमरावती. दिनांक : ०७ सप्टेंबर २०१९, शनिवार / सायंकाळी ६.०० वाजता.

24

पुस्तक प्रकाशन : रई - प्रा. महादेव लुले, व-हाडधन - शिवलिंग काटेकर, सातबारा कोरा - दयाराम निंबोळकर, जागल - अरुण विघ्ने

संमेलन आयोजन समिती: डॉ.मोना चिमोटे, डॉ.मनोज तायडे, डॉ.माधव पुटवाड, डॉ.हेमंत खडके, डॉ. प्रणव कोलते, पुष्पराज गावंडे, प्रा.सदाशिव शेळके, दयाराम निंबोळकर, श्याम ठक, प्रा.महादेव लुले, निलेश कवडे.

समन्वयकः निलेश कवडे, रविंद्र दळवी, तेजस्वी बारब्दे, सौ.अनुराथा धामोडे, सौ.उज्ज्वला कांबळे, विजय पळसपगार, उमेश थोरात, अजय माटे, डॉ. पांडुरंग हटकर

### कोषाध्यक्षः प्रा.महादेव लुले

स्वागत समिती: आबासाहेब कडू, भास्करराव अरबट, शिवलिंग काटेकर, डॉ प्रमोद काकडे, नितिन बरणकार, प्रा.डॉ.सुनिल पखाले, निलेश देवकर, निलू मानकर, ॲड.मोहन जाधव, ॲड.गजेंद्र हिंगणकर, ॲड.ललित कदम, नितेश राऊत, अनिल लाले, विजय बिन्दोड, निळू तायडे, वीणा डिक्कर, युवराज टोपले, सुबोध दळवी, निलेश देशमुख, महेश मुळे

सभागृह तथा मंच व्यवस्थापन समिती: मनीषा माटे, ॲड.रजनी बावस्कर, धनश्री पाटील, सरिता बायस्कर, विद्या बनाफर, अश्विनी घुले, मधुराणी बन्सोड, डॉ. मीना सोसे, डॉ.पुजा फाळके, वनिता गावंडे, माधुरी चौधरी, अलका बोर्डे, कल्पना डोईफोडे, प्रणिता धोटे, रजनी चौखंडे

भोजन समिती: उज्चल विभुते , लक्ष्मण दारमोडे, नितिन वानखडे, सुनिल जुमळे, नितिन बाहकर, भागवत वाढोकार, राम पळसकर

संमेलन समिती सदस्य, मंदा पडवेकर, मनोज मेंडके, अभिजीत मुरूमकर, स्वप्नील इंगळे, धीरज चावरे, अजय इंगळे, किशोर पवार, वैभव इंगळे, प्रज्योत देशमुख, प्रशांत भोंडे, अशोक उघडे, डॉ. शांतीलाल चव्हाण, डॉ. स्वप्नील मानकर गजानन उगले, निखिल गावंडे, राजेश गिन्हे, हर्षा वाघमारे, ईश्वर मते, राम जाधव, मधुकर तराळे, मिलिंद हिवराळे, नितीन ठाकरे, प्रवीण बोपुलकर, रामदास गायधनी, राजेश काटोले, सागर लाहोळकर, संजय रोंगे, प्रा.अशोक सारडा, स्वप्रिल कुलकर्णी, जावेद शेख





कार्याध्यक्ष श्याम ठक (अध्यक्ष, अ.भा. वन्हाडी साहित्य मंच)



संपर्कः श्याम ठकः ९९७५७९२५२० निलेश कवडे : ९८२२३६७७०६



ज्येष्ठ सल्लागार प्रा.सदाशिव शेळके (अ.भा. वन्हाडी साहित्य मंच)



संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती व 'अखिल भारतीय वऱ्हाडी साहित्य मंच, अकोला' द्वारा संयक्त आयोजन

आवतन





अध्यक्ष : डॉ.प्रा.अनुराधा वऱ्हाडे सहभाग: विजय पाटील, प्रा.अशोक मानकर संचालन : निलेश देवकर, आभार: संजय नेमाडे

दुपारी १ ते २ वाजेलोग

सत्र तिसरे

कथाकथन

### जेवनखावन : दुपारी २ ते ३ वाजेलोग

### अध्यक्षः मीराताई ठाकरे

प्रमुख अतिथी: विजया मारोतकर (नागपूर), अरुण विघ्ने (वर्धा), सहभागः नितिन वरणकार (शेगांव), शिवलिंग काटेकर (अकोला), विजय सोसे (परतवाडा), संदिप धावडे (वर्धा), प्रमोद अंबडकार (नाशिक), विजय ढाले (आर्णी), डॉ. संघमिता खंडारे (दर्यापूर), रवींद्र दळवी (नाशिक), डॉ.लक्ष्मण उगले (खामगाव) डॉ. अंबादास घुले, डॉ. गजानन बनसोड, गजानन मानकर संचालनः प्रा.महादेव लुले, पवन नालट आभार: अंकुश इंगळे (मुंबई)



अध्यक्षः डॉ.प्रा.राजेश जयपूरकर प्र.कुलगुरू संमेलनाध्यक्षः नरेंद्र इंगळे, प्रमुख अतिथी: आबासाहेव कडू, दयाराम निंबोळकर, डॉ.ममता इंगोले, निलू मानकर, तेजस्वी बारब्दे,

संचालनः साधना काळबांडे आभारः उमेश थोरात

संध्याकाई ४:३० ते ५:३० वाजेलोग

समारोप

सत्र पाचवे

अध्यक्षः किशोर मुगल प्रमुख अतिथीः विजय बिन्दोड, डॉ.प्रा.सुनिल पखाले, भास्कर इंगळे सहभागः अजय इंगळे, अजय माटे, अजय इंगळे, जावेद शेख, पवन वसे, विनायक काळे, अंकुश इंगळे, श्रीराम वाघ, के. एस. इंगोले अँड. रजनी बावस्कर, सौ. अनुराधा धामोडे, प्रा.अशोक सारडा, सौ.जयश्री रोहनकर, युवराज टोपले, तुकाराम काटे, मधुराणी बन्सोड, उमेश थोरात, खुशाल गुल्हाणे, श्याम राऊत, मोहन जाधव, अजय राऊत, वर्षा बोध, दिनेश सातव,वीणा डिक्कर, दीपकराज खवशी, धीरज चावरे, संजय रोंघे, जितेश इंगळे, हर्षल साखरे, वैष्णवी सानप, शालिनी बेलसरे, अभिजीत इंगळे, प्राची अडलोक, डॉ. शांतीलाल चव्हाण, आशिष कांबळे, सारिका वनवे, प्रशांत भोंडे, अलका देशमुख, मधुकर तराळे, प्रशांत ढोले, सरिता बायस्कर, डॉ. पूजा फाळके, अंकुश चौरपगार संचालन: अनिकेत देशमुख, स्वग्नील इंगळे, आभार: संजय गावंडे



संध्याकाई ५:३० ते ७:३० वाजेलोग



तासमुद्रापार नेन्याचा. .!











संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती व 'अखिल भारतीय वऱ्हाडी साहित्य मंच, अकोला' द्वारा संयुक्त आयोजन

### तिसरे अखिल भारतीय वऱ्हाडी साहित्य संमेलन अमरावती-२०२०

मा. श्री /सौ.

अखिल भारतीय वऱ्हाडी साहित्य मंच, अकोला या आमच्या साहित्यिक मंचानं, वऱ्हाडी बोलीभाषेले जतन कऱ्याचं काम हाती घेतलं ! म्हनून बारोमास काईना काई उपक्रम राबोनं चालूच असते. या सालातई ता .०४ जानेवारी-२०२० शनवार रोजी स्व.उध्दव शेळके साहित्य नगरी, स्व.मनोहर तल्हार विचारपीठ हकश्राव्य सभागृह, मराठी विभाग परिसर, संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती. अधी तिसरे अ . भा . वऱ्हाडी साहित्य संमेलन होनार हाय! तुमाले या कारेक्रमात या लागते म्हनजे याच लागते! कोनताबी बायना चालनार नाई .

### कारेक्रमाचा आराधुरा

सत्र पयले उद्घाटन सत्र

सकावुन १० वा. ते दुपारी १२ वाजेलोग संमेलनाध्यक्ष : मा. नरेंद्र इंगळे (ज्येष्ठ वऱ्हाडी साहित्यिक) स्वागताध्यक्ष : मा.डॉ. मुरलीधर चांदेकर, कुलगुरू (संत गाडगे बाबा अमरावती विद्यापीठ, अमरावती) उदघाटक : मा. प्रा.मधुकर वाकोडे, मावळते अध्यक्ष : मा.प्रकाश पोहरे प्रमुख अतिथी: मा. शिरीषदादा धोत्रे, प्रमुख अतिथी: मा. संदिपजी भारंबे प्रमुख अतिथी: सौ. अनुराधा धामोडे

पयले संमेलनाध्यक्षः मा.पुष्पराज गावंडे, सल्लागारः प्रा.सदाशिव शेळके, मराठी विभागप्रमुखः डॉ.मोना चिमोटे, कार्याध्यक्षः श्याम ठक, संचालनः डॉ.प्रा.हेमंत खडके आभारः दयाराम निंबोळकर

> विषय : संस्कृती संवर्धनासाठी बोलीचे उपयोजन अध्यक्षा : डॉ. अलका गायकवाड सहभाग: प्रा.डॉ. रावसाहेब काळे, डॉ. मीनाताई गावंडे संचालन: निलेश कवडे, आभार: डॉ.मीना सोसे



संत गाडगे बाबा अमरावती विद्यापीठ पान क्रं. /N

क्र.संगाबाअवि/५१७/ **१९५** /२०१९ दिनांकः- २३/१२/२०१९

विषय :- 'मराठी शुद्धलेखन कार्यशाळे'चे संयुक्त विद्यमाने आयोजन करण्याबाबत.

संदर्भ :- शासन परिपत्रक क्र.मभापं-२०१९/प्र.क्र.९३/भाषा-२ दि. ०३.१२.२०१९. संत गाडगे बाबा अमरावती विद्यापीठाच्या पदव्युत्तर मराठी विभागाद्वारे दरवर्षी दिनांक ०१ जानेवारी ते १५ जानेवारी या कालावधीत मराठी भाषा संवर्धन पंधरवाडा साजरा केला जातो. याही वर्षी मराठी भाषा संवर्धन पंधरवाडा साजरा करण्यासंदर्भात शासन आदेश वरील संदर्भाकित पत्रान्वये प्राप्त झाले आहेत.

यावर्षी पदव्युत्तर मराठी विभाग, सं.गा.बा.अमरावती विद्यापीठ व मराठी विभाग शासकीय ज्ञान-विज्ञान संस्था, अमरावती यांच्या संयुक्त विद्यामाने दिनांक ७, ८ व ९ जानेवारी, २०२० रोजी **मराठी शुद्धलेखन कार्यशाळा** आयोजित करण्यासंदर्भात ठरविले आहे. या संदर्भात डॉ. केशव तुपे, विभागप्रमुख शासकीय ज्ञान-विज्ञान संस्था, अमरावती यांच्याशी चर्चा झाली असून त्यांनी संयुक्तपणे कार्यशाळा आयोजित करण्याची संमती दिली आहे. त्यांनी कार्यशाळेसाठी एकूण येणाऱ्या अंदाजे रुपये १,००,०००/- खर्चांपैकी अर्धा खर्च म्हणजेच ५०,०००/- रुपये खर्च उचलण्याची तयारी दर्शविली आहे आणि त्या संदर्भात लेखी पत्र मराठी विभागप्रमुख संत गाडगे बाबा अमरावती विद्यापीठाकडे सुपूर्त केले आहे.

त्या अनुषंगाने पद. मराठी विभागाला कार्यशाळेचे प्रमुख मार्गदर्शक डॉ. अरुण फडके यांचे मानधन व इतर अनुषंगिक खर्च वहन करावयाचा असून तो खालीलप्रमाणे राहील.

- मानधन
- २) बॅनर
- ३) हार व ग्रंथ भेट
- ४) छायाचित्रे
- ५) प्रमाणपत्र
- ६) इतर अनुषंगीक खर्च

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उपरोक्त खर्च रुपये ५०,०००/- हा विद्यापीठ अंदाजपत्रकीय तरतूद ४) साधारण निधी (जनरल फंड) अंतर्गत (H) बेस्ट प्रॅक्टिसेस इन युनिव्हर्सिटी (vi) मराठी भाषा संवर्धन उपक्रम या उपशीर्षांतर्गत करण्यात येईल.

सदर कार्यशाळा आयोजित करण्यासाठी प्रशासकीय व वित्तीय मान्यता मिळावी, ही विनंती.

मोनान्धिम मराठी विभागप्रमुख

83 12/10

मा. कुलग्र

संत गाडगे बाबा अमरावती विद्यापीठ पान क्रं. ol /N प्रशासकीय भाज्यता मिळाली असल्याने वित्तीय मान्यतेसाठी 21122. भोगानिभीटे H2181 ABH121,4 420 M/168 - विल व लेखा आधिकारी DR(Audit) 30/12/19 30/12/209 F-7100 Kuimp Juzou - नियमाने पालन कोपेमून अस्तावित स्रस्तावास २०१९ - २० मध्ये ( र. 50,000 - - भी नोंद घेपून विलीम F-8411 301219 A-27214 3 3. Bizzan - 31-20 Al 3111m 3/1/20 Corrigio brean Breadont - h M Br Stolm 4216 भूजासकीय व विल्तीय माज्यता मिछाछी आहे. उनगीम मिलाया क्रम सादर. મોનાન્વિમોટ - विला व लेखा अधिकारी





Registrar Melghat Tiger Reserve Head, Deptt. of Zoology, SGBAU	<b>COMPETITION (UG / PG Category)</b> organized on the eve of <b>Wildlife Week-201</b> Zoology, Sant Gadge Baba Amravati University Amravati, in collaboration with Melgh Amravati and Center for Animal Rescue and Study, Amravati on 6 <sup>th</sup> October 2017.	This is to certify thatMr./Ms./ Mrsof	<section-header></section-header>	Sant Gadge Baba Amravati University NAAC Accredited with "A" Grade Wildlife Week-2017
rkar Shri.S. M. Deshmukh GBAU CARS, Amravati	<b>Neek-2017</b> by Department of 1 with Melghat Tiger Reserve, r 2017.	prize in SEMINAR/POSTER		rsity

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Wild Life Week 2015- Poster Presentation Competition Sant Gadge Baba Amravati University, Amravati

Poster



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Wild Life Week 2015 Sant Gadge Baba Amravati University, Amravati



SANT GADGE BABA AMRAVATI UNIVERSITY, AMRAVATI DEPARTMENT OF ZOOLOGY

WILDLLIFE WEEK – 2017

# POSTERPRESENTATION

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### SANT GADGE BABA AMRAVATI UNIVERSITY, AMRAVATI DEPARTMENT OF ZOOLOGY WILDLLIFE WEEK – 2017

# POSTERPRESENTATION

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## SANT GADGE BABA AMRAVATI UNIVERSITY, AMRAVATI DEPARTMENT OF ZOOLOGY

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WILDLLIFE WEEK – 2017

# SEMINAR COMPETITION

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	Viney Machard	& Sakghi Washeni	Shiven Your Sky	Selliam Jacker	Genzane Jolli	Sweety Sontakke	Akshay Ramteke	Akshay Sawai	Megha Dharawat	Zahir Abass	Nidhi Kale		. Name of Student
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## Rajiv Gandhi Science and Technology Commission Government of Maharashtra

## No.RGSTC/File-2011/DPP-166/CR-42 / 169

Apeejay House, 3<sup>rd</sup> Floor, Dinshaw Vaccha Road, Near K.C. College, Churchgate, Mumbai-400 020. Tel No:-022-22024755/22024711 Email :-rgstcmaha@rediffmail.com Date:- 24<sup>th</sup> March, 2021.

## To whomsoever it may concern

For advancement, propagation and promotion of applications of Science and Technology for development and to tackle various problems facing the society, Rajiv Gandhi Science and Technology Commission, Government of Maharashtra, Mumbai have evolved a scheme titled, "Assistance for Science and Technology Applications". Under this scheme the Commission had sanctioned a project titled, "An inexpensive method for the determination of oxalate in urine and assessment of hyperoxaluria by nano- based paper strip technique", vide sanction Order No. RGSTC/File-2016/DPP-166/CR-42, dated, 27<sup>th</sup> March, 2017 to Sant Gadge Baba Amravati University (SGBAU), Amravati.

The following faculty members had worked on the said project,

- (1) Dr. Anita Patil, Professor Dept. of Biotechnology Principal Investigator
- (2) Dr. Gajanan Muley, Associate Professor, Department of Physics -Co-Investigator
- (3) Dr. Hariprasad Paikrao, Assistant Professor, Dept. of Forensic Biology, Govt. Institute of Science Aurangabad-Co Investigator

Anna

(A. S. Manekar) Member Secretary



## भारतीय चरागाह एवं चारा अनुसंधान संस्थान INDIAN GRASSLAND AND FODDER RESEARCH INSTITUTE



निकट पहुज बॉध, ग्वालियर रोड, झॉसी (उ.प्र.) – 284003 भारत Near Pahuj Dam, Gwalior Road, JHANSI – 284 003 (U.P.) INDIA Phone : 0510-2730666 Fax : 0510-2730833 email : igfri.director@gmail.com

Dr. Manoj Kumar Srivastava Principal Scientist (Biochemistry-Plant sciences)

03.04.2018

To The Head, Department of Biotechnology Sant Gadge Baba Amravati University, Amravati, Maharastra.

Subject: Request for characterization of nano silver (AgNPs) synthesized using fodder grasses.

Dear Sir,

This is to state that we are working of biosynthesis of silver nano particles using extract of fodder grasses. As your Department has excellence in the facility to characterize nano particles, we request your support in characterization of nanoparticles synthesized here in your lab. For this purpose Mr. Anand Soni, M. Sc. (Biotechnology), who is working for his dissertation work under my supervision will be at your University from **14.04.2018 to 23.04.2018**. In this regard kindly permit him to work under the able supervision of Dr. Aniket Gade, Associate Professor at your Department.

This will be an academic collaboration between both the Institutes.

Thanking you,

With regards,

(Manoj Kumar Srivastava) Principal Scientist (Biochemistry- Plant Sciences) CI Division

srimanoj2009@gmail.com

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